

# 1 Magnetospheric chaos and dynamical complexity response during storm time disturbance

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## 6 Abstract

7 In this study, we examine the magnetospheric chaos and dynamical complexity response to the  
8 disturbance storm time ( $D_{st}$ ) and solar wind electric field ( $VB_s$ ) during different categories of  
9 geomagnetic storm (minor, moderate and major geomagnetic storm). The time series data of the  
10  $D_{st}$  and  $VB_s$  are analyzed for the period of nine years using nonlinear dynamics tools (Maximal  
11 Lyapunov Exponent, MLE, Approximate Entropy, ApEn and Delay Vector Variance, DVV). We  
12 found a significant trend between each nonlinear parameter and the categories of geomagnetic  
13 storm. The MLE and ApEn values of the  $D_{st}$  indicate that chaotic and dynamical complexity  
14 responses are high during minor geomagnetic storms, reduce at moderate geomagnetic storms and  
15 decline further during major geomagnetic storms. However, the MLE and ApEn values obtained  
16 from  $VB_s$  indicate that chaotic and dynamical complexity responses are high with no significant  
17 difference between the periods that are associated with minor, moderate and major geomagnetic  
18 storms. The test for nonlinearity in the  $D_{st}$  time series during major geomagnetic storm reveals the  
19 strongest nonlinearity features. Based on these findings, the dynamical features obtained in the  
20  $VB_s$  as input and  $D_{st}$  as output of the magnetospheric system suggest that the magnetospheric  
21 dynamics is nonlinear and the solar wind dynamics is consistently stochastic in nature.

22 **Keywords:**  $D_{st}$  signals, Solar wind electric field ( $VB_s$ ) signals, Geomagnetic storm, Chaotic  
23 behaviour, Dynamical complexity, Nonlinearity.

24

## 25 **1.0 Introduction**

26 The response of chaos and dynamical complexity behaviour with respect to magnetospheric  
27 dynamics varies (Tsurutani et al., 1990). This is due to changes in the interplanetary electric fields  
28 imposed on the magnetopause and those penetrating the inner magnetosphere and sustaining  
29 convection thereby initiating geomagnetic storm (Dungey, 1961; Pavlos et al., 1992). A prolonged  
30 southward turning of interplanetary magnetic field (IMF,  $B_z$ ), which indicates that solar wind-  
31 magnetosphere coupling is in-progress was confirmed on many occasions for which such  
32 geomagnetic storm was driven by Corotating Interaction Regions (CIRs), or by the sheath  
33 preceding an interplanetary coronal mass ejection (ICME) or by a combination of the sheath and  
34 an ICME magnetic cloud (Gonzalez and Tsurutani, 1987; Tsurutani and Gonzalez, 1987;  
35 Tsurutani et al., 1988; Cowley, 1995; Tsutomu, 2002; Yurchyshyn et al., 2004; Kozyra et al., 2006;  
36 Echer et al., 2008; Meng et al., 2019; Tsurutani et al., 2020). The sporadic magnetic reconnection  
37 between the southward component of the Alfvén waves and the earth's magnetopause leads to  
38 isolated substorms/convection events such as the high intensity long-duration continuous AE  
39 activity (HILDCAA) which are shown to last from days to weeks (Akasofu, 1964; Tsurutani et al.,  
40 1972; Meng et al., 1973; Tsurutani and Gonzalez, 1987; Hajra et al., 2013; Liou et al., 2013;  
41 Mendes et al., 2017; Hajra and Tsurutani, 2018; Tsurutani and Hajra, 2021). Notably, the  
42 introduction of Disturbance Storm Time ( $D_{st}$ ) index (Sugiura, 1964; Sugiura and Kamei, 1991)  
43 unveiled the quantitative measure of the total energy of the ring current particles. Therefore, the  
44  $D_{st}$  index remains one of the most popular global indicators that can precisely reveal the severity  
45 of a geomagnetic storm (Dessler and Parker, 1959).

46 The  $D_{st}$  fluctuations exhibit different signatures for different categories of geomagnetic storm.  
47 Ordinarily, one can easily anticipate that fluctuations in a  $D_{st}$  signal appear chaotic and complex.

48 These may arise from the changes in the interplanetary electric fields driven by the solar wind-  
49 magnetospheric coupling processes. At different categories of geomagnetic storm, fluctuations in  
50 the  $D_{st}$  signals differ (Oludehinwa et al., 2018). One obvious reason is that as the intensity of the  
51 geomagnetic storm increases, the fluctuation behaviour in the  $D_{st}$  signal becomes more complex  
52 and nonlinear in nature. It has been established that the electrodynamic response of the  
53 magnetosphere to solar wind drivers are non-autonomous in nature (Price and Prichard, 1993;  
54 Price et al., 1994; Johnson and Wings, 2005). Therefore, the chaotic analysis of the  
55 magnetospheric time series must be related to the concept of input-output dynamical process  
56 (Russell et al., 1974; Burton et al., 1975; Gonzalez et al., 1989; Gonzalez et al., 1994).  
57 Consequently, it is necessary to examine the chaotic behaviour of the solar wind electric field  
58 ( $VB_s$ ) as input signals and the magnetospheric activity index ( $D_{st}$ ) as output during different  
59 categories of geomagnetic storms.

60 Several works have been presented on the chaotic and dynamical complexity behaviour of the  
61 magnetospheric dynamics based on autonomous concept, i.e using the time series data of  
62 magnetospheric activity alone such as auroral electrojet (AE), Amplitude Lower (AL) and  $D_{st}$   
63 index (Vassiliadis et al., 1990; Baker and Klimas, 1990; Vassiliadis et al., 1991; Shan et al., 1991;  
64 Pavlos et al., 1994; Klimas et al., 1996; Valdivia et al., 2005; Mendes et al., 2017; Consolini,  
65 2018). They found evidence of low-dimensional chaos in the magnetospheric dynamics. For  
66 instance, the report by Vassiliadis et al. (1991) shows that the computation of Lyapunov exponent  
67 for AL index time series gives a positive value of Lyapunov exponent indicating the presence of  
68 chaos in the magnetospheric dynamics. Unnikrishnan, (2008) studied the deterministic chaotic  
69 behaviour in the magnetospheric dynamics under various physical conditions using AE index time  
70 series and found that the seasonal mean value of Lyapunov exponent in winter season during quiet

71 periods ( $0.7 \pm 0.11 \text{ min}^{-1}$ ) is higher than that of the stormy periods ( $0.36 \pm 0.09 \text{ min}^{-1}$ ).  
72 Balasis et al. (2006) examined the magnetospheric dynamics in the  $D_{st}$  index time series from  
73 pre-magnetic storm to magnetic storm period using fractal dynamics. They found that the transition  
74 from anti-persistent to persistent behaviour indicates that the occurrence of an intense geomagnetic  
75 storm is imminent. Balasis et al. (2009) further reveal the dynamical complexity behaviour in the  
76 magnetospheric dynamics using various entropy measures. They reported a significant decrease in  
77 dynamical complexity and an accession of persistency in the  $D_{st}$  time series as the magnetic storm  
78 approaches. Recently, Oludehinwa et al. (2018) examined the nonlinearity effects in  $D_{st}$  signals  
79 during minor, moderate and major geomagnetic storm using recurrence plots and recurrence  
80 quantification analyses. They found that the dynamics of the  $D_{st}$  signal is stochastic during minor  
81 geomagnetic storm periods and deterministic as the geomagnetic storm increases.

82 Also, studies describing the solar wind and magnetosphere as a non-autonomous system have been  
83 extensively investigated. Price et al. (1994) examine the nonlinear input-output analysis of AL  
84 index and different combinations of interplanetary magnetic field (IMF) with solar wind  
85 parameters as input functions. They found that only a few of the input combinations show any  
86 evidence whatsoever for nonlinear coupling between the input and output for the interval  
87 investigated. Pavlos et al. (1999) presented further evidence of magnetospheric chaos. They  
88 compared the observational behaviour of the magnetospheric system with the results obtained by  
89 analyzing different types of stochastic and deterministic input-output systems and asserted that a  
90 low dimensional chaos is evident in magnetospheric dynamics. Devi et al. (2013) studied the  
91 magnetospheric dynamics using AL index and the southward component of IMF( $B_z$ ). They  
92 observed that the magnetosphere and turbulent solar wind have values corresponding to nonlinear  
93 dynamical system with chaotic behaviour. The modeling and forecasting approach have been

94 applied to magnetospheric time series using nonlinear models (Valdivia et al., 1996; Vassiliadis et  
95 al., 1999; Vassiliadis, 2006; Balikhin et al., 2010). These efforts have improved our understanding  
96 that the concept of nonlinear dynamics can reveal some hidden dynamical information in the  
97 observational time series. In addition to these nonlinear effects in  $D_{st}$  signals, a measure of the  
98 exponential divergence and convergence within the trajectories of a phase space known as  
99 Maximal Lyapunov Exponent (MLE), which has the potential to depict the chaotic behavior in the  
100  $D_{st}$  and  $VB_s$  time series during a minor, moderate and major geomagnetic storm have not been  
101 investigated. In addition, to the best of our knowledge, computation of Approximate Entropy  
102 (ApEn) that depicts the dynamical complexity behaviour during different categories of  
103 geomagnetic storm has not been reported in the literature. The test for nonlinearity through delay  
104 vector variance (DVV) analysis that reveals the nonlinearity features in  $D_{st}$  and  $VB_s$  time series  
105 during minor, moderate and major geomagnetic storms is not well known. It is worth to note that  
106 understanding the dynamical characteristics in the  $D_{st}$  and  $VB_s$  signals at different categories of  
107 geomagnetic storms will provide useful diagnostic information to different conditions of space  
108 weather phenomenon. Consequently, this study attempts to carry out comprehensive numerical  
109 analyses to unfold the chaotic and dynamical complexity behaviour in the  $D_{st}$  and  $VB_s$  signals  
110 during minor, moderate and major geomagnetic storm. In section 2, our methods of data  
111 acquisition are described. Also, the nonlinear analysis that we employed in this investigation are  
112 detailed. In section 3, we unveiled our results and engage the discussion of results in section 4.

