Magnetospheric chaos and dynamical complexity response during storm time disturbance

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Abstract

In this study, we examine the magnetospheric chaos and dynamical complexity response to the disturbance storm time (D_{st}) and solar wind electric field (VB_s) during different categories of geomagnetic storm (minor, moderate and major geomagnetic storm). The time series data of the D_{st} and VB_s are analyzed for the period of nine years using nonlinear dynamics tools (Maximal Lyapunov Exponent, MLE, Approximate Entropy, ApEn and Delay Vector Variance, DVV). We found a significant trend between each nonlinear parameter and the categories of geomagnetic storm. The MLE and ApEn values of the D_{st} indicate that chaotic and dynamical complexity responses are high during minor geomagnetic storms, reduce at moderate geomagnetic storms and decline further during major geomagnetic storms. However, the MLE and ApEn values obtained from VB_s indicate that chaotic and dynamical complexity response are high with no significant difference between the periods that are associated with minor, moderate and major geomagnetic storms. The test for nonlinearity in the D_{st} time series during major geomagnetic storm reveals the strongest nonlinearity features. Based on these findings, the dynamical features obtained in the VB_s as input and D_{st} as output of the magnetospheric system suggest that the magnetospheric dynamics is nonlinear and the solar wind dynamics is consistently stochastic in nature.

Keywords: D_{st} signals, Solar wind electric field (VB_s) signals, Geomagnetic storm, Chaotic behaviour, Dynamical complexity, Nonlinearity.

1.0 Introduction

The response of chaos and dynamical complexity behaviour with respect to magnetospheric dynamics varies. This is due to changes in the interplanetary electric fields imposed on the magnetopause and those penetrating the inner magnetosphere and sustaining convection thereby initiating geomagnetic storm (Pavlos et al. 1992). A prolonged southward turning of interplanetary magnetic field (IMF,B_z), which indicates that solar wind-magnetosphere coupling is in-progress was confirmed on many occasions for which such geomagnetic storm was driven by Corotating Interaction Regions (CIRs), or by the sheath preceding an interplanetary coronal mass ejection (ICME) or by a combination of the sheath and an ICME magnetic cloud (Russell et al. 1974; Burton et al.1975; Gonzalez and Tsurutani, 1987; Tsurutani et al. 1988; Cowley, 1995; Tsutomu, 2002; Yurchyshyn et al. 2004; Kozyra et al. 2006; Echer et al. 2008; Meng et al. 2019; Tsurutani et al. 2020). Notably, the introduction of Disturbance Storm Time (D_{st}) index (Sugiura, 1964; Sugiura and Kamei, 1991) unveil the quantitative measure of the total energy of the ring current particles. Therefore, the D_{st} index remains one of the most popular global indicators that can precisely reveal the severity of a geomagnetic storm (Dessler and Parker, 1959).

The D_{st} fluctuations exhibit different signatures for different categories of geomagnetic storm. Ordinarily, one can easily anticipate that fluctuations in a D_{st} signal appear chaotic and complex. These may arise from the changes in the interplanetary electric fields driven by the solar wind-magnetospheric coupling processes. At different categories of geomagnetic storm, fluctuations in the D_{st} signals differ (Oludehinwa et al. 2018). One obvious reason is that as the intensity of the geomagnetic storm increases, the fluctuation behaviour in the D_{st} signal becomes more complex and nonlinear in nature. It has been established that the electrodynamic response of the magnetosphere to solar wind driver are non-autonomous in nature (Price and Prichard, 1993; Price

et al. 1994; Johnson and Wings, 2005). Therefore, the chaotic analysis of the magnetospheric time series must be related to the concept of input-output dynamical process. Consequently, it is necessary to examine the chaotic behaviour of the solar wind electric field (VB_s) as input signals and the magnetospheric activity index (D_{st}) as output during different categories of geomagnetic storms.

Several works have been presented on the chaotic and dynamical complexity behaviour of the magnetospheric dynamics based on autonomous concept, i.e using the time series data of magnetospheric activity alone such as auroral electrojet (AE), Amplitude Lower (AL) and D_{st} index (Vassiliadis et al.1990; Baker and Klimas, 1990; Vassiliadis et al.1991; Shan et al. 1991; Pavlos et al. 1994; Klimas et al. 1996; Valdivia et al. 2005; Mendes et al. 2017; Consolini, 2018). They found evidence of low-dimensional chaos in the magnetospheric dynamics. For instance, the report by Vassiliadis et al. (1991) shows that the computation of Lyapunov exponent for AL index time series gives a positive value of Lyapunov exponent indicating the presence of chaos in the magnetospheric dynamics. Unnikrishnan, (2008) studied the deterministic chaotic behaviour in the magnetospheric dynamics under various physical conditions using AE index time series and found that the seasonal mean value of Lyapunov exponent in winter season during quiet periods ($0.7 \pm 0.11\ min^{-1}$) is higher than that of the stormy periods ($0.36 \pm 0.09\ min^{-1}$). Balasis et al. (2006) examined the magnetospheric dynamics in the D_{st} index time series from pre-magnetic storm to magnetic storm period using fractal dynamics. They found that the transition from anti-persistent to persistent behaviour indicates that the occurrence of an intense geomagnetic storm is imminent. Balasis et al. (2009) further reveal the dynamical complexity behaviour in the magnetospheric dynamics using various entropy measures. They reported a significant decrease in dynamical complexity and an accession of persistency in the D_{st} time series as the magnetic storm

approaches. Recently, Oludehinwa et al. (2018) examined the nonlinearity effects in D_{st} signals during minor, moderate and major geomagnetic storm using recurrence plot and recurrence quantification analysis. They found that the dynamics of the D_{st} signal is stochastic during minor geomagnetic storm periods and deterministic as the geomagnetic storm increases.

Also, studies describing the solar wind and magnetosphere as non-autonomous system have been extensively investigated. Price et al. (1994) examine the nonlinear input-output analysis of AL index and different combinations of interplanetary magnetic field (IMF) with solar wind parameters as input function. They found that only a few of the input combinations show any evidence whatsoever for nonlinear coupling between the input and output for the interval investigated. Pavlos et al. (1999) presented further evidence of magnetospheric chaos. They compared the observational behaviour of the magnetospheric system with the results obtained by analyzing different types of stochastic and deterministic input-output systems and asserted that a low dimensional chaos is evident in magnetospheric dynamics. Devi et al. (2013) studied the magnetospheric dynamics using AL index with the southward component of IMF, (Bz) and observed that the magnetosphere and turbulent solar wind have values corresponding to nonlinear dynamical system with chaotic behaviour. The modeling and forecasting approach have been applied to magnetospheric time series using nonlinear models (Valdivia et al. 1996; Vassiliadis et al. 1999; Vassiliadis, 2006; Balikhin et al. 2010). These efforts have improved our understanding with regards to the facts that nonlinear dynamics can reveal some hidden dynamical information in the observational time series. In addition to these nonlinear effects in D_{st} signals, a measure of the exponential divergence and convergence within the trajectories of a phase space known as Maximal Lyapunov Exponent (MLE), which has the potential to depict the chaotic behavior in the D_{st} and VB_s time series during a minor, moderate and major geomagnetic storm have not been

investigated. In addition, to the best of our knowledge, computation of Approximate Entropy (ApEn) that depicts the dynamical complexity behaviour during different categories of geomagnetic storm has not been reported in the literature. The test for nonlinearity through delay vector variance (DVV) analysis that establishes the degree at which nonlinearity response in D_{st} time series during minor, moderate and major geomagnetic storms is not well known. It is worth to note that understanding the dynamical characteristics in the D_{st} and VB_s signals at different categories of geomagnetic storms will provide useful diagnostic information to different conditions of space weather phenomenon. Consequently, this study attempts to carry out comprehensive numerical analysis to unfold the chaotic and dynamical complexity behaviour in the D_{st} and VB_s signals during minor, moderate and major geomagnetic storm. In section 2, our methods of data acquisition are described. Also, the nonlinear analysis that we employed in this investigation are detailed. In section 3, we unveiled our results and engage the discussion of results in section 5.

2.0 Description of the Data and Nonlinear Dynamics

The D_{st} index is derived by measurements from ground-based magnetic stations at low-latitudes observatories around the world and depicts mainly the variation of the ring current, as well as the Chapman-Ferraro Magnetopause currents, and tail currents to a lesser extent (Sugiura, 1964; Love and Gannon, 2009). Due to its global nature, D_{st} time series provides a measure of how intense a geomagnetic storm was (Dessel and Parker, 1959). In this study, we considered D_{st} data for the period of nine years from January to December between 2008 and 2016 which were downloaded from the World Data Centre for Geomagnetism, Kyoto, Japan (http://wdc.kugi-kyoto-u.ac.jp/Dstae/index.html). The sampling time of D_{st} and VB_s time series data was 1-hour. We use the classification of geomagnetic storms as proposed by Gonzalez et al. (1994) such that D_{st} index value in the ranges $0 \leq Dst \leq -50nT$, $-50nT \leq Dst \leq -100nT$, $-100nT \leq Dst \leq -250nT$

are classified as minor, moderate and major geomagnetic storms respectively and each month is being classified based on its minimum Dst value. The solar wind electric field (VB_s) data are archived from the National Aeronautics and Space Administration, Space Physics Facility (http://omniweb.gsfc.nasa.gov). It is well known that the dynamics of the solar wind contribute to the driving of the magnetosphere (Burton et al. 1975). Furthermore, we took the solar wind electric field (VB_s) as the input signal (Price and Prichard, 1993; Price et al. 1994). The VB_s was categorized according to the periods of minor, moderate and major geomagnetic storm. Then, the D_{st} and VB_s time series were subjected to a variety of nonlinear analytical tools explained as follow:

2.1 Phase Space Reconstruction and Observational time series

An observational time series can be defined as a sequence of scalar measurements of some quantity, which is a function of the current state of the system taken at multiples of a fixed sampling time. In nonlinear dynamics, the first step in analyzing an observational time series data is to reconstruct an appropriate state space of the system. Takens, (1981) and Mane, (1981) stated that one time series or a few simultaneous time series are converted to a sequence of vectors. This reconstructed phase space has all the dynamical characteristic of the real phase space provided the time delay and embedding dimension are properly specified.

$$X(t) = [x(t), x(t+\tau), x(t+2\tau), \dots, x(t+(m-1)\tau]^T \quad (1)$$

Where $X(t)$ is the reconstructed phase space, $x(t)$ is the original time series data, τ is the time delay and m is the embedding dimension. An appropriate choice of τ and m are needed for the reconstruction phase space which is determined by average mutual information and false nearest neighbour respectively.

