



- 1 Singular spectrum and principal component analysis of soil radon
- 2 (Rn-222) emanation for better detection and correlation of seismic
- **induced anomalies**
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Abstact. In the recent years there are several reporting's of anomalous seismic induced 10 temporal changes in soil radon emanation. It is however well known that radon anomalies apart 11 from seismic activity are also governed and controlled by meteorological parameters. This is 12 the major complication which arise for isolating the seismic induced precursory signals. Here 13 in the investigation the soil radon emanations temporal variability at MPGO, Tezpur, is 14 15 scrutinized in the lime light of singular spectrum analysis (SSA). Further prior applying SSA Digital filter (Butterworth low pass) is applied to remove the high frequency quasi periodic 16 component in the time series of soil radon emanation. It was scrutinized that sum of just 9 17 eigenfunctions were sufficient enough for reproducing the prominent characteristics of the 18 19 overall variation. This perhaps also evinces that more significantly produced fluctuations are mostly free from natural variations. The variations in soil temperature was observed to be 20 21 dominated by daily variations similar to radon variation which account to 97.99 % whereas soil pressure accounts for 100 % of the total variance which suggests that daily variations of soil 22 23 radon (Rn-222) emanation are controlled by soil pressure in MPGO, Tezpur during the investigation period followed by soil temperature. The study concludes that SSA eliminates 24 diurnal and semidiurnal components from time series of soil radon emanation for better 25 correlation of soil radon emanation with earthquakes. 26





27 1 Introduction

28 Radon (Rn-222) is a noble gas, a decay product of radium with atomic number, Z=86 and a half-life of nearly 3.8 days. Because of its short decay time, amount changes in its 29 production from the rock is quite evidenced. Ulomov and Mavashev in the year 1971 (Ulomov 30 and Mavashev, 1971) first suggested the correlation of radon concentration with earthquakes. 31 It has been scrutinized that the radon concentration has correlation to earthquakes and volcanic 32 eruptions (Walia et al. 2006, Singh et al. 2005). Significant radon concentration anomalies were 33 also observed prior to the Uttarkashi earthquake of 20th October, 1991; mb~6.5 (Virk and Singh 34 1994). Radon concentrations was monitored in the North West Himalaya for earthquake 35 prediction studies and empirical equation between earthquake magnitude, epicentral distance 36 and precursory time were examined (Kumar et al. 2009). Earthquake prediction depending 37 entirely on precursory phenomena is empirical and comprises many applied problems. Among 38 various geophysical parameters soil radon is preferred and used for earthquake precursory 39 40 studies because of its ease of detectability. Radon in nature has different isotopes: Rn-222 (halflife~3.8 days), Rn-220 (Thoron, half-life~54.5 s) and Rn-219 (half-life~3.92 s). The utmost 41 prominent is Rn-222 because of its longer half-life which is a product of Ra-226 decay process. 42 The Rn-222 emanates from soil or crust either by diffusion or convection and reaches the 43 44 atmosphere. The soil radon emanation concentration is generally assigned to developments of micro-cracks, fissure and fracture due to dilatancy prior to earthquake. This process enhances 45 the transportation of Radon from its original enclosure following the cracks into the 46 atmosphere. Significant radon concentration anomalies were also observed prior to the 47 Uttarkashi earthquake of 20th October, 1991; mb~6.5 (Virk and Singh, 1994). 48

49 North-East India (NE India) is highly vulnerable to earthquake and lies in seismic zone
50 V (BIS 2002) and frequent occurrence of earthquake facilitates the probability of finding
51 precursory phenomena which may lead to a successful prediction in near future. With this





52 objective a Multiparametric Geophysical Observatory (MPGO) in Ouguri Hills (Latitude 26.61°; Longitude 92.78°, Elevation~82m), Tezpur, Assam, India with the installation of 53 several geophysical instruments collecting data simultaneously, portray an opportunity towards 54 55 identification of precursory signatures prior to earthquakes. Earthquake precursory and prediction studies advanced in late 1970s and Heicheng earthquake of 4th February is a land-56 57 mark which was in short-term successfully predicted in 1975 in China (Adams, 1976). The accomplished medium term forecast of M~7.5 earthquake on 6th August, 1988 in northeast 58 59 Indian region (Gupta and Singh, 1988) encouraged to strengthen such studies in India. Another 60 successful short term prediction was done of $(M \ge 4)$ in Koyna region of western India, famous for Reservoir Triggered Seismicity (Verma and Bansal, 2012). 61

62 The study tries to correlate radon emanation in soil gas with earthquakes within the 63 epicentral distance of 100 km of mb > 3.1 from MPGO, Tezpur which is situated in a highly 64 tectonically strained and seismically active region. The major problem arises is the removal of quasi periodic, diurnal (mostly due to temperature) and semidiurnal (mostly due to pressure) 65 components (Kumar et al., 2015). Radon anomalies are governed by seismic activity as well as 66 by meteorological parameters (soil temperature, pressure, rainfall, moisture and even wind for 67 atmospheric radon; Stranden et al., 1984; Kumar et al., 2009; Walia et al., 2005) which in turn 68 makes it more complex for identifying the seismic induced anomalies. Here in the investigation 69 70 characteristics features of temporal soil radon concentrations variability at MPGO, Tezpur is scrutinized by applying singular spectrum analysis (SSA) in concern to the objective of filtering 71 the meteorological parameters on radon emanation. SSA is a relatively innovative and powerful 72 advanced method which has been used across many applied problems for different scientific 73 fields (e.g., Fraedrich, 1986, Serita et al., 2005). The foremost concept of SSA is applying PCA 74 on trajectory matrix acquired from the original time series following time series reconstruction. 75