## 113 **2.0 Description of the Data and Nonlinear Dynamics**

114 The  $D_{st}$  index is derived by measurements from ground-based magnetic stations at low-latitudes  
115 observatories around the world and depicts mainly the variation of the ring current, as well as the  
116 Chapman-Ferraro Magnetopause currents, and tail currents to a lesser extent (Sugiura, 1964;

117 Feldstein et al., 2005; Feldstein et al., 2006; Love and Gannon, 2009). Due to its global nature,  $D_{st}$   
118 time series provides a measure of how intense a geomagnetic storm was (Dessel and Parker, 1959).  
119 In this study, we considered  $D_{st}$  data for the period of nine years from January to December  
120 between 2008 and 2016 which were downloaded from the World Data Centre for Geomagnetism,  
121 Kyoto, Japan (<http://wdc.kugi-kyoto-u.ac.jp/Dstae/index.html>). We use the classification of  
122 geomagnetic storms as proposed by Gonzalez et al. (1994) such that  $D_{st}$  index value in the ranges  
123  $0 \leq D_{st} \leq -50nT$ ,  $-50nT \leq D_{st} \leq -100nT$ ,  $-100nT \leq D_{st} \leq -250nT$  are classified as  
124 minor, moderate and major geomagnetic storms respectively and each time series is being  
125 classified based on its minimum  $D_{st}$  value. The solar wind electric field ( $VB_s$ ) data are archived  
126 from the National Aeronautics and Space Administration, Space Physics Facility  
127 (<http://omniweb.gsfc.nasa.gov>). The sampling time of  $D_{st}$  and  $VB_s$  time series data was 1-hour. It  
128 is well known that the dynamics of the solar wind contribute to the driving of the magnetosphere  
129 (Burton et al. 1975). Furthermore, we took the solar wind electric field ( $VB_s$ ) as the input signal  
130 (Price and Prichard, 1993; Price et al., 1994). The  $VB_s$  was categorized according to the periods  
131 of minor, moderate and major geomagnetic storm. Then, the  $D_{st}$  and  $VB_s$  time series were  
132 subjected to a variety of nonlinear analytical tools explained as follow:

### 133 **2.1 Phase Space Reconstruction and Observational time series**

134 An observational time series can be defined as a sequence of scalar measurements of some  
135 quantity, which is a function of the current state of the system taken at multiples of a fixed sampling  
136 time. In nonlinear dynamics, the first step in analyzing an observational time series data is to  
137 reconstruct an appropriate state space of the system. Takens (1981) and Mane (1981) stated that  
138 one time series or a few simultaneous time series are converted to a sequence of vectors. This

139 reconstructed phase space has all the dynamical characteristic of the real phase space provided the  
 140 time delay and embedding dimension are properly specified.

$$141 \quad X(t) = [x(t), x(t + \tau), x(t + 2\tau), \dots, x(t + (m - 1)\tau)]^T \quad (1)$$

142 Where  $X(t)$  is the reconstructed phase space,  $x(t)$  is the original time series data,  $\tau$  is the time  
 143 delay and  $m$  is the embedding dimension. An appropriate choice of  $\tau$  and  $m$  are needed for the  
 144 reconstruction of phase space which is determined by average mutual information and false nearest  
 145 neighbor, respectively.

## 146 **2.2 Average Mutual Information (AMI)**

147 The method of Average Mutual Information (AMI) is one of the nonlinear techniques used to  
 148 determine the optimal time delay ( $\tau$ ) required for phase space reconstruction in observational time  
 149 series. The time delay mutual information was proposed by Fraser and Swinney, (1986) instead of  
 150 an autocorrelation function. This method takes into account nonlinear correlations within the time  
 151 series data. It measures how much information can be predicted about one time series point, given  
 152 full information about the other. For instance, the mutual information between  $x_i$  and  $x_{(i+\tau)}$   
 153 quantifies the information in state  $x_{(i+\tau)}$  under the assumption that information at the state  $x_i$  is  
 154 known. The AMI for a time series,  $x(t_i)$ ,  $i = 1, 2, \dots, N$  is calculated as:

$$155 \quad I(T) = \sum_{x(t_i), x(t_i+T)} P(x(t_i), x(t_i + T)) \times \log_2 \left[ \frac{P(x(t_i), x(t_i+T))}{P(x(t_i)) P(x(t_i+T))} \right] \quad (2)$$

156 where  $x(t_i)$  is the  $i$ th element of the time series,  $T = k\Delta t$  ( $k = 1, 2, \dots, k_{max}$ ),  $P(x(t_i))$  is the  
 157 probability density at  $x(t_i)$ ,  $P(x(t_i), x(t_i + T))$  is the joint probability density at the pair  
 158  $x(t_i), x(t_i + T)$ . The time delay ( $\tau$ ) of the first minimum of AMI is chosen as optimal time delay  
 159 (Fraser and Swinney, 1986). Therefore, the AMI was applied to the  $D_{st}$  and  $VB_s$  time series and

160 the plot of AMI versus time delay is shown in Figure (3). We notice that the AMI showed the first  
 161 local minimum at roughly  $\tau = 15hr$ . Furthermore, the values of  $\tau$  near this value of  $\sim 15hr$   
 162 maintain constancy for both  $VB_S$  and  $D_{st}$ . In the analysis  $\tau = 15hr$  was used as the optimal time  
 163 delay for the computation of maximal Lyapunov exponent.

### 164 **2.3 False Nearest Neighbor (FNN)**

165 In determining the optimal choice of embedding dimension( $m$ ), the false nearest neighbor method  
 166 was used in the study. The method was suggested by Kennel et al. (1992). The concept is based  
 167 on how the number of neighbors of a point along a signal trajectory changes with increasing  
 168 embedding dimension. With increasing embedding dimension, the false neighbor will no longer  
 169 be neighbors, therefore by examining how the number of neighbors changes as a function of  
 170 dimension, an appropriate embedding dimension can be determined. For instance, suppose we  
 171 have a one-dimensional time series. We can construct a time series  $y(t)$  of  $D$ -dimensional points  
 172 from the original one-dimensional time series  $x(t)$  as follows:

$$173 \quad y(t) = (x(t), x(t + \tau), \dots, x(t + (D - 1)\tau)) \quad (3)$$

174 where  $\tau$  and  $D$  are time delay and embedding dimension. Using the formular from Kennel et al.  
 175 (1992); Wallot and Monster, (2018), if we have a  $D$ -dimensional phase space and denote the  $r$ th  
 176 nearest neighbor of a coordinate vector  $y(t)$  by  $y^{(r)}(t)$ , then the square of the Euclidean distance  
 177 between  $y(t)$  and the  $r$ th nearest neighbor is:

$$178 \quad R_D^2(t, r) = \sum_{k=0}^{D-1} [x(t + k\tau) - x^{(r)}(t + k\tau)]^2 \quad (4)$$



179 Now applying the logic outlined above, we can go from a  $D$ -dimensional phase space to  $(D + 1)$   
180 dimensional phase space by time-delay embedding, adding a new coordinate to  $y(t)$ , and ask what  
181 is the squared distance between  $y(t)$  and the same  $r$ th nearest neighbor:

$$182 \quad R_{D+1}^2(t, r) = R_D^2(t, r) + [x(t + D\tau) - x^{(r)}(t + D\tau)]^2 \quad (5)$$

183 As explained above, if the one-dimensional time series is already properly embedded in  $D$   
184 dimensions, then the distance  $R$  between  $y(t)$  and the  $r$ th nearest neighbor should not change  
185 appreciably by some distance criterion  $R_{tol}$  (i.e.  $R < R_{tol}$ ). Moreover, the distance of the nearest  
186 neighbor when embedded into the next higher dimension relative to the size of the attractor should  
187 be less than some criterion  $A_{tol}$  (i.e.  $R_{D+1} < A_{tol}$ ). Doing this for the nearest neighbor of each  
188 coordinate will result on many false nearest neighbors when embedding is insufficient or in few  
189 (or no) false neighbors when embedding is sufficient. In the analysis, the FNN was applied to the  
190  $D_{st}$  and  $VB_s$  time series to detect the optimal value of embedding dimension ( $m$ ). Figure (4) shows  
191 a sample plot of the percentage of false nearest neighbor against embedding dimension in one of  
192 the months under investigation (other months show similar results, thus for brevity we depict only  
193 one of the results). We notice that the false nearest neighbor attains its minimum value at  $m \geq 5$   
194 indicating that embedding dimension ( $m$ ) from  $m \geq 5$  are optimal values. Therefore,  $m = 5$  was  
195 used for the computation of maximal Lyapunov exponent.