2.2 Average Mutual Information (AMI)

The method of Average Mutual Information (AMI) is one of the nonlinear techniques used to determine the optimal time delay (τ) required for phase space reconstruction in observational time series. The time delay mutual information was proposed by Fraser and Swinney, (1986) instead of autocorrelation function. This method takes into account nonlinear correlations within the time series data. It measures how much information can be predicted about one time series point, given full information about the other. For instance, the mutual information between x_i and $x_{(i+\tau)}$ quantifies the information in state $x_{(i+\tau)}$ under the assumption that information at the state x_i is known. The AMI for a time series, $x(t_i), \quad i = 1,2, \dots, N$ is calculated as:

$$I(T) = \sum_{x(t_i),\ x(t_i+T)} P(x(t_i), x(t_i+T)) \times \log_2 \left[\frac{P(x(t_i), x(t_i+T))}{P(x(t_i))\, P(x(t_i+T))}\right] \qquad (2)$$

Where $x(t_i)$ is the ith element of the time series, $T = k\Delta t \quad (k = 1,2, \dots, k_{max})$, $P(x(t_i))$ is the probability density at $x(t_i)$, $P(x(t_i), x(t_i+T))$ is the joint probability density at the pair $x(t_i), x(t_i+T)$. The time delay (τ) of the first minimum of AMI is chosen as optimal time delay (Fraser and Swinney, 1986). Therefore, the AMI was applied to the D_{st} and VB_s time series and the plot of AMI against time delay is shown in Figure (3). We notice that the AMI showed the first local minimum at roughly $(\tau = 15hr)$. Furthermore, the values of τ near this value of (~15hr) maintain constancy for both VBs and D_{st}. In the analysis $(\tau = 15hr)$ was used as the optimal time delay for the computation of maximal Lyapunov exponent.

2.3 False Nearest Neighbour (FNN)

In determining the optimal choice of embedding dimension(m), the false nearest neighbour method was used in the study. It was suggested by Kennel et al. (1992). The concept is based on how the number of neighbours of a point along a signal trajectory changes with increasing embedding dimension. With increasing embedding dimension, the false neighbour will no longer be neighbours, therefore by examining how the number of neighbours changes as a function of dimension, an appropriate embedding dimension can be determined. For instance, suppose we have a one-dimensional time series. We can construct a time series $y(t)$ of D-dimensional points from the original one-dimensional time series $x(t)$ as follows:

$$y(t) = (x(t), x(t+\tau), \dots, x(t+(D-1)\tau) \qquad (3)$$

Where τ and D are time delay and embedding dimension. Using the formular from Kennel et al. (1992); Wallot and Monster, (2018). If we have a D-dimensional phase space and denote the rth nearest neighbour of a coordinate vector $y(t)$ by $y^{(r)}(t)$, then the square of the Euclidean distance between $y(t)$ and the rth nearest neighbor is:

$$R_D^2(t,r) = \sum_{k=0}^{D-1} \left[x(t+k\tau) - x^{(r)}(t+k\tau)\right]^2 \qquad (4)$$

Now applying the logic outlined above, we can go from a D-dimensional phase space to $(D+1)$ dimensional phase space by time-delay embedding, adding a new coordinate to $y(t)$, and ask what is the squared distance between $y(t)$ and the same rth nearest neighbour:

$$R_{D+1}^2(t,r) = R_D^2(t,r) + \left[x(t+D\tau) - x^{(r)}(t+D\tau)\right]^2 \qquad (5)$$

As explained above, if the one-dimensional time series is already properly embedded in D dimensions, then the distance R between $y(t)$ and the rth nearest neighbour should not appreciably change by some distance criterion $R_{tol}(i.e\ R < R_{tol})$. Moreover, the distance of the nearest neighbour when embedded into the next higher dimension relative to the size of the attractor should be less than some criterion $A_{tol}(i.e\ R_{D+1} < A_{tol})$. Doing this for the nearest neighbour of each coordinate will result on many false nearest neighbours when embedding is insufficient or in few (or no) false neighbours when embedding is sufficient. In the analysis, the FNN was applied to the D_{st} and VB_s time series to detect the optimal value of embedding dimension(m). Figure (4) shows a sample plot of the percentage of false nearest neighbour against embedding dimension in one of the months under investigation (other months show similar results, thus for brevity we depict only one of the results). We notice that the false nearest neighbor attains its minimum value at $m \geq 5$ indicating that embedding dimension (m) from $m \geq 5$ are optimal values. Therefore, $m = 5$ was used for the computation of maximal Lyapunov exponent.

2.4 Maximal Lyapunov Exponent (MLE)

The Maximal Lyapunov Exponent (MLE) is one of the most popular nonlinear dynamics tool used for detecting chaotic behaviour in a time series data. It describes how small changes in the state of a system grow at an exponential rate and eventually dominate the behaviour. An important indication of chaotic behavior of a dissipative deterministic system is the existence of a positive Lyapunov Exponent. A positive MLE signifies divergence of trajectories in one direction or expansion of an initial volume in this direction. On the other hand, a negative MLE exponent implies convergence of trajectories or contraction of volume along another direction. The algorithm proposed by Wolf et al. (1985) for estimating MLE is employed to compute the chaotic behavior of the D_{st} and VB_s time series at minor, moderate and major geomagnetic storm. Other

methods of determining MLE includes Rosenstein's method, Kantz's method and so on. In this study, the MLE at minor, moderate and major geomagnetic storms periods was computed with $m = 5$ and $\tau = 15hr$ as shown in figures (5 & 6-bar plots) for D_{st} and VB_s. The calculation of MLE is explained as follows: given a sequence of vector $x(t)$, an m-dimensional phase space is formed from the observational time series through embedding theorem as

$$\{x(t), x(t+\tau), \dots, x(t+(m-1)\tau)\} \quad (6)$$

Where m and τ are as defined earlier, after reconstructing the observational time series, the algorithm locates the nearest neighbour (in Euclidean sense) to the initial point $\{x(t_0), \dots, x(t_0 + (m-1)\tau\}$ and denote the distance between these two points $L(t_0)$. At a later point t_1, the initial length will have evolved to length $L'(t_1)$. Then the MLE is calculated as:

$$\lambda = \frac{1}{t_M - t_0} \sum_{k=1}^{M} \log_2 \frac{L'(t_k)}{L(t_{k-1})} \quad (7)$$

M is the total number of replacement steps. We look for a new data point that satisfies two criteria reasonably well: its separation, $L(t_1)$, from the evolved fiducial point is small. If an adequate replacement point cannot be found, we retain the points that were being used. This procedure is repeated until the fiducial trajectory has traversed the entire data

2.5 Approximate Entropy (ApEn)

Approximate Entropy (ApEn) is one of the nonlinear dynamics tools that measure the dynamical complexity in observational time series. The concept was proposed by Pincus, (1991) which provides a generalized measure of regularity, such that it accounts for the logarithm likehood in the observational time series. For instance, a dataset of length, N, that repeat itself for m points within a boundary will again repeat itself for $m + 1$ points. Because of its computational

advantage, ApEn has been widely used in many areas of disciplines to study dynamical complexity (Pincus and Kalman (2004); Pincus and Goldberger (1994); McKinley et al. (2011); Kannathan et al. (2005); Balasis et al. (2009); Shujuan and Weidong, (2010); Moore and Marchant (2017)). The ApEn is computed using the formula below:

$$ApEn(m,r,N) = \frac{1}{N-m+1}\sum_{i=1}^{N-m+1}\log C_i^m(r) - \frac{1}{N-m}\sum_{i=1}^{N-m}\log C_i^m(r) \quad (8)$$

where $C_i^m(r) = \frac{1}{N-m+1}\sum_{j=1}^{N-m+1}\Theta\left(r - \left\|x_i - x_j\right\|\right)$ is the correlation integral, m is the embedding dimension and r is the tolerance. To compute the ApEn for the D_{st} and VB_s time series classified as minor, moderate and major geomagnetic storm from 2008 to 2016, we choose $(m = 3, \tau = 1hr)$. We refer the works of Pincus, (1991); Kannathal et al. (2005); and Balasis et al. (2009) to interested readers where all the computational steps regarding ApEn were explained in details. Figures (5 & 6) depict the stem plot of ApEn for D_{st} and (VB_s) from 2008 to 2016.

2.6 Delay Vector Variance (DVV) analysis

The Delay Vector Variance (DVV) is a unified approach in analyzing and testing for nonlinearity in a time series (Gautama et al. 2004; Mandic et al. 2007). The basic idea of the DVV is that, if two delay vectors of a predictable signal are close to each other in terms of the Euclidean distance, they should have similar target. For instance, when a time delay (τ) is embedded into a time series $x(k),\ k = 1,2,\dots,N$, then a reconstructed phase space vector is formed which represents a set of delay vectors (DVs) of a given dimension.

$$X(k) = [X_{k-m\tau}, \dots, X_{k-\tau}]^T \quad (9)$$

Reconstructing the phase space, a set (λ_k) is generated by grouping those DVs that are with a certain Euclidean distance to DVs $\big(X(k)\big)$. For a given embedding dimension (m), a measure of unpredictability $\sigma *^2$ is computed over all pairwise Euclidean distance between delay vector as

$$d(i,j) = \|x(i) - x(j)\| \quad (i \neq j) \tag{10}$$

Then, sets $\lambda_k(r_d)$ are generated as the sets which consist of all delay vectors that lie closer to $x(k)$ than a certain distance r_d.

$$\lambda_k(r_d) = \{x(i) \| x(k) - x(i)\| \leq r_d\} \tag{11}$$

For every set $\lambda_k(r_d)$, the variance of the corresponding target $\sigma *^2 (r_d)$ is

$$\sigma *^2 (r_d) = \frac{\frac{1}{N}\sum_{k=1}^{N} \sigma_k^2(r_d)}{\sigma_k} \tag{12}$$

where $\sigma *^2 (r_d)$ is target variance against the standardized distance indicating that Euclidean distance will be varied in a manner standardized with respect to the distribution of pairwise distance between DVs. Iterative Amplitude Adjusted Fourier Transform (IAAFT) method is used to generate the surrogate time series (Kugiumtzis, 1999). If the surrogate time series yields DV plots similar to the original time series and the scattered plot coincides with the bisector line, then the original time series can be regarded as linear (Theiler et al. 1992; Gautama et al.2004; Imitaz, 2010; Jaksic et al. 2016). On the other hand, if the surrogate time series yield DV plot that is not similar to that of the original time series, then the deviation from the bisector lines indicates nonlinearity. The deviation from the bisector lines grows as a result of the degree of nonlinearity in the observational time series.

$$t^{DVV} = \sqrt{\langle(\sigma^{*2}(r_d) - \frac{\sum_{i=1}^{N} \sigma_{s,i}^{*2}}{N_s})\rangle} \quad (13)$$

where $\sigma_{s,i}^{*2}(r_d)$ is the target variance at the span r_d for the i^{th} surrogate. To carry out the test for nonlinearity in the D_{st} signals, $m = 3$ and $n_d = 3$, the number of reference DVs=200, and number of surrogate, $N_s = 25$ was used in all the analysis. Then we examined the nonlinearity response at minor, moderate and major geomagnetic storm.