## 196 **2.4 Maximal Lyapunov Exponent (MLE)**

197 The Maximal Lyapunov Exponent (MLE) is one of the most popular nonlinear dynamics tool used  
198 for detecting chaotic behaviour in a time series data. It describes how small changes in the state of  
199 a system grow at an exponential rate and eventually dominate the behaviour. An important  
200 indication of chaotic behavior of a dissipative deterministic system is the existence of a positive

201 Lyapunov Exponent. A positive MLE signifies divergence of trajectories in one direction or  
 202 expansion of an initial volume in this direction. On the other hand, a negative MLE exponent  
 203 implies convergence of trajectories or contraction of volume along another direction. The  
 204 algorithm proposed by Wolf et al. (1985) for estimating MLE is employed to compute the chaotic  
 205 behavior of the  $D_{st}$  and  $VB_S$  time series at minor, moderate and major geomagnetic storm. Other  
 206 methods of determining MLE includes Rosenstein's method, Kantz's method and so on. In this  
 207 study, the MLE at minor, moderate and major geomagnetic storms periods was computed with  
 208  $m = 5$  and  $\tau = 15hr$  as shown in figures (5 & 6-bar plots) for  $D_{st}$  and  $VB_S$ . The calculation of  
 209 MLE is explained as follows: given a sequence of vector  $x(t)$ , an  $m$ -dimensional phase space is  
 210 formed from the observational time series through embedding theorem as

$$211 \quad \{x(t), x(t + \tau), \dots, x(t + (m - 1)\tau)\} \quad (6)$$

212 where  $m$  and  $\tau$  are as defined earlier, after reconstructing the observational time series, the  
 213 algorithm locates the nearest neighbor (in Euclidean sense) to the initial point  $\{x(t_0), \dots, x(t_0 +$   
 214  $(m - 1)\tau\}$  and denote the distance between these two points  $L(t_0)$ . At a later point  $t_1$ , the initial  
 215 length will have evolved to length  $L'(t_1)$ . Then the MLE is calculated as:

$$216 \quad \lambda = \frac{1}{t_M - t_0} \sum_{k=1}^M \log_2 \frac{L'(t_k)}{L(t_{k-1})} \quad (7)$$

217  $M$  is the total number of replacement steps. We look for a new data point that satisfies two criteria  
 218 reasonably well: its separation,  $L(t_1)$ , from the evolved fiducial point is small. If an adequate  
 219 replacement point cannot be found, we retain the points that were being used. This procedure is  
 220 repeated until the fiducial trajectory has traversed the entire data.

221

## 222 2.5 Approximate Entropy (ApEn)

223 Approximate Entropy (ApEn) is one of the nonlinear dynamics tools that measure the dynamical  
224 complexity in observational time series. The concept was proposed by Pincus, (1991) which  
225 provides a generalized measure of regularity, such that it accounts for the logarithm likelihood in  
226 the observational time series. For instance, a dataset of length,  $N$ , that repeat itself for  $m$  points  
227 within a boundary will again repeat itself for  $m + 1$  points. Because of its computational  
228 advantage, ApEn has been widely used in many areas of disciplines to study dynamical complexity  
229 (Pincus and Kalman (2004); Pincus and Goldberger (1994); McKinley et al. (2011); Kannathan et  
230 al. (2005); Balasis et al. (2009); Shujuan and Weidong, (2010); Moore and Marchant (2017)). The  
231 ApEn is computed using the formula below:

$$232 \quad ApEn(m, r, N) = \frac{1}{N-m+1} \sum_{i=1}^{N-m+1} \log C_i^m(r) - \frac{1}{N-m} \sum_{i=1}^{N-m} \log C_i^m(r) \quad (8)$$

233 where  $C_i^m(r) = \frac{1}{N-m+1} \sum_{j=1}^{N-m+1} \Theta(r - \|x_i - x_j\|)$  is the correlation integral,  $m$  is the embedding  
234 dimension and  $r$  is the tolerance. To compute the ApEn for the  $D_{st}$  and  $VB_s$  time series classified  
235 as minor, moderate and major geomagnetic storm from 2008 to 2016, we choose ( $m = 3, \tau =$   
236  $1hr$ ). We refer the works of Pincus, (1991); Kannathal et al. (2005); and Balasis et al. (2009) to  
237 interested readers where all the computational steps regarding ApEn were explained in details.  
238 Figures (5 & 6) depict the stem plot of ApEn for  $D_{st}$  and ( $VB_s$ ) from 2008 to 2016.

## 239 2.6 Delay Vector Variance (DVV) analysis

240 The Delay Vector Variance (DVV) is a unified approach in analyzing and testing for nonlinearity  
241 in a time series (Gautama et al., 2004; Mandic et al., 2007). The basic idea of the DVV is that, if  
242 two delay vectors of a predictable signal are close to each other in terms of the Euclidean distance,

243 they should have similar target. For instance, when a time delay ( $\tau$ ) is embedded into a time series  
 244  $x(k)$ ,  $k = 1, 2, \dots, N$ , then a reconstructed phase space vector is formed which represents a set of  
 245 delay vectors (DVs) of a given dimension.

$$246 \quad X(k) = [X_{k-m\tau}, \dots, X_{k-\tau}]^T \quad (9)$$

247 Reconstructing the phase space, a set ( $\lambda_k$ ) is generated by grouping those DVs that are with a  
 248 certain Euclidean distance to DVs ( $X(k)$ ). For a given embedding dimension ( $m$ ), a measure of  
 249 unpredictability  $\sigma^{*2}$  is computed over all pairwise Euclidean distance between delay vector as

$$250 \quad d(i, j) = \|x(i) - x(j)\| \quad (i \neq j) \quad (10)$$

251 Then, sets  $\lambda_k(r_d)$  are generated as the sets which consist of all delay vectors that lie closer to  $x(k)$   
 252 than a certain distance  $r_d$ .

$$253 \quad \lambda_k(r_d) = \{x(i) \| x(k) - x(i)\| \leq r_d\} \quad (11)$$

254 For every set  $\lambda_k(r_d)$ , the variance of the corresponding target  $\sigma^{*2}(r_d)$  is

$$255 \quad \sigma^{*2}(r_d) = \frac{\frac{1}{N} \sum_{k=1}^N \sigma_k^2(r_d)}{\sigma_k} \quad (12)$$

256 where  $\sigma^{*2}(r_d)$  is target variance against the standardized distance indicating that Euclidean  
 257 distance will be varied in a manner standardized with respect to the distribution of pairwise  
 258 distance between DVs. Iterative Amplitude Adjusted Fourier Transform (IAAFT) method is used  
 259 to generate the surrogate time series (Kugiumtzis, 1999). If the surrogate time series yields DV  
 260 plots similar to the original time series and the scattered plot coincides with the bisector line, then  
 261 the original time series can be regarded as linear (Theiler et al., 1992; Gautama et al., 2004; Imitaz,  
 262 2010; Jaksic et al., 2016). On the other hand, if the surrogate time series yields DV plot that is not

263 similar to that of the original time series, then the deviation from the bisector lines indicates  
264 nonlinearity. The deviation from the bisector lines grows as a result of the degree of nonlinearity  
265 in the observational time series.

$$266 \quad t^{D_{VV}} = \sqrt{\langle (\sigma^{*2}(r_d) - \frac{\sum_{i=1}^N \sigma_{s,i}^{*2}}{N_s}) \rangle} \quad (13)$$

267 where  $\sigma_{s,i}^{*2}(r_d)$  is the target variance at the span  $r_d$  for the  $i^{th}$  surrogate. To carry out the test for  
268 nonlinearity in the  $D_{st}$  signals,  $m = 3$  and  $n_d = 3$ , the number of reference DVs=200, and number  
269 of surrogate,  $N_s = 25$  was used in all the analysis. Then we examined the nonlinearity response at  
270 minor, moderate and major geomagnetic storm.

### 271 3.0 Results

272 In this study,  $D_{st}$  and  $VB_s$  time series from January to December were analyzed for the period of  
273 nine years (2008 to 2016) to examine the chaotic and dynamical complexity response in the  
274 magnetospheric dynamics during minor, moderate and major geomagnetic storms. Figures (1) &  
275 (2), display the samples of fluctuation signatures of  $D_{st}$  and  $VB_s$  signals classified as (a): minor,  
276 (b): moderate and (c): major geomagnetic storms. The plot of Average Mutual information against  
277 time delay ( $\tau$ ) shown in Figure (3) depicts that the first local minimum of the AMI function was  
278 found to be roughly at  $\tau = 15$ hr. Furthermore, we notice that the values of  $\tau$  near this value of  
279 ( $\sim 15$ hr) maintain constancy for both  $VB_s$  and  $D_{st}$ . Also, in Figure (4), we display the plot of the  
280 percentage of false nearest neighbor against embedding dimension ( $m$ ). It is obvious that a  
281 decrease in false nearest neighbor when increasing the embedding dimension drop steeply to zero  
282 at the optimal dimension ( $m = 5$ ), thereafter the false neighbors stabilizes at that  $m = 5$  for  $VB_s$   
283 and  $D_{st}$ . Therefore,  $m = 5$  and  $\tau = 15$ hr was used for the computation of MLE at different  
284 categories of geomagnetic storm, while  $m = 3$  and  $\tau = 1$ hr are applied for the computation of  
285 ApEn values.