3.0 Results

In this study, D_{st} and VB_s time series from January to December was analyzed for the period of nine years (2008 to 2016) to examine the chaotic and dynamical complexity response in the magnetospheric dynamics during the month of minor, moderate and major geomagnetic storms activity. Figures (1) & (2), display the samples of fluctuation signatures of D_{st} and VB_s signals classified as (a): the month of minor, (b): the month of moderate and (c): the month of major geomagnetic storm activity. The plot of Average Mutual information against time delay (τ) shown in Figure (3) depicts that the first local minimum of the AMI function was found to be roughly at $\tau = 15$hr. Furthermore, we notice that the values of τ near this value of (~15hr) maintain constancy for both VB_s and D_{st}. Also, in Figure (4), we display the plot of the percentage of false nearest neighbour against embedding dimension (m). It is obvious that a decrease in false nearest neighbour when increasing the embedding dimension drop steeply to zero at the optimal dimension($m = 5$), thereafter the false neighbours stabilizes at that $m = 5$ for VB_s and D_{st}. Therefore, $m = 5$ and $\tau = 15$hr was used for the computation of MLE at different categories of geomagnetic storm, while $m = 3$ and $\tau = 1$hr are applied for the computation of ApEn values.

The results of MLE (bar plot) and ApEn (stem plot) for D_{st} at the month of minor, moderate and major geomagnetic storms activity are shown in Figure 5. During the month of minor geomagnetic storms activity, we notice that the value of MLE ranges between 0.07 and 0.14 for most of the months classified as minor geomagnetic storm. Similarly, the ApEn (stem plot) ranges between 0.59 and 0.83. It is obvious that strong chaotic behaviour with high dynamical complexity are associated with minor geomagnetic storms. During the month of moderate geomagnetic storm activity, (see b part of Figure 5), we observe a reduction in MLE values (0.04~0.07) compared to minor geomagnetic storm periods. Within the observed values of MLE during the month of

289 moderate geomagnetic storms activity, we found a slight rise of MLE in the following months
290 (Mar 2008), (Apr 2011), (Jan 2012, Feb 2012, Apr 2012), (Jul 2015, Aug 2015, Sept 2015,
291 Oct2015, Nov 2015) and (Nov 2016). Also, the ApEn revealed a reduction in values between 0.44
292 and 0.57 at the month of moderate geomagnetic storms activity. The lowest values of ApEn were
293 noticed in the following months: May 2010, Mar 2011, and Jan 2016. During major geomagnetic
294 storm as shown in Figure 5, the minimum and maximum value of MLE is respectively 0.03 and
295 0.04 implying a very strong reduction of chaotic behaviour compared with the month of minor and
296 moderate geomagnetic storm activity. The lowest values of MLE were found in the months of Jul
297 2012, Jun 2013 and Mar 2015. Interestingly, further reduction in ApEn value (0.29~0.40) was as
298 well noticed during this period. Thus, during the month of major geomagnetic storm activity,
299 chaotic behaviour and dynamical complexity subsides significantly.

300 We display in Figure 6, the results of MLE and ApEn computation for the VB_s which has been
301 categorized according to the month of minor, moderate and major geomagnetic storm activity. The
302 values of MLE (bar plot) were between 0.06 and 0.20 for VB_s. The result obtained indicate strong
303 chaotic behaviour with no significant difference in chaoticity during minor, moderate and major
304 geomagnetic storm. Similarly, the results obtained from computation of ApEn (stem plot) for VB_s
305 depict a minimum value of 0.60 and peak value of 0.87 as shown in Figure 6. The ApEn values of
306 VB_s indicates high dynamical complexity response with no significant difference during the
307 periods of the three categories of geomagnetic storm investigated.

308 The test for nonlinearity in the D_{st} signals during the month of minor, moderate and major
309 geomagnetic storms activity was analyzed through the DVV analysis. Shown in Figure 9 is the
310 DVV plot and DVV scatter plot during minor geomagnetic storm for January 2009 and January
311 2014. We found that the DVV plots during the month of minor geomagnetic storms activity reveals

a slight separation between the original and surrogate data. Also, the DVV scatter plots shows a slight deviation from the bisector line between the original and surrogate data which implies nonlinearity. Also, during the month of moderate geomagnetic storm activity, we notice that the DVV plot depicts a wide separation between the original and the surrogate data. Also, a large deviation from the bisector line between the original and the surrogate data was also noticed in the DVV scatter plot as shown in Figure (8) thus indicating nonlinearity. In Figure (9), we display samples of DVV plot and DVV scatter plot during major geomagnetic storm for Oct 2011 and Dec 2015. The original and the surrogate data showed a very large separation in the DVV plot during the month of major geomagnetic storm activity. While the DVV scatter plot depict the greatest deviation from the bisector line between the original and the surrogate data which is also an indication of nonlinearity. The DVV analysis of the VB_s time series during the month of minor, moderate and major geomagnetic storm activity shown in Figures (10-12) revealed a slight separation between the original and surrogate data with no significant difference between the month of minor, moderate and major geomagnetic activity.

4.0 Discussion of Results

4.1 The chaotic and dynamical complexity response in D_{st} at minor, moderate and major geomagnetic storms

Our result shows that the values of MLE for D_{st} during the month of minor geomagnetic storm activity are higher, indicating significant chaotic response during minor geomagnetic stormy periods (bar plot, Figure 5). This increase in chaotic behaviour for D_{st} signals during minor geomagnetic storm may be as a result of asymmetry features in the longitudinal distribution of solar source region for the Corotating Interaction Regions (CIR) signatures responsible for the development of geomagnetic storms (Turner et al. 2006; Kozyra et al. 2006). CIR generated

magnetic storms are generally weaker than ICME/MC generated storms (Richardson and Cane, 2011). Therefore, we suspect that the increase in chaotic behaviour during minor geomagnetic storm is strongly associated with the asymmetry features in the longitudinal distribution of solar source region for the Corotating Interaction Regions (CIR) signatures. For most of these periods of moderate geomagnetic storms, the values of MLE decreases compared to the month of minor geomagnetic storms activity. This revealed that as geomagnetic stormy events build up, the level of unpredictability and sensitive dependence on initial condition (chaos) begin to decrease (Lorentz, 1963; Stogaz, 1994). The chaotic behaviour during the month of major geomagnetic storm decreases significantly compared with the month of moderate geomagnetic storm activity. The reduction in chaotic response during the month of moderate and its further declines at major geomagnetic storm activity may be attributed to the disturbance in the interplanetary medium driven by sheath preceding an interplanetary coronal mass ejection (ICME) or combination of the sheath and an ICME magnetic cloud (Echer et al. 2008; Tsurutani et al. 2003; Meng et al. 2019). Notably, the dynamics of the solar wind-magnetospheric interaction are dissipative chaotic in nature (Pavlos, 2012); and, the electrodynamics of the magnetosphere due to the flux of interplanetary electric fields had a significant impact on the state of the chaotic signatures. For instance, the observation of strong chaotic behaviour during the month of minor geomagnetic storm activity suggests that the dynamics was characterized by a weak magnetospheric disturbance. While the reduction in chaotic behaviour at moderate and major geomagnetic storm period reveals the dynamical features with regards to when a strong magnetospheric disturbance begins to emerge. Therefore, our observation of chaotic signatures at different categories of geomagnetic storm has potential capacity to give useful diagnostic information about monitoring space weather events. It is important to note that the features of D_{st} chaotic behaviour at different

categories of geomagnetic storm has not been reported in the literature. For example, previous study of Balasis et al. (2009, 2011) investigate dynamical complexity behaviour using different entropy measures and revealed the existence of low dynamical complexity in the magnetospheric dynamics and attributed it to ongoing large magnetospheric disturbance (major geomagnetic storm). The work of Balasis et al. (2009, 2011) where certain dynamical characteristic evolved in the D_{st} signal was revealed was limited to one year data (2001). It is worthy to note that the year 2001, according to sunspot variations is a period of high solar activity during solar cycle 23. It is characterized by numerous and strong solar eruptions that were followed by significant magnetic storm activities. This confirms that on most of the days in year 2001, the geomagnetic activity is strongly associated with major geomagnetic storm. The confirmation of low dynamical complexity response in the D_{st} signal during major geomagnetic storm agree with our current study. However, the idea of comparing the dynamical complexity behaviour at different categories of geomagnetic storm and reveal its chaotic features was not reported. This is the major reason why our present investigation is crucial to the understanding of the level of chaos and dynamical complexity involved during different categories of geomagnetic storm. As an extension to the single-year investigation done by Balasis et al. (2009, 2011) during a major geomagnetic storm, we further investigated nine years data of D_{st} that covered minor, moderate and major geomagnetic storm (see Figure 5, stem plots) and unveiled their dynamical complexity behaviour. During major geomagnetic stormy periods, we found that the ApEn values decrease significantly, indicating reduction in the dynamical complexity behaviour. This is in agreement with the low dynamical complexity reported by Balasis et al. (2009, 2011) during a major geomagnetic period. Finally, based on the method of DVV analysis, we found that test of nonlinearity in the D_{st} time series during major geomagnetic storm reveals the strongest nonlinearity features.

4.2 The chaotic and dynamical complexity behaviour in the VB_s as input signals.

The results of the MLE values for VB_s revealed a strong chaotic behaviour during the three categories of geomagnetic storm. Comparing these MLE values during the month of minor to those observed during moderate and major geomagnetic storm activity, the result obtained did not indicate any significant difference in chaoticity (bar plots, Figure 6). Also, the ApEn values of VB_s during the periods associated with minor, moderate and major geomagnetic storm revealed high dynamical complexity behaviour with no significant difference between the three categories of geomagnetic storm investigated. These observation of high chaotic and dynamical complexity behaviour in the dynamics of VB_s may be due to interplanetary discontinuities caused by the abrupt changes in the interplanetary magnetic field direction and plasma parameters (Tsurutani et al. 2010). Also, the indication of high chaotic and dynamical complexity behaviour in VB_s signifies that the solar wind electric field is stochastic in nature. The DVV analysis for VB_s revealed nonlinearity features with no significant difference between the month of minor, moderate and major geomagnetic storm activity. It is worth mentioning that the dynamical complexity behaviour for VB_s is different from what was observed for D_{st} time series data. For instance, our results for D_{st} times series revealed that the chaotic and dynamical complexity behaviour of the magnetospheric dynamics are high during minor geomagnetic storm, reduce at moderate geomagnetic storm and further decline during major geomagnetic storm. While the VB_s signal revealed a high chaotic and dynamical complexity behaviour at all the categories of geomagnetic storm period. Therefore, these dynamical features obtained in the VB_s as input signal and the D_{st} as the output in describing the magnetosphere as a non-autonomous system further support the finding of Donner et al. (2019) that found increased or not changed in dynamical complexity behaviour for VB_s and low dynamical complexity behaviour during storm using recurrence

method. Thus, suggesting that the magnetospheric dynamics is nonlinear and the solar wind dynamics is consistently stochastic in nature.

5.0 Conclusions

This work has examined the magnetospheric chaos and dynamical complexity behaviour in the disturbance storm time (D_{st}) and solar wind electric field (VB_s) as input during different categories of geomagnetic storm. The chaotic and dynamical complexity behaviour at the month of minor, moderate and major geomagnetic storm activity for solar wind electric field (VB_s) as input and D_{st} as output of the magnetospheric system were analyzed for the period of 9 years using nonlinear dynamics tools. Our analysis has shown a noticeable trend of these nonlinear parameters (MLE and ApEn) and the categories of geomagnetic storm (minor, moderate and major). The MLE and ApEn values of the D_{st} have indicated that the chaotic and dynamical complexity behaviour are high during the month of minor geomagnetic storm, low during moderate geomagnetic storm and further reduced during major geomagnetic storm activity. The values of MLE and ApEn obtained from VB_s indicate that chaotic and dynamical complexity are high with no significant difference during the periods of minor, moderate and major geomagnetic storm. Finally, the test for nonlinearity in the D_{st} time series during major geomagnetic storm reveals the strongest nonlinearity features. Based on these findings, the dynamical features obtained in the VB_s as input and D_{st} as output of the magnetospheric system suggest that the magnetospheric dynamics is nonlinear and the solar wind dynamics is consistently stochastic in nature.

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Declaration of Interest statement

The authors declare that there is no conflicts of interest.

References

Baker, D.N., Klimas, A.J., (1990). The evolution of weak to strong geomagnetic activity: An interpretation in terms of deterministic chaos. J. GeoPhys. Res. Letts. Vol. 17, No. 1, PP. 41-44.

Balasis, G., Daglis, I.A., Anastasiadia, A., Eftaxias, K., (2011). Detection of dynamical complexity changes in Dst time series using entropy concepts and rescaled range analysis. W.Liu, M. Fujimoto (eds.), The Dynamics Magnetosphere, IAGA Special Sopron Book Series 3, doi: 10.1007/978-94-007-0501-2_12, Springer Science+Business Media B.V. 2011.

Balasis, G., I.A. Daglis, C. Papadimitriou, M. Kalimeri, A. Anastasiadis, K. Eftaxias (2009). Investigating dynamical complexity in the magnetosphere using various entropy measures, J.Geophys.Res., 114, A0006, doi: 10.1029/2008JA 014035.

Balasis, G., I.A. Daglis, P. Kapiris, M.Mandea, D. Vassilliadis, K. Eftaxias (2006). From pre-storm activity to magnetic storms: a transition described in terms of fractal dynamics, Ann.Geophys., 24, 3557-3567, www.ann-geophys.net/24/3557/2006.

446 Balikhin, M.A., Boynton, R.J., Billings, S.A., Gedalin, M., Ganushkina, N., Coca, D., (2010). Data
447 based quest for solar wind-magnetosphere coupling function, Geophys.Res.Lett, 37, L24107,
448 doi: 10.1029/2010GL045733.

449 Burton, R.K., McPherron, R.L., Russell, C.T., (1975). An empirical relationship between
450 interplanetary conditions and Dst. Journal of Geophysical Research, Vol.80, No.31.

451 Consolini, G., (2018), Emergence of dynamical complexity in the Earth's magnetosphere, Machine
452 learning techniques for space weather, PP. 177-202, doi: 10.1016/B978-0-12-811788-0.00007-
453 X

454 Cowley, S.W.H., (1995). The earth's magnetosphere: A brief beginner's guide, EOS
455 Trans.Am.GeoPhys.Union, 76, 525.

456 Dessler, A.J., Parker, E.N., (1959). Hydromagnetic theory of magnetic storm. J. GeoPhys. Res, 64,
457 PP 2239-2259.

458 Devi, S.P., Singh, S.B., Sharma, A.S., (2013). Deterministic dynamics of the magnetosphere:
459 results of the 0-1 test. Nonlin. Processes Geophys., 20, 11-18, 2013, www.nonlin-processes-
460 geophys.net/20/11/2013, doi: 10. 5194/npg-20-11-2013.

461 Donner, R.V., Balasis, G., Stolbova,V., Georgiou, M., Weiderman, M., Kurths, J. (2019).
462 Recurrence-based quantification of dynamical complexity in the earth's magnetosphere at
463 geospace storm time scales. Journal of Geophysical Research: Space Physics, 124, 90-108,
464 doi: 10.1029/2018JA025318.

Echer, E., Gonzalez, D., Alves, M.V., (2006). On the geomagnetic effects of solar wind interplanetary magnetic structures. Space Weather, Vol.4, S06001, doi:10.1029/2005SW000200.

Echer, E., Gonzalez, W.D., Tsurutani, B.T., Gonzalez, A. L. C., (2008), Interplanetary conditions causing intense geomagnetic storms ($D_{st} \leq -100nT$) during solar cycle 23 (1996–2006), J. Geophys. Res., 113, A05221, doi:10.1029/2007JA012744.

Fraser, A.M., (1986). Using mutual information to estimate metric entropy, dimension and entropies in chaotic system, Springer-Verlag, 1986, PP: 82-91.

Fraser, A.M., Swinney, H.L., (1986). Independent coordinates for strange attractors from mutual information, Phys.Rev.A 33, 1134-1140.

Gautama, T., Mandic, D.P., Hulle, M.M.V., (2004). The delay vector variance method for detecting determinism and nonlinearity in time series. Physica D, 190, 167-176, doi: 10.1016/j.physd.2003.11.001.

Gonzalez, W.D., Joselyn, J.A., Kamide, Y., Kroehl, H.W., Rostoker, G., Tsurutani, B.T., Vasyliunas, V.M., (1994). What is a geomagnetic storm? J. Geophys. Res.: Space Physics, Vol. 99, issue A4, pg. 5771-5792, doi: 10.1029/93JA02867.

Gonzalez, W.D., Tsurutani, B.T., (1987). Criteria of interplanetary parameters causing intense magnetic storm ($Dst < -100nT$). Planetary and Space Science, https://ntrs.nasa.gov/search.jsp?R=198800068.

484 Horne, R.H., Glauert, S.A., Meredith, N.P., Boscher, D., Maget, V., Heynderickx, D., and Pitcford,
485 D. (2013). Space weather impacts on satellites and forecasting the Earths electron radiation
486 belts with SPACECAST. Space weather, **11,** 169 - 186

487 Imtiaz, A., (2010). Detection of nonlinearity and stochastic nature in time series by delay vector
488 variance method, International journal of Engineering & Technology, Vol. 10, No. 02.

489 Jaksic, V., Mandic, D.P., Ryan, K., Basu, B., Pakrashi V., (2016). A Comprehensive Study of the
490 Delay Vector Variance Method for Quantification of Nonlinearity in dynamical Systems.
491 *R.Soc.OpenSci.*, 2016: 3:150493, http://dx.doi.org/10.1098/rsos.150493.

492 Johnson, J.R., Wing, S., (2005), A solar cycle dependence of nonlinearity in magnetospheric
493 activity. J. Geophys. Res., 110, A04211, doi: 10.1029/2004JA010638.

494 Kannathal, N., M.L. Choo, U. R. Acharya, P. K. Sadasivan (2005). Entropies for detecting of
495 epilepsy in EEG, Computer Methods and Programs in Biomedicine (2005) 80, 187-194,
496 www.intl.elsevierhealth.com/journals/cmpb.

497 Kennel, M.B., R. Brown, H.D.I. Abarbanel (1992). Determining embedding dimension for phase-
498 space reconstruction using a geometrical construction, PHYSICAL REVIEW A, Volume 45,
499 Number 6.

500 Klimas, A.J., Vassilliasdis, D., Baker. D.N., Roberts, D.A., (1996). The organized nonlinear
501 dynamics of the magnetosphere. J. GeoPhys. Res. Vol.101, No. A6, PP 13089-13113.

502 Kozyra, J.U., Crowley, G., Emery, B.A., Fang, X., Maris, G., Mlynczak, M.G., Niciejewski, R.J.,
503 Palo, S.E., Paxton, L.J., Randall, C.E., Rong, P.P., Russell, J.M., Skinner, W., Solomon, S.C.,
504 Talaat, E.R., Wu, Q., Yee, J.H., (2006), Response of the Upper/Middle Atmosphere to Coronal

505 Holes and Powerful High-Speed Solar Wind Streams in 2003, Geophysical Monograph Series
506 167, 10.1029/167GM24.

507 Kugiumtzis, D., (1999). Test your surrogate before you test your nonlinearity, Phys. Rev. E, 60,
508 2808-2816.

509 Lorenz, E.N., (1963). Determining nonperiodic flow. J. Atmos.Sci.,20,130.

510 Love, J.J., Gannon, J.L. (2009). Revised Dst and the epicycles of magnetic disturbance: 1958-2007.
511 Ann.GeoPhys., 27, 3101-3131.

512 Mandic, D.P., Chen, M., Gautama, T., Van Hull, M.M., Constantinides, A., (2007). On the
513 Characterization of the Deterministic/Stochastic and Linear/Nonlinear Nature of Time Series.
514 *Proc.R.Soc*, 2008: A464, 1141-1160, doi: 10.1098/rspa. 2007.0154.

515 Mane, R., (1981). On the dimension of the compact invariant sets of certain nonlinear maps, D.Rand
516 and L.S.Young, eds, 1981.

517 Mckinley, R.A., Mclntire, L.K., Schmidt, R., Repperger, D.W., Caldwell, J.A., (2011). Evaluation
518 of Eye Metrics as a Detector of Fatigue. Human factor, 53 (4): 403-414, doi:
519 10.1177/0018720811411297.

520 Mendes, O., Dominques, M.O., Echer, E., Hajra, R., Menconi, V.E., (2017), Characterization of
521 high-intnesity, long-duration continuous auroral activitity (HILDCAA) events using
522 recurrence quantification analysis. Nonlin. Processes Geophys.,24,407-417, doi:10.5194/npg-
523 24-407-2017.

524 Meng, X., Tsurutani, B. T., Mannucci, A. J. (2019). The solar and interplanetary causes of
525 superstorms (minimum Dst ≤ −250 nT) during the space age. Journal of Geophysical Research:
526 Space Physics, 124, 3926–3948. https://doi.org/10.1029/2018JA026425.

527 Millan, H., Gharbarian-Alavijeh, B., Garcia-Fornaris, I., (2010). Nonlinear dynamics of mean daily
528 temperature and dewpoint time series at Babolsar, Iran, 1961-2005. Elsevier, Atmospheric
529 Research 98 (2010) 89-101, doi: 10.1016/j.atmosres.2010.06.001.

530 Moore, C., Marchant, T., (2017). The approximate entropy concept extended to three dimensions
531 for calibrated, single parameter structural complexity interrogation of volumetric images.
532 Physics in Medicine & Biology, 62(15).