286 The results of MLE (bar plot) and ApEn (stem plot) for  $D_{st}$  at minor, moderate and major  
287 geomagnetic storms are shown in Figure 5. During minor geomagnetic storms, we notice that the  
288 value of MLE ranges between 0.07 and 0.14 for most of the months classified as minor  
289 geomagnetic storm. Similarly, the ApEn (stem plot) ranges between 0.59 and 0.83. It is obvious  
290 that strong chaotic behaviour with high dynamical complexity are associated with minor  
291 geomagnetic storms. During moderate geomagnetic storms, (see b part of Figure 5), we observe a  
292 reduction in MLE values (0.04 – 0.07) compared to minor geomagnetic storm periods. Within  
293 the observed values of MLE during moderate geomagnetic storms, we found a slight rise of MLE

294 in the following months (Mar 2008), (Apr 2011), (Jan 2012, Feb 2012, Apr 2012), (Jul 2015, Aug  
295 2015, Sept 2015, Oct 2015, Nov 2015) and (Nov 2016). Also, the ApEn revealed a reduction in  
296 values between 0.44 and 0.57 during moderate geomagnetic storms. The lowest values of ApEn  
297 were noticed in the following months: May 2010, Mar 2011, and Jan 2016. During major  
298 geomagnetic storms as shown in Figure 5, the minimum and maximum value of MLE is  
299 respectively 0.03 and 0.04 implying a very strong reduction of chaotic behaviour compared with  
300 minor and moderate geomagnetic storms. The lowest values of MLE were found in the months of  
301 Jul 2012, Jun 2013 and Mar 2015. Interestingly, further reduction in ApEn value (0.29 – 0.40)  
302 was as well noticed during this period. Thus, during major geomagnetic storms, chaotic behaviour  
303 and dynamical complexity subside significantly.

304 We display in Figure 6, the results of MLE and ApEn computation for the  $VB_s$  which has been  
305 categorized according to the periods of minor, moderate and major geomagnetic storms. The  
306 values of MLE (bar plot) were between 0.06 and 0.20 for  $VB_s$ . The result obtained indicate strong  
307 chaotic behaviour with no significant difference in chaoticity during minor, moderate and major  
308 geomagnetic storm. Similarly, the results obtained from computation of ApEn (stem plot) for  $VB_s$   
309 depict a minimum value of 0.60 and peak value of 0.87 as shown in Figure 6. The ApEn values of  
310  $VB_s$  indicates high dynamical complexity response with no significant difference during the  
311 periods of the three categories of geomagnetic storm investigated.

312 The test for nonlinearity in the  $D_{st}$  signals during minor, moderate and major geomagnetic storms  
313 was analyzed through the DVV analysis. Shown in Figure 9 is the DVV plot and DVV scatter plot  
314 during minor geomagnetic storm for January 2009 and January 2014. We found that the DVV  
315 plots during minor geomagnetic storms reveals a slight separation between the original and  
316 surrogate data. Also, the DVV scatter plots shows a slight deviation from the bisector line between

317 the original and surrogate data which implies nonlinearity. Also, during moderate geomagnetic  
318 storms, we notice that the DVV plot depicts a wide separation between the original and the  
319 surrogate data. Also, a large deviation from the bisector line between the original and the surrogate  
320 data was also noticed in the DVV scatter plot as shown in Figure (8) thus indicating nonlinearity.  
321 In Figure (9), we display samples of DVV plot and DVV scatter plot during major geomagnetic  
322 storm for Oct 2011 and Dec 2015. The original and the surrogate data showed a very large  
323 separation in the DVV plot during major geomagnetic storm. While the DVV scatter plot depict  
324 the greatest deviation from the bisector line between the original and the surrogate data which is  
325 also an indication of nonlinearity. The DVV analysis of the  $VB_s$  time series during minor, moderate  
326 and major geomagnetic storms shown in Figures (10-12) revealed a separation between the original  
327 and surrogate data with no significant difference between the periods of minor, moderate and major  
328 geomagnetic storm.

## 329 **4.0 Discussion of Results**

### 330 **4.1 The chaotic and dynamical complexity response in $D_{st}$ at minor, moderate and major** 331 **geomagnetic storms**

332 Our result shows that the values of MLE for  $D_{st}$  during minor geomagnetic storm are higher,  
333 indicating significant chaotic response during minor geomagnetic stormy periods (bar plot, Figure  
334 5). This increase in chaotic behaviour for  $D_{st}$  signals during minor geomagnetic storms may be as  
335 a result of asymmetry features in the longitudinal distribution of solar source region for the  
336 Corotating Interaction Regions (CIR) signatures responsible for the development of geomagnetic  
337 storms (Turner et al. 2006; Kozyra et al. 2006). CIR generated magnetic storms are generally  
338 weaker than ICME/MC generated storms (Gonzalez et al., 1994; Tsurutani et al., 1995; Feldstein  
339 et al., 2006; Richardson and Cane, 2011). Therefore, we suspect that the increase in chaotic



340 behaviour during minor geomagnetic storms is strongly associated with the asymmetry features in  
341 the longitudinal distribution of solar source region for the Corotating Interaction Regions (CIR)  
342 signatures. For most of these periods of moderate geomagnetic storms, the values of MLE  
343 decreases compared to minor geomagnetic storms. This revealed that as geomagnetic stormy  
344 events build up, the level of unpredictability and sensitive dependence on initial condition (chaos)  
345 begin to decrease (Lorentz, 1963; Stogaz, 1994). The chaotic behaviour during major geomagnetic  
346 storms decreases significantly compared with moderate geomagnetic storms. The reduction in  
347 chaotic response during moderate and its further declines at major geomagnetic storms may be  
348 attributed to the disturbance in the interplanetary medium driven by sheath preceding an  
349 interplanetary coronal mass ejection (ICME) or combination of the sheath and an ICME magnetic  
350 cloud (Echer et al., 2008; Tsurutani et al., 2003; Meng et al., 2019). Notably, the dynamics of the  
351 solar wind-magnetospheric interaction are dissipative chaotic in nature (Pavlos, 2012); and, the  
352 electrodynamic of the magnetosphere due to the flux of interplanetary electric fields had a  
353 significant impact on the state of the chaotic signatures. For instance, the observation of strong  
354 chaotic behaviour during minor geomagnetic storms suggests that the dynamics was characterized  
355 by a weak magnetospheric disturbance. While the reduction in chaotic behaviour at moderate and  
356 major geomagnetic storm period reveals the dynamical features with regards to when a strong  
357 magnetospheric disturbance begins to emerge. Therefore, our observation of chaotic signatures at  
358 different categories of geomagnetic storm has potential capacity to give useful diagnostic  
359 information about monitoring space weather events. It is important to note that the features of  $D_{St}$   
360 chaotic behaviour at different categories of geomagnetic storm has not been reported in the  
361 literature. For example, previous study of Balasis et al. (2009, 2011) investigate dynamical  
362 complexity behaviour using different entropy measures and revealed the existence of low

363 dynamical complexity in the magnetospheric dynamics and attributed it to ongoing large  
364 magnetospheric disturbance (major geomagnetic storm). The work of Balasis et al. (2009, 2011)  
365 where certain dynamical characteristic evolved in the  $D_{st}$  signal was revealed was limited to one  
366 year data (2001). It is worthy to note that the year 2001, according to sunspot variations is a period  
367 of high solar activity during solar cycle 23. It is characterized by numerous and strong solar  
368 eruptions that were followed by significant magnetic storm activities. This confirms that on most  
369 of the days in year 2001, the geomagnetic activity is strongly associated with major geomagnetic  
370 storms. The confirmation of low dynamical complexity response in the  $D_{st}$  signal during major  
371 geomagnetic storms agree with our current study. However, the idea of comparing the dynamical  
372 complexity behaviour at different categories of geomagnetic storms and reveal its chaotic features  
373 was not reported. This is the major reason why our present investigation is crucial to the  
374 understanding of the level of chaos and dynamical complexity involved during different categories  
375 of geomagnetic storms. As an extension to the single-year investigation done by Balasis et al.  
376 (2009, 2011) during a major geomagnetic storm, we further investigated nine years data of  $D_{st}$   
377 that covered minor, moderate and major geomagnetic storms (see Figure 5, stem plots) and  
378 unveiled their dynamical complexity behaviour. During major geomagnetic stormy periods, we  
379 found that the ApEn values decrease significantly, indicating reduction in the dynamical  
380 complexity behaviour. This is in agreement with the low dynamical complexity reported by Balasis  
381 et al. (2009, 2011) during a major geomagnetic period. Finally, based on the method of DVV  
382 analysis, we found that test of nonlinearity in the  $D_{st}$  time series during major geomagnetic storms  
383 reveals the strongest nonlinearity features.

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#### 386 **4.2 The chaotic and dynamical complexity behaviour in the $VB_s$ as input signals.**

387 The results of the MLE values for  $VB_s$  revealed a strong chaotic behaviour during the three  
388 categories of geomagnetic storms. Comparing these MLE values during minor to those observed  
389 during moderate and major geomagnetic storms, the result obtained did not indicate any significant  
390 difference in chaoticity (bar plots, Figure 6). Also, the ApEn values of  $VB_s$  during the periods  
391 associated with minor, moderate and major geomagnetic storms revealed high dynamical  
392 complexity behaviour with no significant difference between the three categories of geomagnetic  
393 storms investigated. These observation of high chaotic and dynamical complexity behaviour in the  
394 dynamics of  $VB_s$  may be due to interplanetary discontinuities caused by the abrupt changes in the  
395 interplanetary magnetic field direction and plasma parameters (Tsurutani et al., 2010). Also, the  
396 indication of high chaotic and dynamical complexity behaviour in  $VB_s$  signifies that the solar wind  
397 electric field is stochastic in nature. The DVV analysis for  $VB_s$  revealed nonlinearity features with  
398 no significant difference between the minor, moderate and major geomagnetic storms. It is worth  
399 mentioning that the dynamical complexity behaviour for  $VB_s$  is different from what was observed  
400 for  $D_{st}$  time series data. For instance, our results for  $D_{st}$  times series revealed that the chaotic and  
401 dynamical complexity behaviour of the magnetospheric dynamics are high during minor  
402 geomagnetic storms, reduce at moderate geomagnetic storms and further decline during major  
403 geomagnetic storms. While the  $VB_s$  signal revealed a high chaotic and dynamical complexity  
404 behaviour at all the categories of geomagnetic storm period. Therefore, these dynamical features  
405 obtained in the  $VB_s$  as input signal and the  $D_{st}$  as the output in describing the magnetosphere as a  
406 non-autonomous system further support the finding of Donner et al. (2019) that found increased  
407 or not changed in dynamical complexity behaviour for  $VB_s$  and low dynamical complexity

408 behaviour during storm using recurrence method. Thus, suggesting that the magnetospheric  
409 dynamics is nonlinear and the solar wind dynamics is consistently stochastic in nature.