533 Oludehinwa, I.A., Olusola, O.I., Bolaji, O.S., Odeyemi, O.O., (2018). Investigation of nonlinearity
534 effect during storm time disturbance, Adv. Space. Res, 62 (2018) 440-456, doi:
535 10.1016/j.asr.2018.04.032.

536 Omkar, P.T., Verma, P.L., (2013). Solar features and solar wind plasma parameters with
537 geomagnetic storms during the period of 2002-2006. Indian Journal of Applied Research,
538 Vol.3, Issue.5, ISSN-2249-555X.

539 Pavlos, G.P., (1994). The magnetospheric chaos: a new point of view of the magnetospheric
540 dynamics. Historical evolution of magnetospheric chaos hypothesis the past two decades.
541 Conference Proceeding of the 2nd Panhellenic Symposium held in Democritus University of
542 Thrace, Greece, 26-29, April, edited 1994.

543 Pavlos, G.P., (2012). Magnetospheric dynamics and Chaos theory

544 Pavlos, G.P., Athanasiu, M.A., Diamantidis, D., Rigas, A.G., Sarri, E.T., (1999). Comments and
545 new results about the magnetospheric chaos hypothesis. Nonlinear Processes in Geophysics
546 (1999) 6: 99-127.

547 Pavlos, G.P., Rigas, A.G., Dialetis, D., Sarris, E.T., Karakatsanis, L.P., Tsonis, A.A., (1992).
548 Evidence of chaotic dynamics in the outer solar plasma and the earth magnetosphere. Chaotic
549 dynamics: Theory and Practice, Edited by T. Bountis, Plenum Press, New York, Page. 327-
550 339, doi:10.1007/978-1-4615-3464-8_30.

551 Pincus, S.M., (1991). Approximate entropy as a measure of system complexity, Proc.Natl.Acad.Sci.
552 USA, Vol.88, PP. 2297-2301.

553 Pincus, S.M., Goldberger, A.L., (1994). Physiological time series analysis: what does regularity
554 quantify, The American Journal of Physiology, 266 (4): 1643-1656.

555 Pincus, S.M., Kalman, E.K., (2004). Irregularity, volatility, risk, and financial market time series,
556 Proceedings of the National Academy of Sciences, 101 (38): 13709-13714, doi:
557 10.1073/pnas.0405168101.

558 Price, C.P., Prichard, D., Bischoff, J.E., (1994). Nonlinear input/output analysis of the auroral
559 electrojet index. Journal of Geophysical Research, Vol.99, No: A7, PP: 227-238.

560 Price,C.P., Prichard, D., (1993). The Non-linear response of the magnetosphere: 30 October, 1978.
561 Geophysical Research Letters, Vol.20.

562 Richardson, I.G., Cane, H.V., (2011), Geoeffectiveness (D_{st} and K_p) of interplanetary coronal
563 mass ejections during 1995-2009 and implication for storm forecasting, Space Weather, 9,
564 S07005, doi:10.1029/2011SW000670.

565 Russell, C.T., (2001). Solar wind and Interplanetary Magnetic Field: A Tutorial. Space Weather,
566 Geophysical Monograph 125, Page: 73-89.

567 Russell, C.T., McPherron, R.L., Burton, R.K. (1974). On the cause of geomagnetic storms,
568 J.GeoPhys.Res., 79, 1105-1109.

569 Shujuan G., Weidong, Z., (2010). Nonlinear feature comparison of EEG using correlation
570 dimension and approximate entropy, 3rd international conference on biomedical engineering
571 and informatics.

572 Strogatz, S.H., (1994), Nonlinear dynamics and chaos with Application to physics, Biology,
573 chemistry and Engineering, New York, John Wiley & Sons.

574 Sugiura, M. (1964). Hourly Values of equatorial Dst for the IGY, Ann.Int. GeoPhys. Year, 35, 9-
575 45.

576 Takens, F., (1981). Detecting Strange Attractors in Turbulence in Dynamical Systems, D.*Rand &*
577 *L.Young Eds*, 1981: 898, 366-381.

578 Theiler, J., Eubank, S., Longtin, A., Galdrikian, B., Farmer, J.D., (1992), Testing for nonlinearity
579 in time series: The method of surrogate data, Physica D, 58, 77.

580 Tsurutani, B.T., Gonzalez, W. D., Tang, F., Akasofu, S.I., Smith, E.J., (1988), Origin of
581 Interplanetary Southward Magnetic Fields Responsible for Major Magnetic Storms Near Solar
582 Maximum (1978-1979), Journal of Geophysical Research, Vol. 93, No. A8, Pages 8519-8531,
583 Paper number 7A9404, 0148-0227/88/007 A-9404$05.00.

584 Tsurutani, B.T., Gonzalez, W.D., Lakhina, G.S. Alex, S., (2003). The extreme magnetic storm of
585 1-2 September 1859, J. Geophys. Res. 108(A7), doi: 10.1029/2002JA009504.

586 Tsurutani, B.T., Lakhina, G.S., Hajra, R., (2020), The physics of space weather/solar-terrestrial
587 physics (STP): what we know now and what the current and future challenges are, Nonlin.
588 Processes Geophys., 27, 75–119, doi:10.5194/npg-27-75-2020.

589 Tsurutani, B.T., Lakhina, G.S., Verkhoglyadova, O.P., Gonzalez, W.D., Echer, E., Guarnieri, F.L.,
590 (2010). A review of interplanertary discontinuities and their geomagnetic effects. Journal of
591 Atmospheric and Solar-Terrestrial Physics, doi: 10.1016/j.jastp.2010.04.001.

592 Tsutomu, N., (2002). Geomagnetic storms. Journal of communications Research Laboratory, Vol.
593 49, No.3.

594 Turner, N.E., Mitchell, E.J., Knipp, D.J., Emery, B.A., (2006), Energetics of magnetic storms
595 driven by corotating interaction regions: a study of geoeffectiveness, Geophysical Monograph
596 Series 167, 10.1029/167GM11.

597 Unikrishnan, K., (2008). Comparison of chaotic aspect of magnetosphere under various physical
598 conditions using AE index time series. Ann. Geophys., 26, 941-953, www.ann-
599 geophys.net/26/941/2008.

600 Unikrishnan, K., Ravindran, S., (2010). A study on chaotic behaviour of equatorial/low latitude
601 ionosphere over indian sub-continent, using GPS-TEC time series, J. Atmos. Sol. Ter. Phys.,
602 72, 1080-1089.

603 Valdivia, J.A., Rogan, J., Munoz, V., Gomberoff, L., Klimas, A., Vassilliasdis, D., Uritsky, V.,
604 Sharma, S., Toledo, B., Wastaviono, L. (2005). The magnetosphere as a complex system. Adv.
605 Space. Res, 35, 961-971.

606 Valdivia, J.A., Sharma, A.S., Papadopoulos, K., (1996). Prediction of magnetic storms by nonlinear
607 models. Geophysical Research Letters, 23(21), 2899-2902, doi: 10.1029/96GL02828.

608 Vassiliadis, D., (2006). Systems theory for geospace plasma dynamics, Rev.Geophys., 44, RG2002,
609 doi: 10.1029/2004RG000161.

610 Vassiliadis, D., Klimas, A.J., Valdivia, J.A., Baker, D.N., (1999). The geomagnetic response as a
611 function of storm phase and amplitude and solar wind electric field. Journal of Geophysical
612 Research, 104(A11), 24957-24976, doi: 10.1029/1999JA900185.

613 Vassilliadis, D., Sharma, A.S., Papadopoulos, K., (1991). Lyapunov exponent of magnetospheric
614 activity from AL time series. J. GeoPhys. Letts, Vol. 18, No.8, PP. 1643-1646.

615 Vassilliadis, D.V., Sharma, A.S., Eastman, T.E., Papadopoulou, K., (1990). Low-dimensional
616 chaos in magnetospheric activity from AE time series. J. GeoPhys.Res.Lett, 17, 1841-1844.

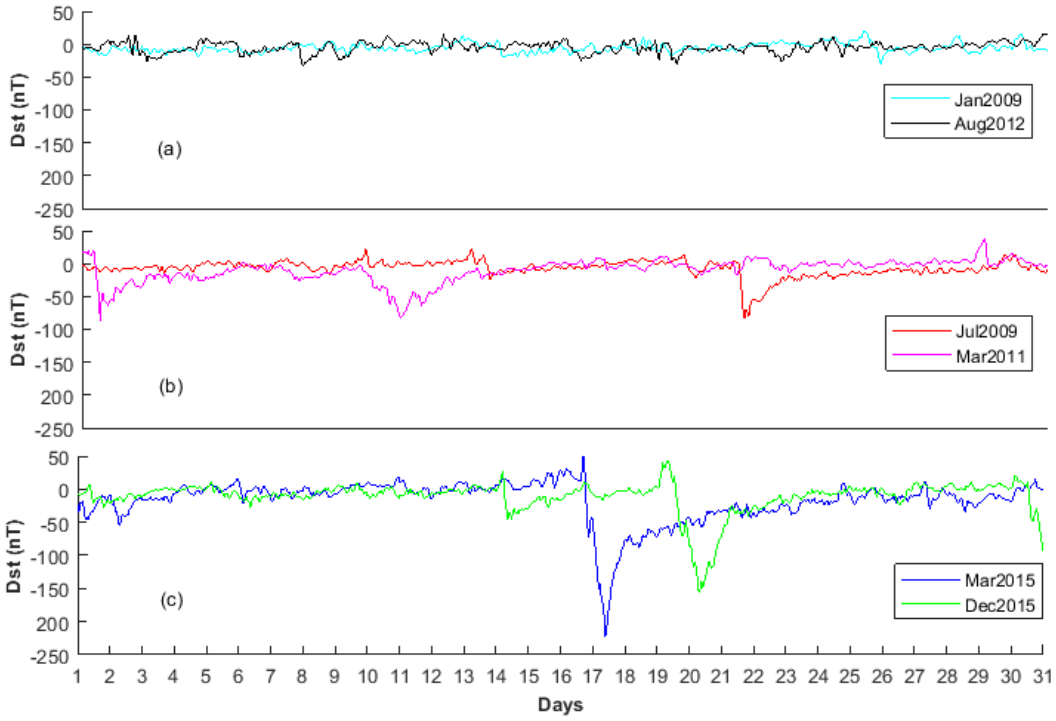
617 Wallot, S., Monster, D., (2018). Calculation of Average Mutual Information (AMI) and False
618 Nearest Neigbhours (FNN) for the estimation of embedding parameters of multidimensional
619 time series in MATLAB. Front. Physchol. 9:1679, doi: 10.3389/fpsyg.2018.01679.

620 Watari, S., (2017). Geomagnetic storms of cycle 24 and their solar sources, Earth, Planets and
621 Space, PP: 69:70, doi: 10.1186/s40623-017-0653-z.

622 Wolf, A., Swift, J. B., Swinney, H. L., and Vastano, J. A. (1985). Determining Lyapunov exponents
623 from a time series, *Physica D, 16*, 285–317, doi:10.1016/0167-2789(85)90011-9.

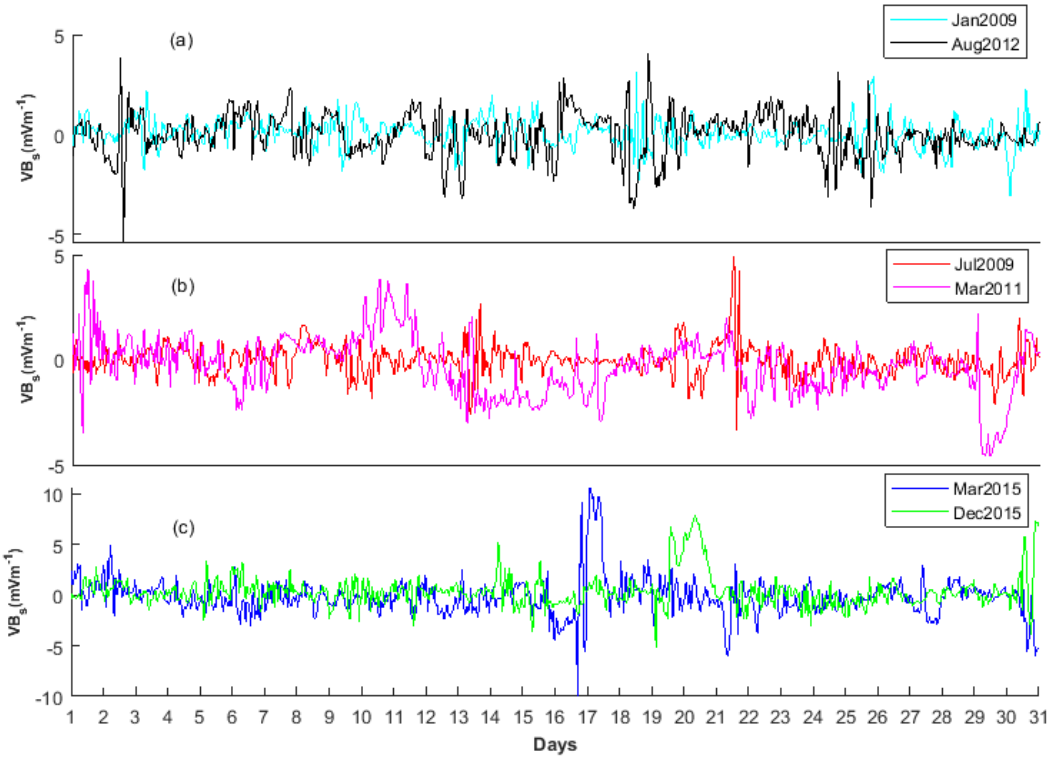
624 Zhang, J., Dere, K.P., Howard, R.A., Bothmer, V., (2002), Identification of solar sources of major
625 geomagnetic storms between 1996 and 2000. Astrophysical Journal, 582:520-53.

626



627

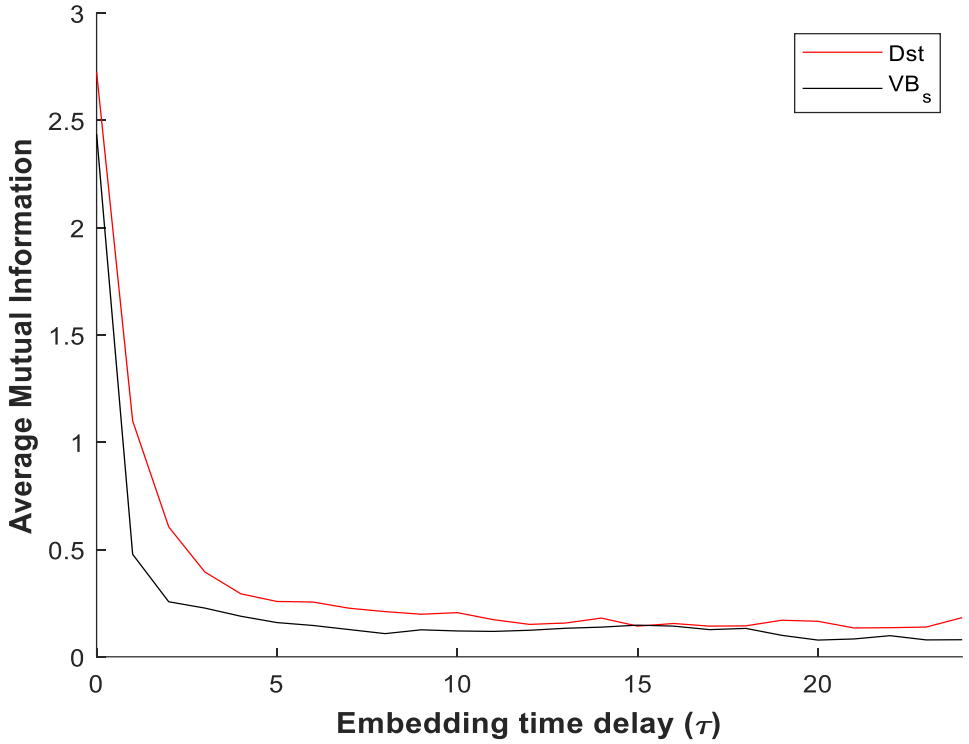
628 Figure 1: Samples of Dst signals classified as (a) Month of Minor, (b) Month of Moderate and (c)
629 Month of Major geomagnetic storm activity



630

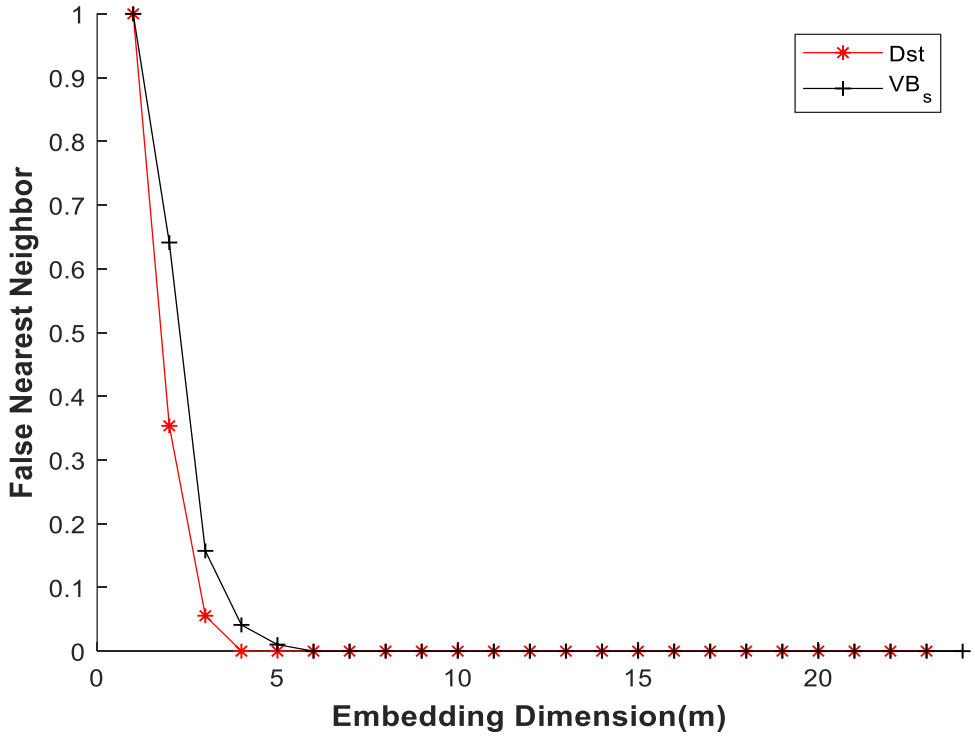
631 Figure 2: Samples of (VB_s) during (a) Month of Minor, (b) Month of Moderate and (c) Month of
632 Major geomagnetic storm activity.

633



634

635 Figure 3: The plot AMI against embedding time delay (τ)



636

637 Figure 4: The plot of FNN against embedding dimension (m)

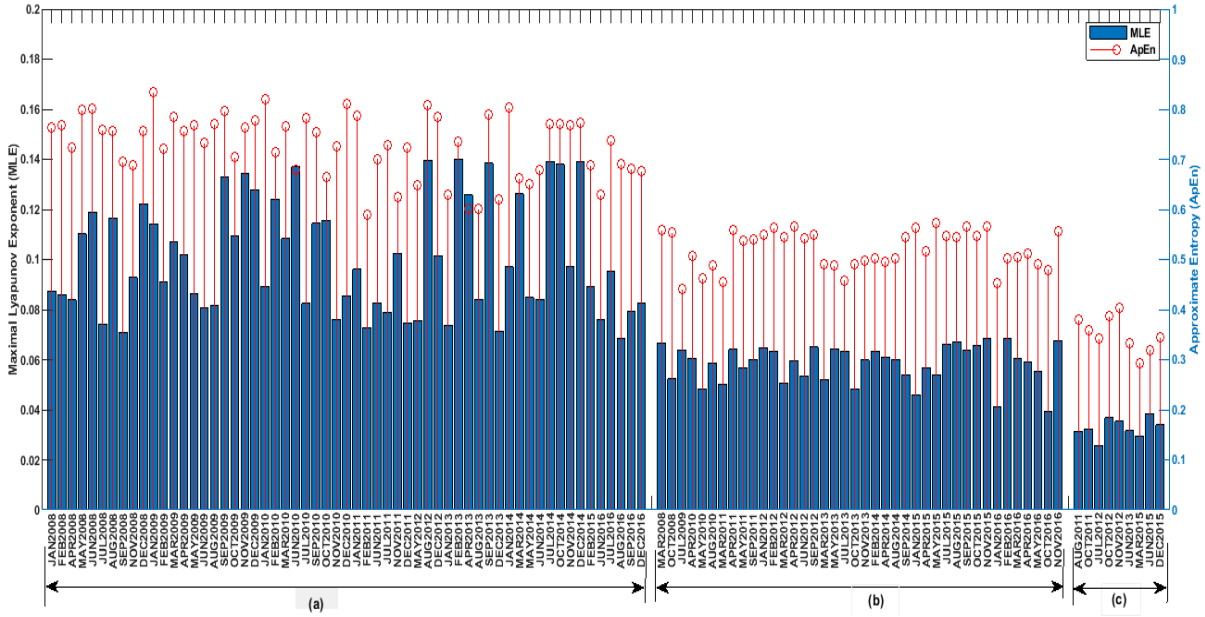


Figure 5: The MLE (bar plot) and ApEn (stem plot) of Dst at: (a) Month of Minor, (b) Month of Moderate and (c) Month of Major geomagnetic storm activity