## 410 **5.0 Conclusions**

411 This work has examined the magnetospheric chaos and dynamical complexity behaviour in the  
412 disturbance storm time ( $D_{st}$ ) and solar wind electric field ( $VB_s$ ) as input during different categories  
413 of geomagnetic storms. The chaotic and dynamical complexity behaviour at minor, moderate and  
414 major geomagnetic storms for solar wind electric field ( $VB_s$ ) as input and  $D_{st}$  as output of the  
415 magnetospheric system were analyzed for the period of 9 years using nonlinear dynamics tools.  
416 Our analysis has shown a noticeable trend of these nonlinear parameters (MLE and ApEn) and the  
417 categories of geomagnetic storm (minor, moderate and major). The MLE and ApEn values of the  
418  $D_{st}$  have indicated that the chaotic and dynamical complexity behaviour are high during minor  
419 geomagnetic storms, low during moderate geomagnetic storms and further reduced during major  
420 geomagnetic storms. The values of MLE and ApEn obtained from  $VB_s$  indicate that chaotic and  
421 dynamical complexity are high with no significant difference during the periods of minor,  
422 moderate and major geomagnetic storms. Finally, the test for nonlinearity in the  $D_{st}$  time series  
423 during major geomagnetic storms reveals the strongest nonlinearity features. Based on these  
424 findings, the dynamical features obtained in the  $VB_s$  as input and  $D_{st}$  as output of the  
425 magnetospheric system suggest that the magnetospheric dynamics is nonlinear and the solar wind  
426 dynamics is consistently stochastic in nature.

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## 430 **7.0 Acknowledgement**

431 The authors would like to acknowledge the World Data Centre for Geomagnetism, Kyoto, and the  
432 National Aeronautics and Space Administration, Space Physics Facility (NASA) for making the  
433 Dst data and solar wind plasma data available for research purpose. Also, the efforts of anonymous  
434 reviewers in ensuring that the quality of the work is enhanced are highly appreciated.

## 435 **Declaration of Interest statement**

436 The authors declare that there is no conflicts of interest.

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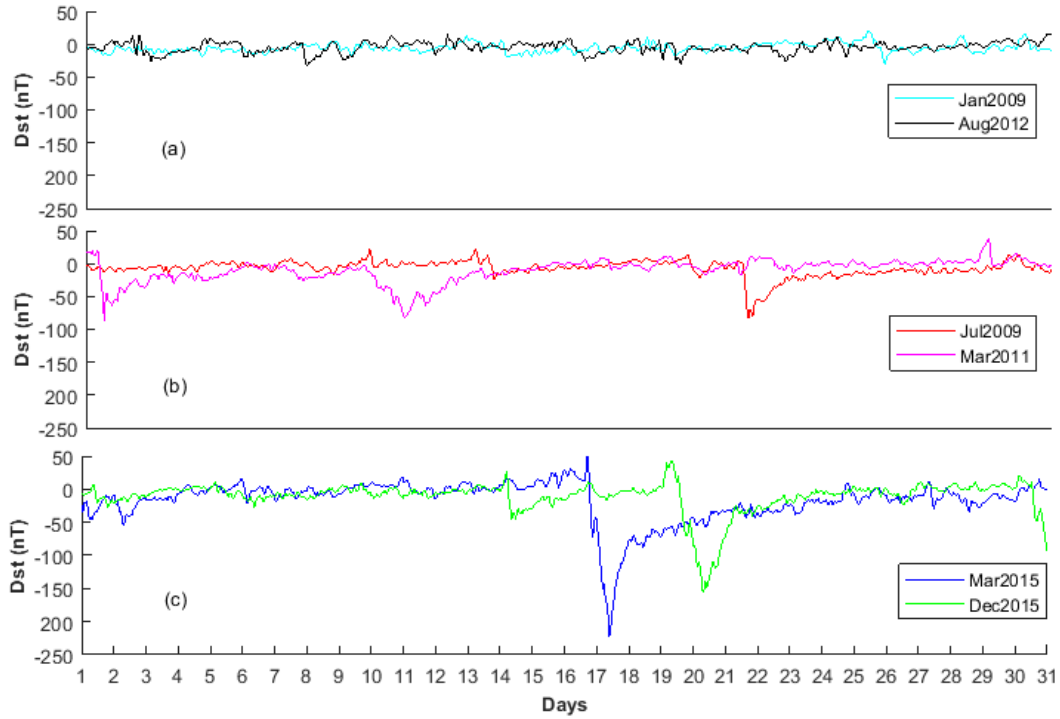
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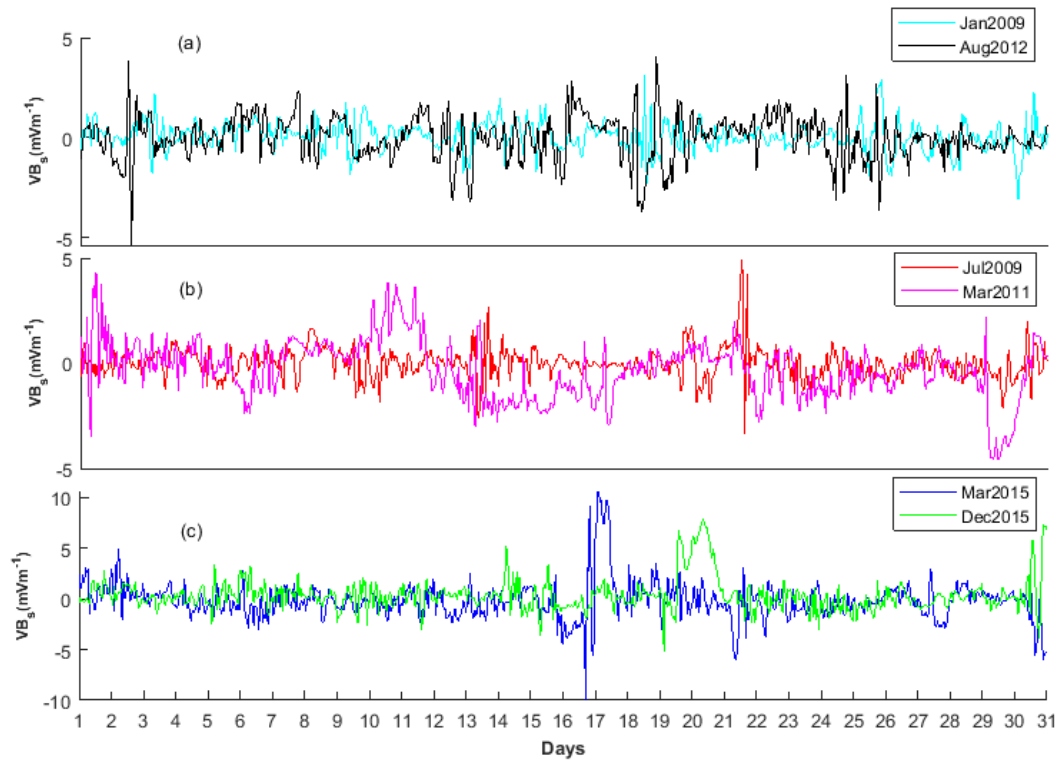


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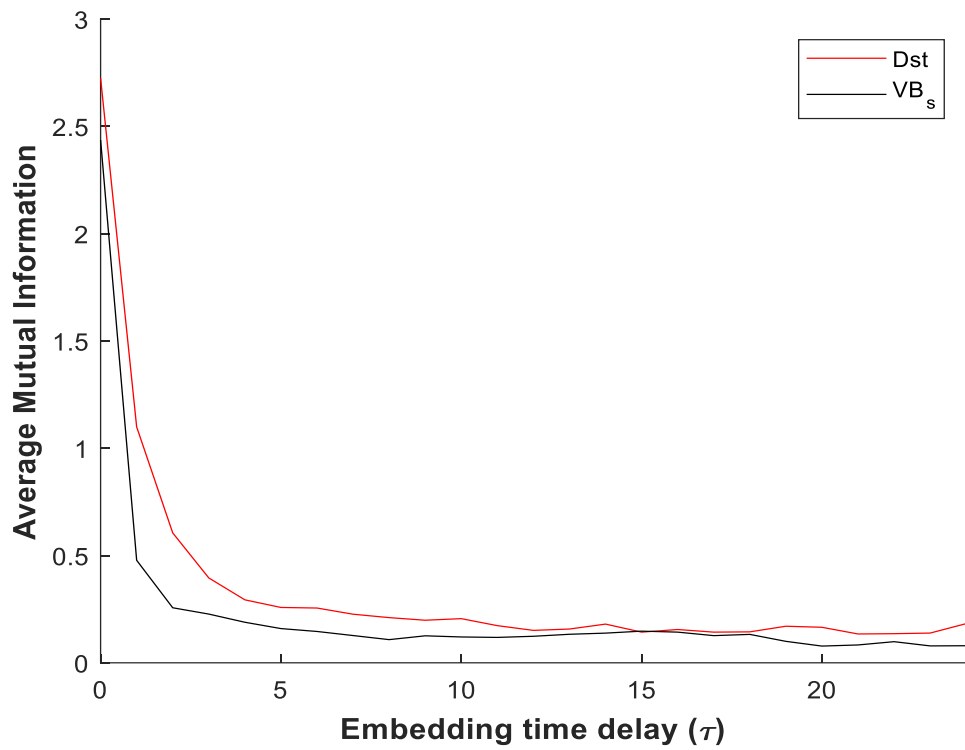
676 Figure 1: Samples of  $D_{st}$  signals classified as (a) Minor, (b) Moderate and (c) Major geomagnetic  
 677 storm