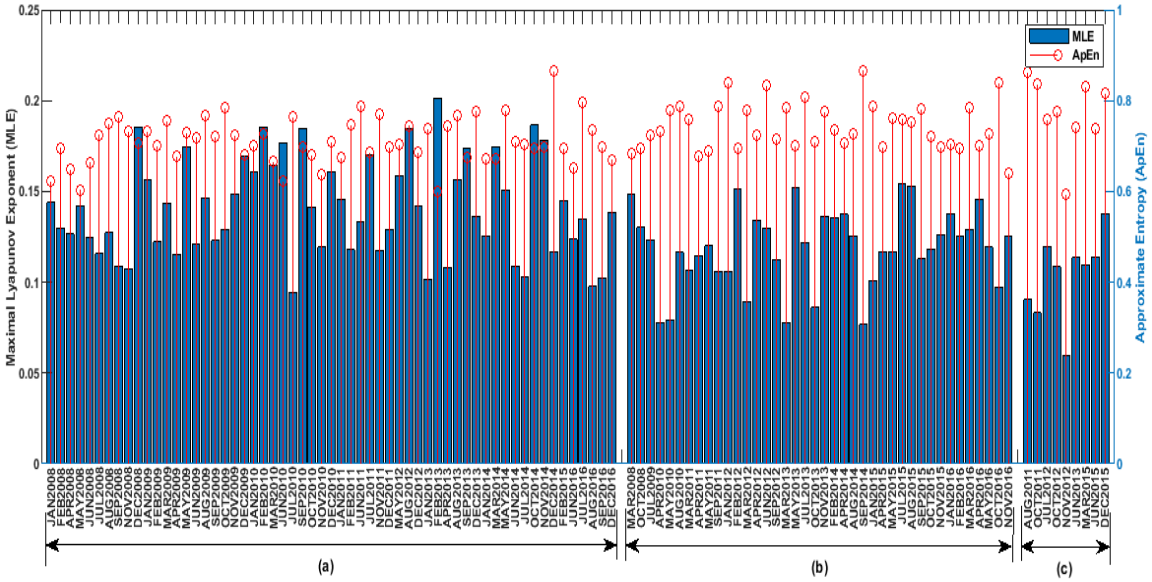
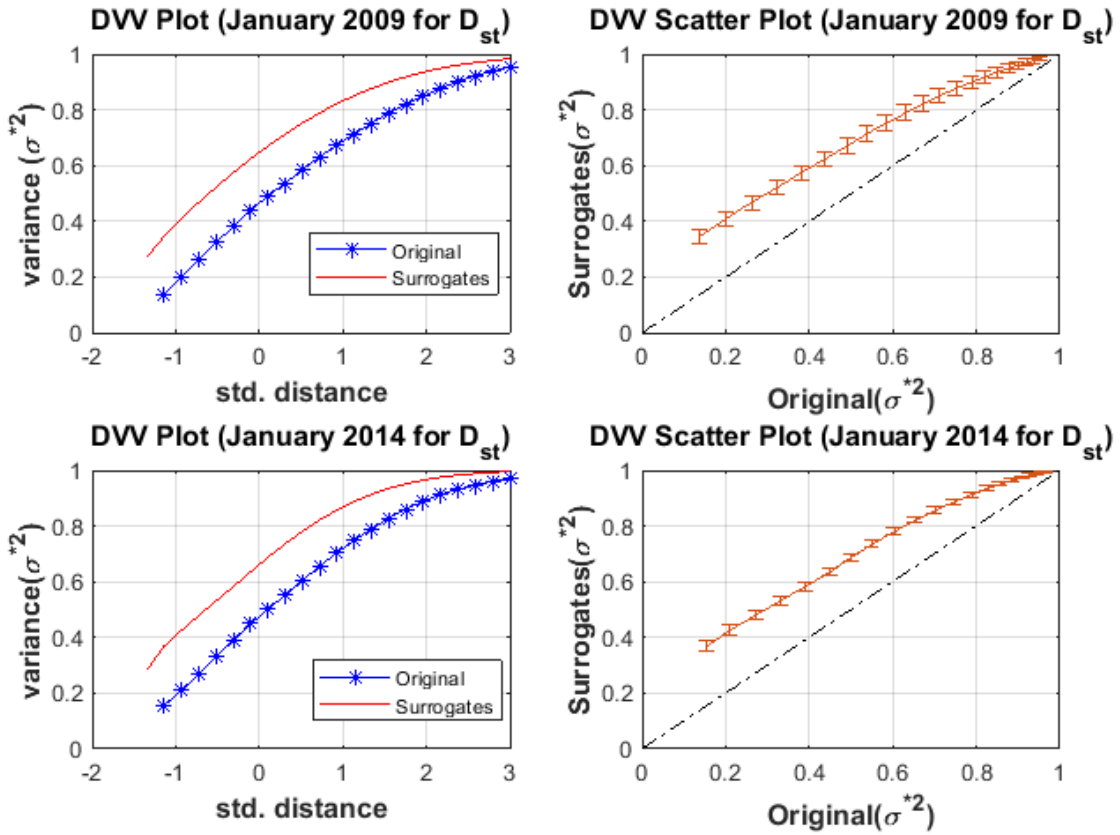


Figure 6: The MLE (bar plot) and ApEn (stem plot) of solar wind electric field (VB_s) during: (a) Month of Minor, (b) Month of Moderate and (c) Month of Major geomagnetic storm activity.



664

665 Figure 7: The DVV plot and Scatter plot for D_{st} during the month of minor geomagnetic storm
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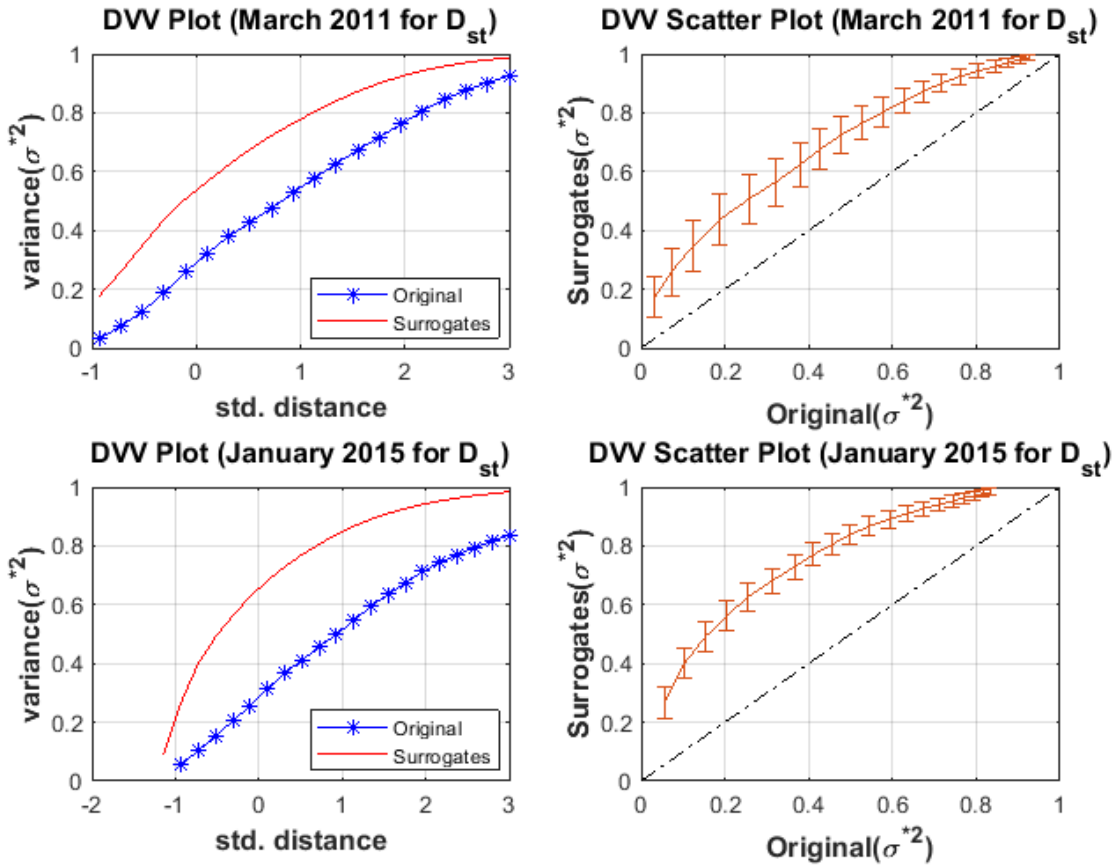

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668

669 Figure 8: The DVV plot and Scatter plot for D_{st} during the month of moderate geomagnetic storm
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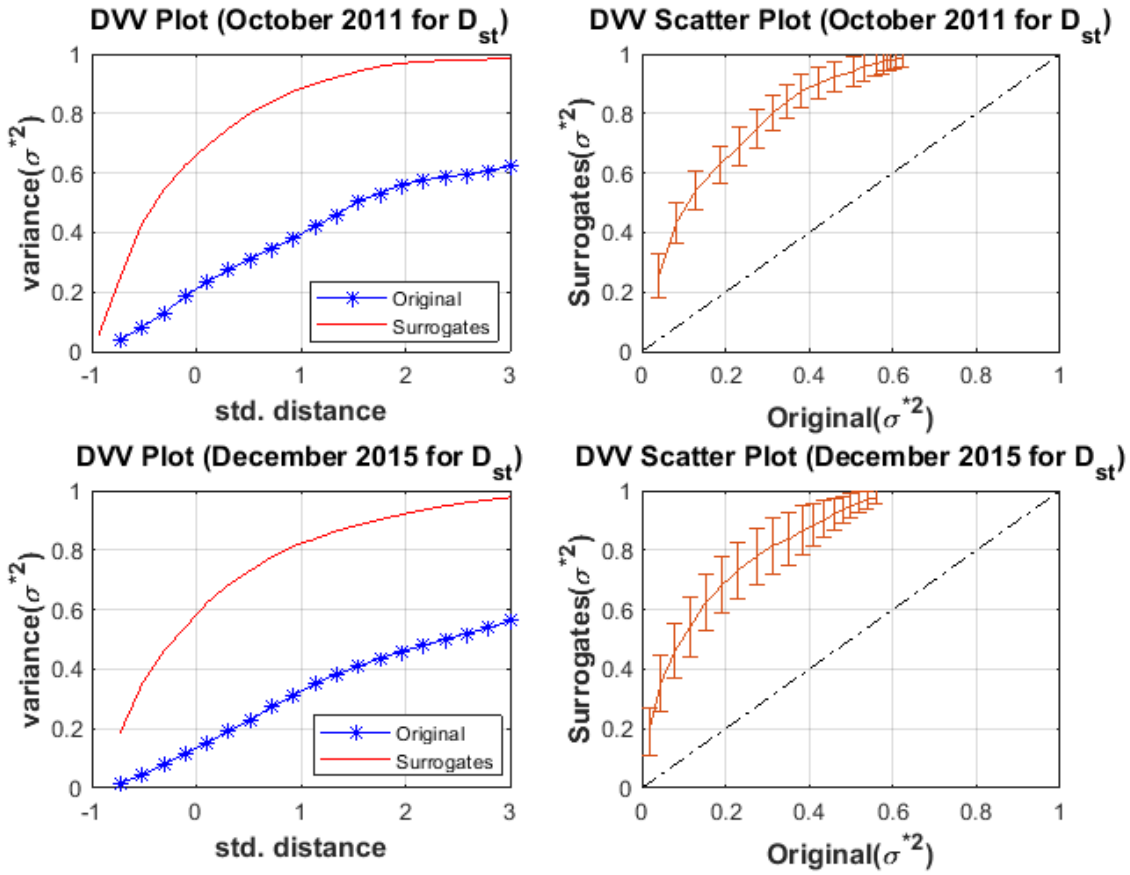
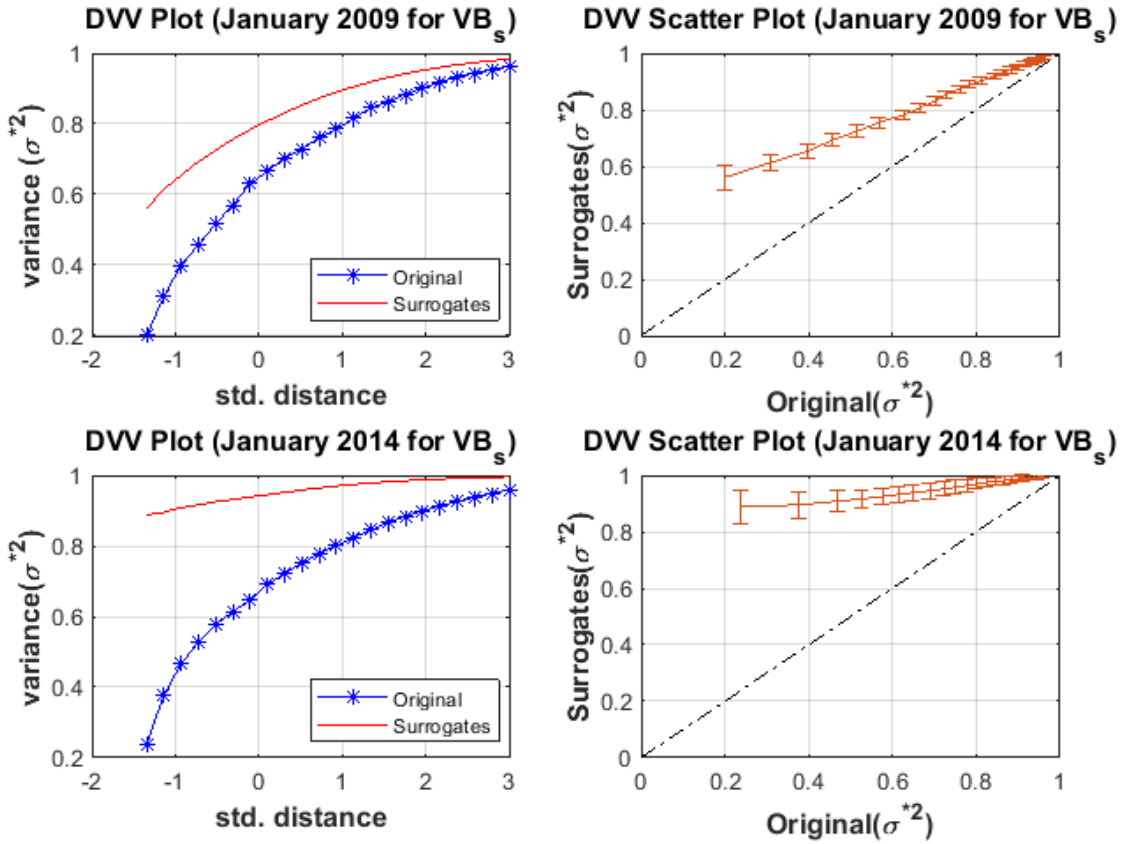
671



672

673 Figure 9: The DVV plot and Scatter plot for D_{st} during the month of major geomagnetic storm
674 for October 2011 and December 2015.

675

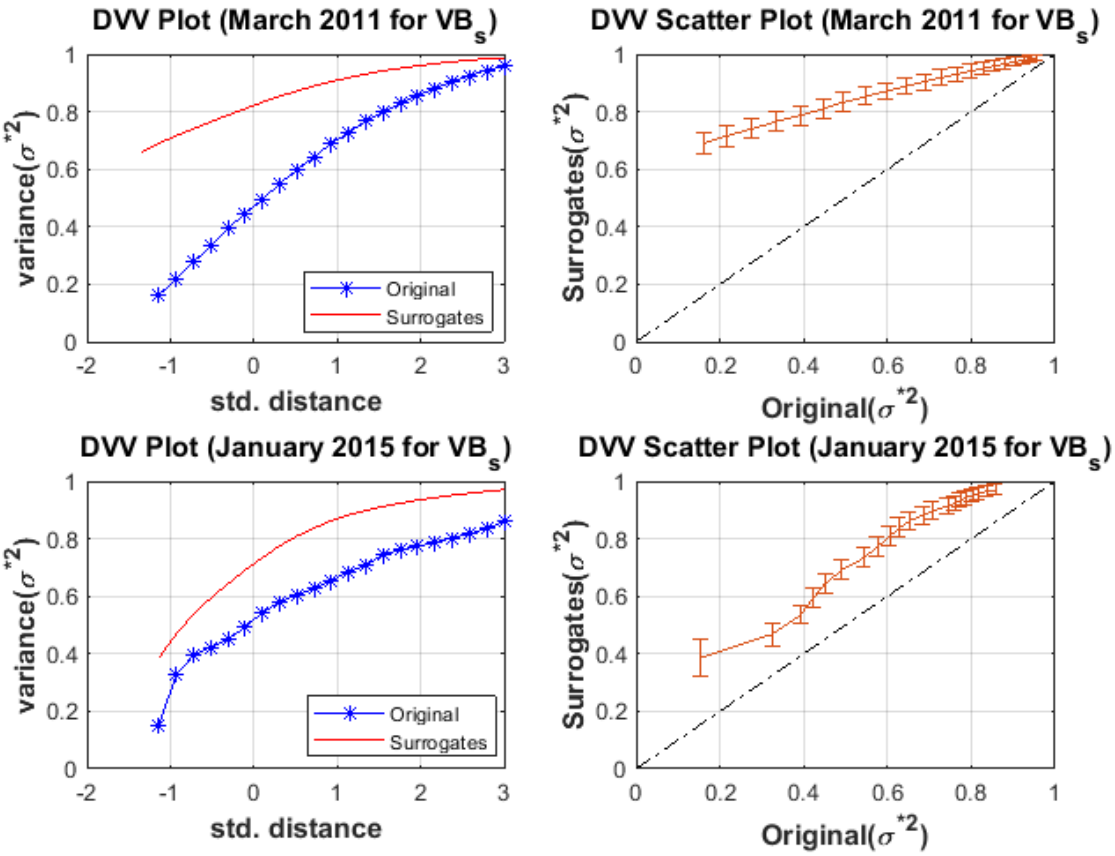


676

677 Figure 10: The DVV plot and Scatter plot for VB_s during the month of minor geomagnetic storm
678 for January 2009 and January 2014.

679

680

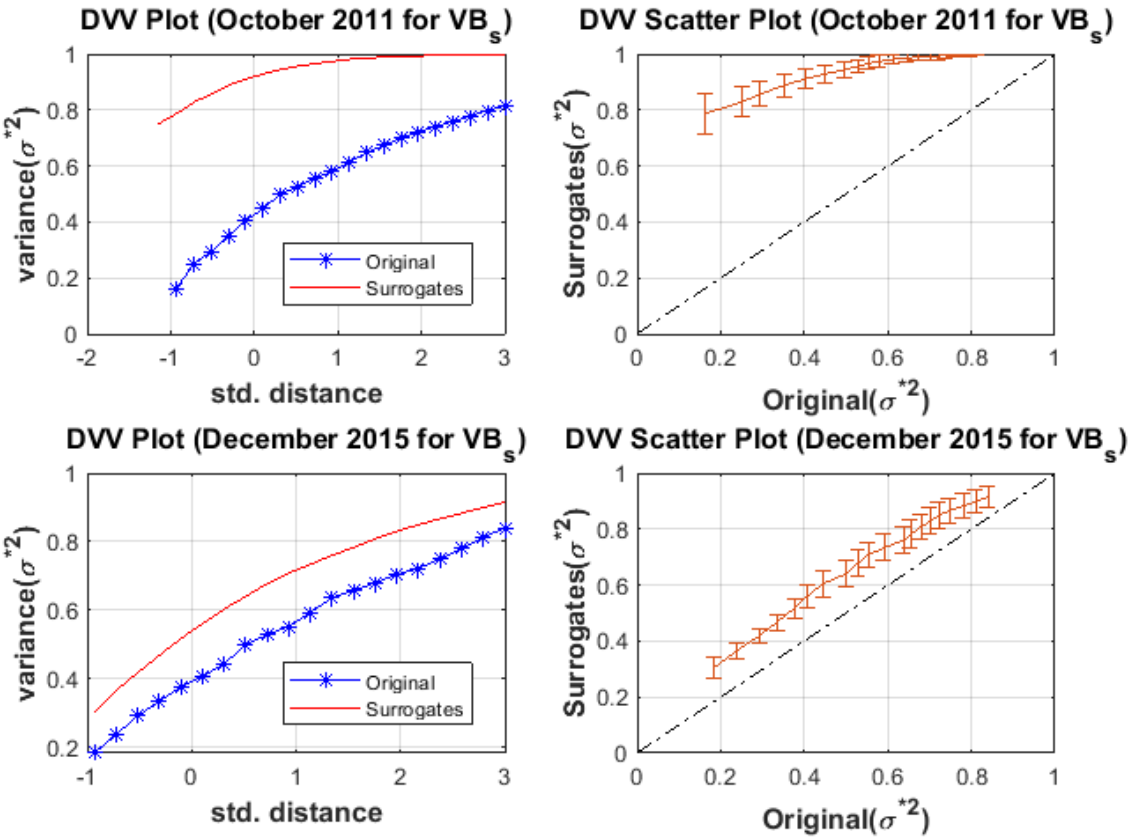


681

682 Figure 11: The DVV plot and Scatter plot for VB_s during the month of moderate geomagnetic
683 storm for March 2011 and January 2015.

684

685



686

687 Figure 12: The DVV plot and Scatter plot for VB_s during the month of major geomagnetic storm
688 for October 2011 and December 2015.

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