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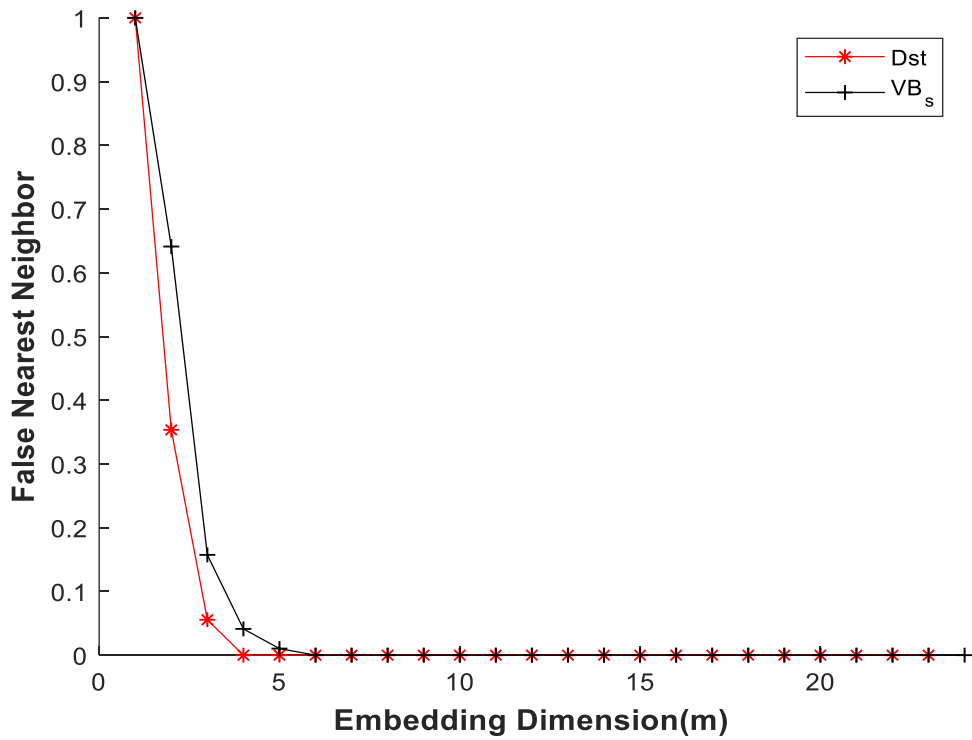
679 Figure 2: Samples of solar wind electric fields ( $VB_s$ ) during (a) Minor, (b) Moderate and (c) Major  
 680 geomagnetic storm.

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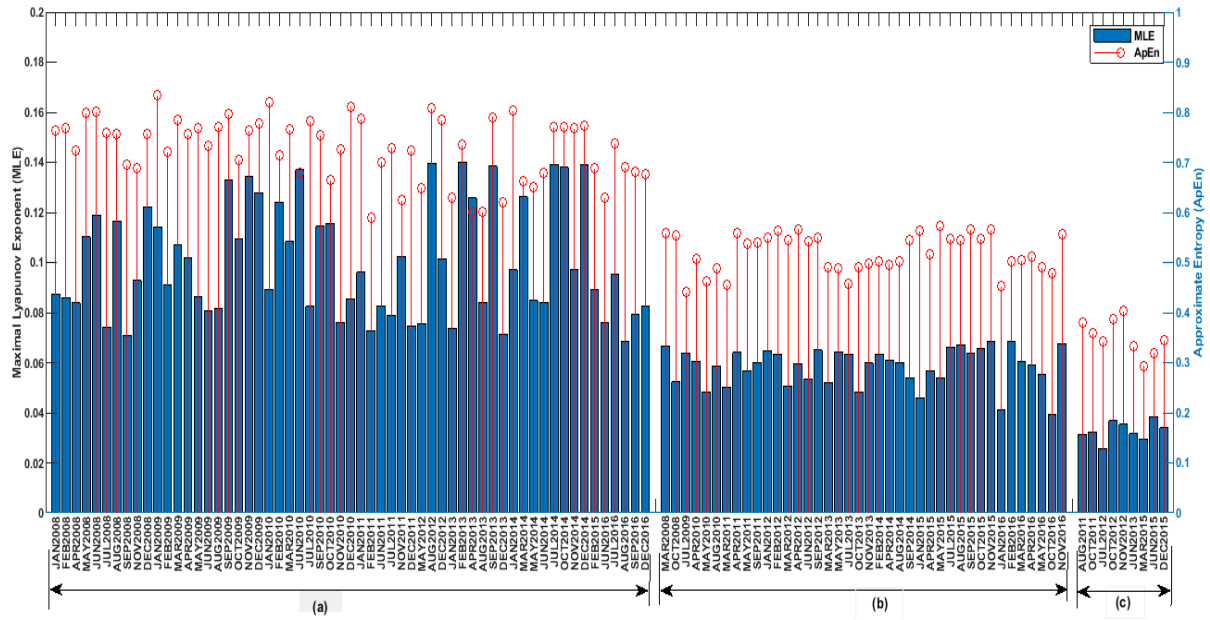
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683 Figure 3: The plot Average Mutual Information against embedding time delay ( $\tau$ )



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685 Figure 4: The plot of Percentage of False Nearest Neighbors against embedding dimension ( $m$ )



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687 Figure 5: The MLE (bar plot) and ApEn (stem plot) of  $D_{st}$  at: (a) Minor, (b) Moderate and (c)  
 688 Major geomagnetic storm.

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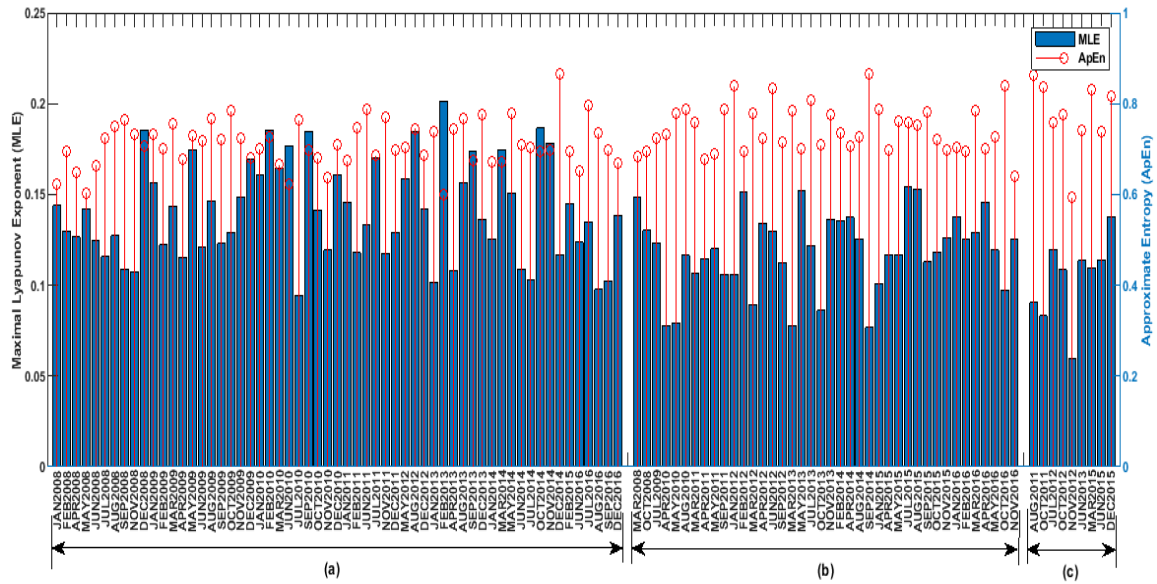
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697 Figure 6: The MLE (bar plot) and ApEn (stem plot) of solar wind electric field ( $VB_s$ ) during: (a)  
 698 Minor, (b) Moderate and (c) Major geomagnetic storm.

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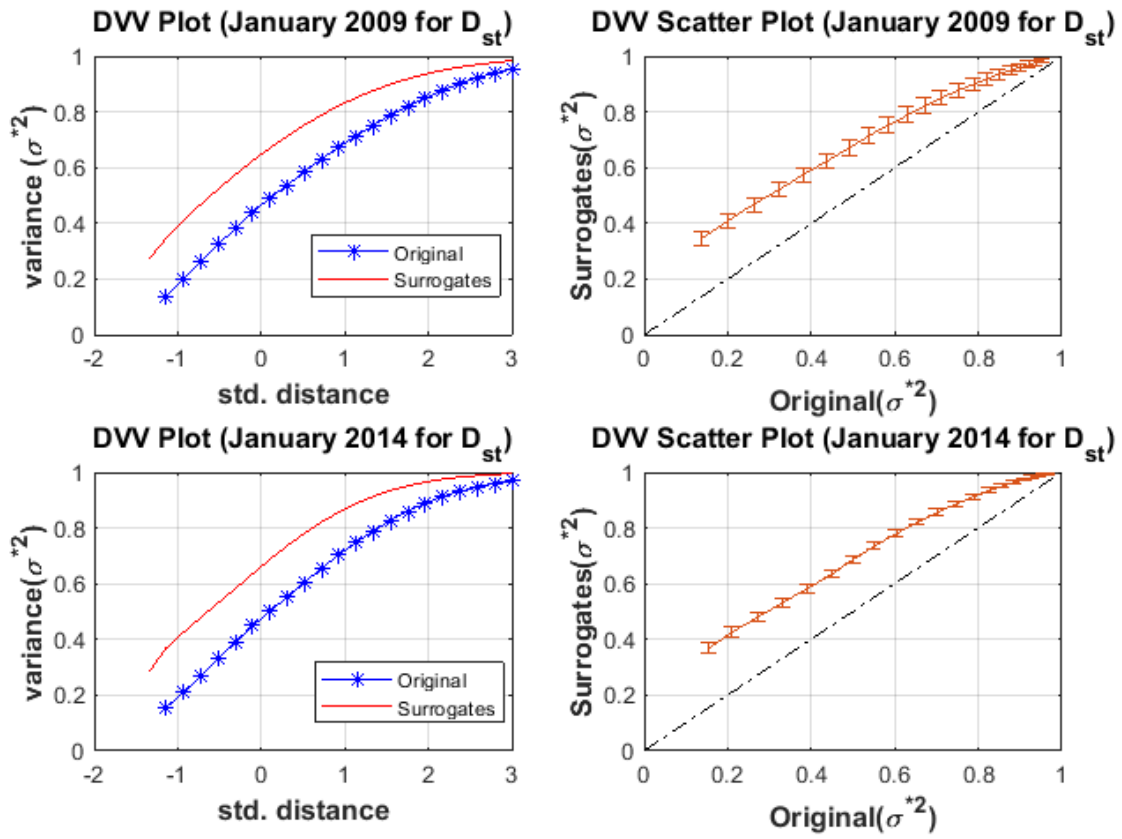
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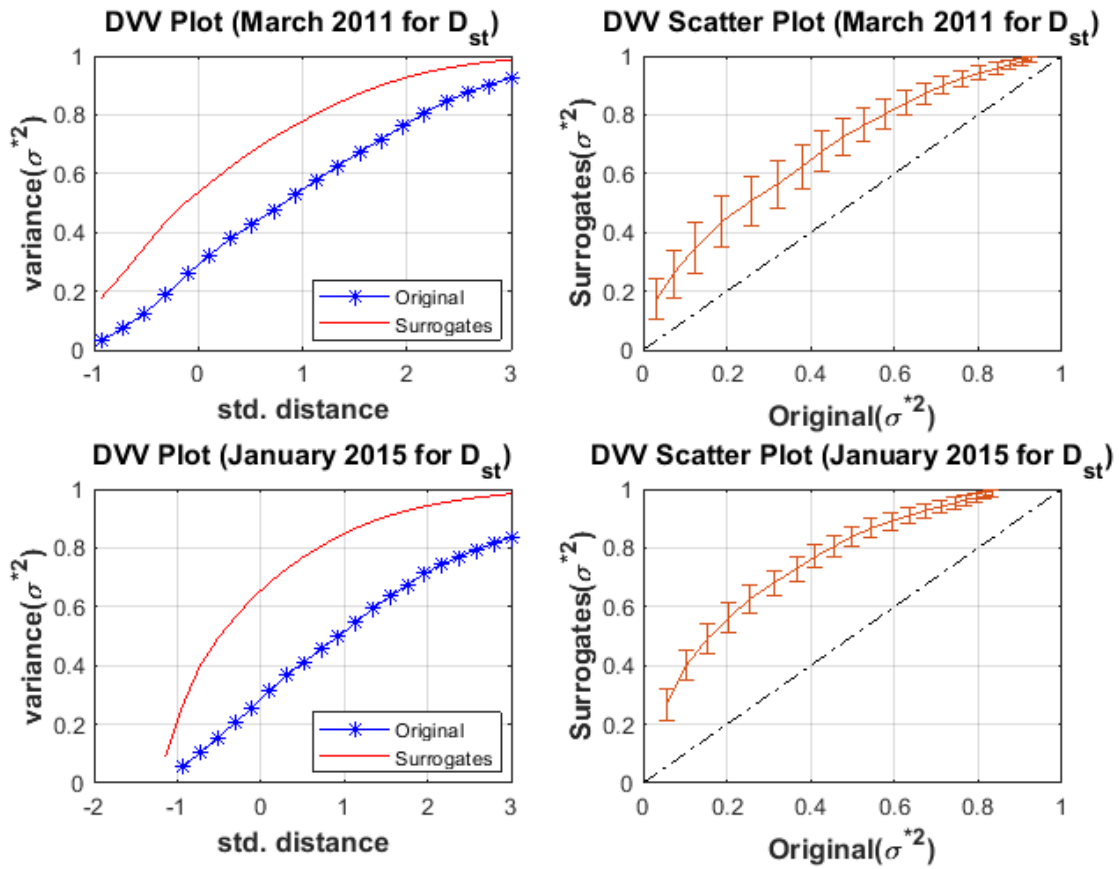
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713 Figure 7: The DVV plot and Scatter plot for  $D_{st}$  during minor geomagnetic storm for January  
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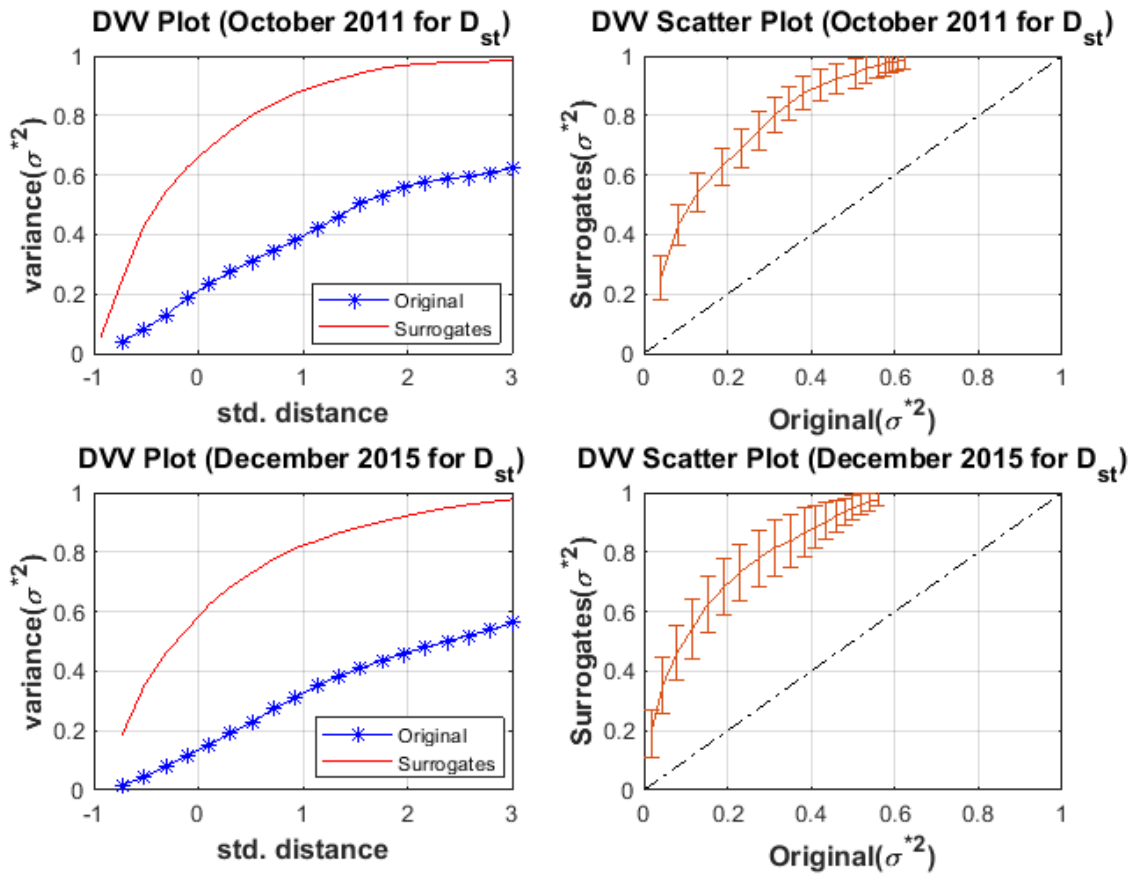


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717 Figure 8: The DVV plot and Scatter plot for  $D_{st}$  during moderate geomagnetic storm for March  
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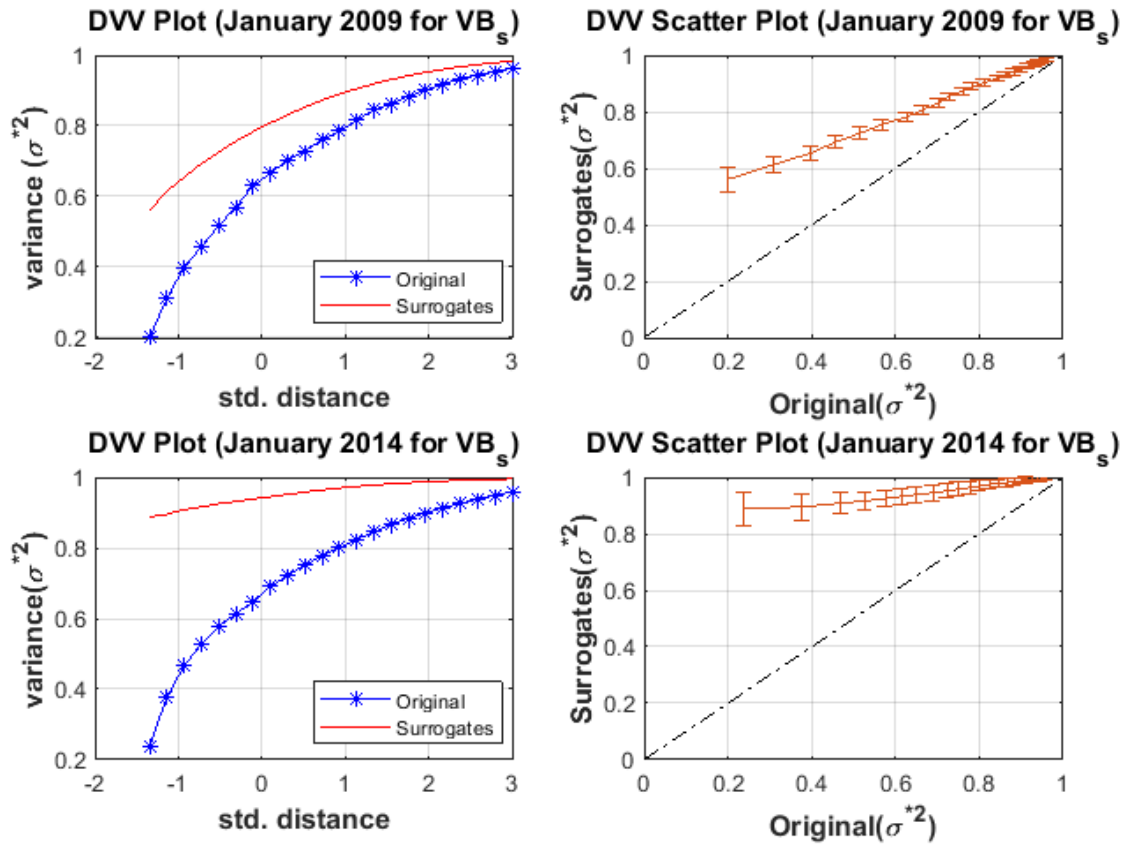


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721 Figure 9: The DVV plot and Scatter plot for  $D_{st}$  during major geomagnetic storm for October  
 722 2011 and December 2015.

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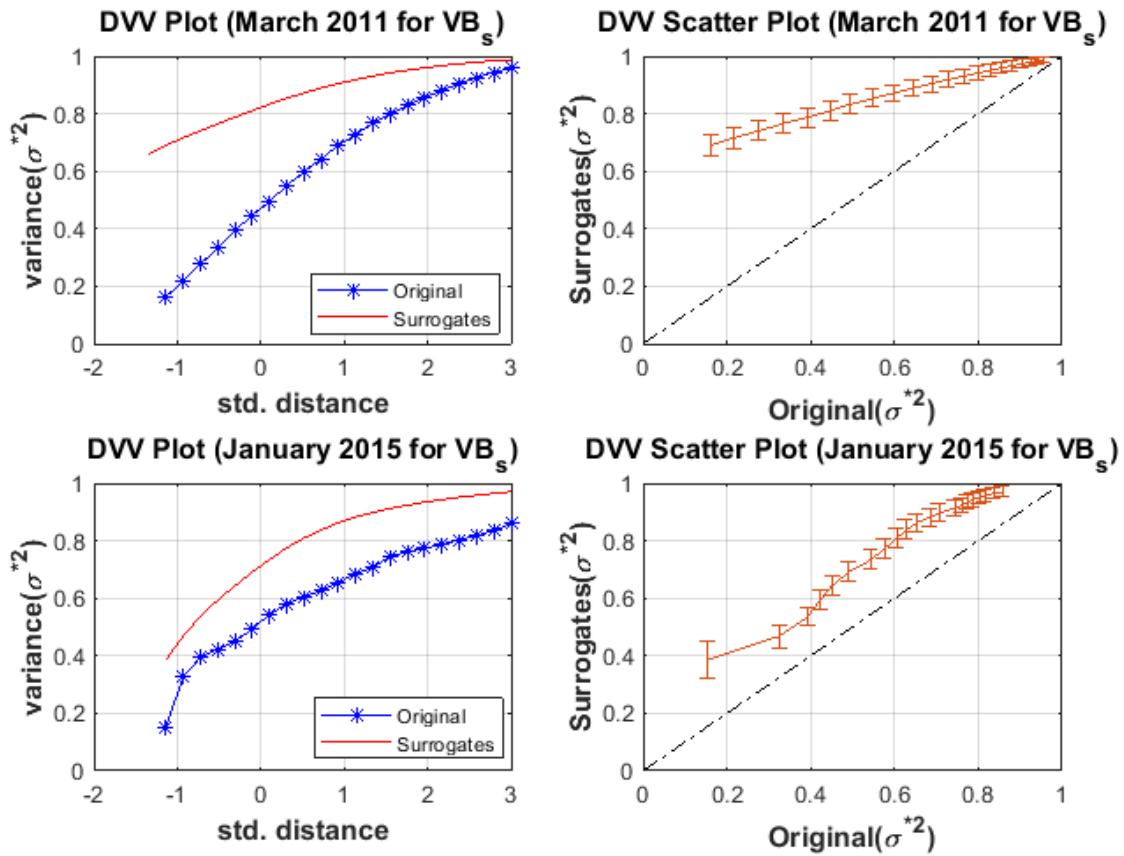


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725 Figure 10: The DVV plot and Scatter plot for  $VB_s$  during minor geomagnetic storm for January  
 726 2009 and January 2014.

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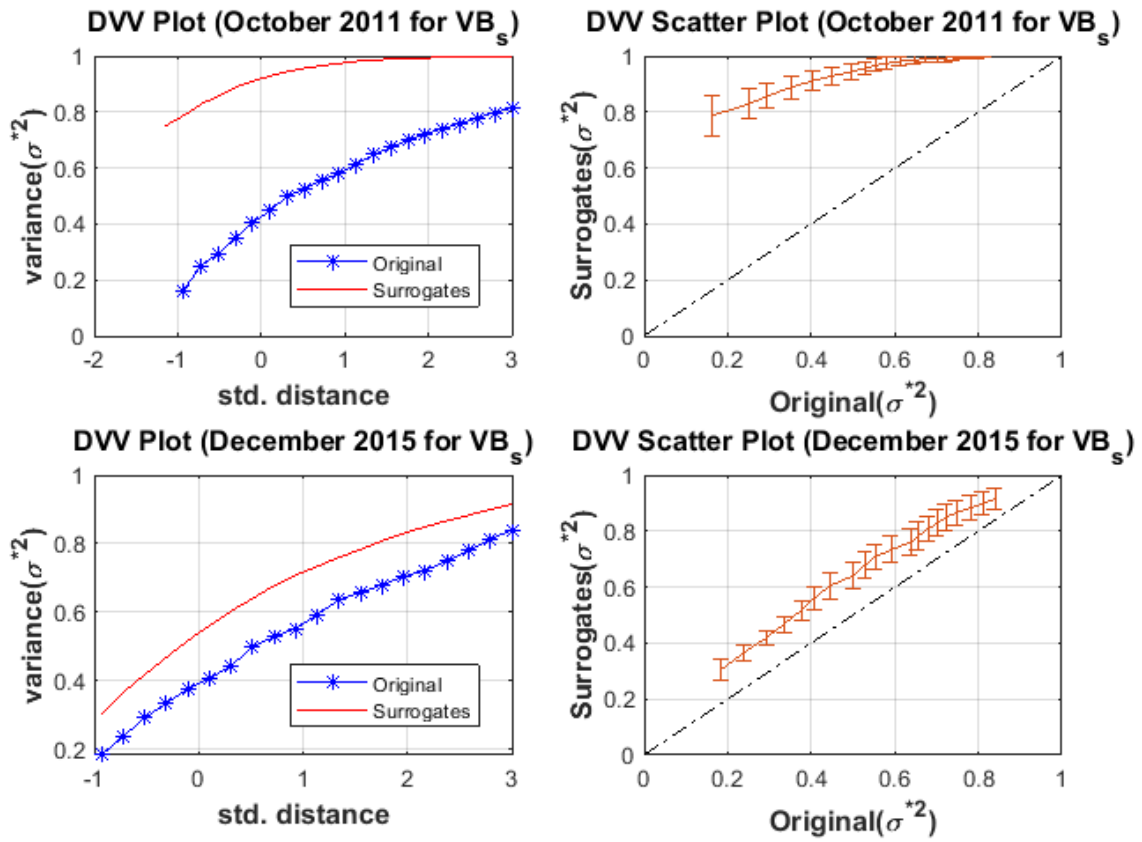


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730 Figure 11: The DVV plot and Scatter plot for  $VB_s$  during moderate geomagnetic storm for March  
 731 2011 and January 2015.

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735 Figure 12: The DVV plot and Scatter plot for  $VB_s$  during major geomagnetic storm for October  
 736 2011 and December 2015.

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