77 2 Seismotectonics of the region

78 In middle of the active Kopili and Bomdila fault, the MPGO is situated. The Kopili and Bomdila faults comprise Neogene-Quaternary sediments, which directly were deposited 79 over the Archean basement. The Kopili fault zone in an approx is 100 km in width and 300 km 80 in length. It is a NW-SE trending strike-slip fault (Kayal et al., 2006, Bhattacharya et al., 2008, 81 82 2010). The two Precambrian massifs - the Shillong Plateau and the Mikir Hills is delineated by the tectonic disposition of the Kopili fault. MPGO is bounded to the north by the Main 83 Boundary Thrust (MBT) and to the south by the NE-SW trending Belt of Schuppen (Figure 1). 84 85 The Bomdila fault is strike slip fault of about 400 km in length which trends along WNW-ESE 86 direction. The northern portion of the fault mostly lies in the Gondwana, Paleogene and Neogene sediments. This fault is surrounded by the Belt of Schuppen to the east and south by 87 the Mikir massif and to the west. The fault cuts across the Himalayan fold belt towards the 88 89 north (Nandy and Dasgupta, 1991).

The Kopili Fault has produced two large earthquakes (Figure 1) $M_w \sim 7.7$, 1869 event 90 (Figure 1) in the south eastern edge of the fault contravening the Naga-Disang thrust and 91 M_w~7.2, 1943 earthquake which occurred farther to the north of 1869 event within a period of 92 nearly 75 years (Kayal, 2008). It is highly active with strong seismicity discernible down up to 93 depth of about 50 km, and which extends to the Main Central Thrust (MCT) in the Bhutan 94 Himalaya. Even if MCT is dormant (Ni and Barazangi, 1984), intense activity is observed at 95 the region where Kopili Fault meets the MBT and MCT (Nandy, 2001, Kayal et al., 2010, 96 2012). This is demonstrated by the August 19, 2009 Earthquake ($M_w \sim 5.1$) in the Assam Valley 97 that occurred in the center of the Kopili fault zone and the September 21, 2009 strong Bhutan 98 Himalaya Earthquake ($M_w \sim 6.3$) that occurred at the northern end of the Kopili fault where it 99 connects with the MCT (Kayal et al., 2012). Both the earethquakes are shallow focus (depth ~ 100 101 10 km) showing right lateral strike-slip faulting (Kayal et al., 2010) which suggests that the





102 Kopili fault zone is experiencing compressional stress due to the Indo-Burma arc and Himalyan arc to the east and the north respectively which is characterized by transverse tectonics. The 103 Bomdila Fault inter-weaves across three major tectonic domains of the NE-India, namely 104 105 MCT, MBT and Naga-Disang thrust along NW-SE direction. The earthquake events along 106 Bomdila fault occur in a diffused pattern having post-collisional intracratonic characteristics 107 (Nandy and Dasgupta, 1991). It is observed that, the Upper Brahmaputra Valley stretching between the Bomdila Fault and almost near NW-trending Mishmi Thrust in the northeast is 108 109 seismically dormant, and is recognized as the Assam Gap (Khattri, 1983).

110 **3** Method and techniques

111 3.1 Soil radon (Rn-222) time series

112 A BMC2 barasol manufactured by Algade is into operation for soil radon emanation time series data in MPGO for earthquake prediction and precursory studies. Soil gas radon 113 114 emanation every 15 minutes is being continuously monitored. The barasol probe is kept fitted inside a plastic tube (length 1.5m and diameter of 0.0635m) with open bottom dug inside the 115 ground in such a way that the detection unit (detector sensitivity-0.02 pulses/h for 1 Bq/m³ and 116 saturation volumetric activity- $3MBq/m^3$) which is at the bottom of the probe lies 1m from the 117 ground level. A silicon alpha detector detects the radon gas which enters the detection chamber 118 when it emanates from the soil. The radon pass in a detection volume over three cellulose filters 119 120 which allows to trap all the solid daughter products of radon. The sensor is a fixed silicon detector with a depleted depth of 100µm and 400 mm² of sensitive area. It performs the 121 counting by atomic spectrometry of radon (Rn-222) and daughter products created in the 122 detection volume (with window customized between 1.5 MeV and 6 MeV). The probe system 123 124 is embedded with soil pressure and temperature sensor. The sensor calibration permits the 125 volumic activity of the radon (Rn-222) to be evaluated. In the present investigation the soil





radon emanation temporal variability at MPGO, Tezpur the radon data were prudently checked for no gaps or discontinuous jump. Digital filter (Butterworth) is applied to eliminate the high frequency quasi periodic components form the soil radon time series for better discernibility of seismic induced anomalies.

130 3.2 Singular Spectrum Analysis

131 The SSA results and graphs in the investigation are acquired by using Caterpillar-SSA 3.40 software (Alexandrov and Golyandina, 2004). Window selection rule applied to the time 132 series is one half of the length of the time series to meet the theoretical requirements for the 133 134 investigation (Golyandina, 2010, Khan, 2011, Hassani, 2007). The singular value 135 decomposition (SVD) alogorithm was applied as it is more accurate than QR iteration which 136 are the common most algorithm for solving of eigenvalues and singular value problems (Demmel, 1992). The main objective of SSA is decomposing the original time series into sum 137 138 of series such that each of the component in this sum can be known (either as a trend, periodic or quasi-periodic components) or noise. This is accomplished by decomposition and 139 140 reconstruction. At the first the time series is decomposed following the reconstruction of the original time series (which is without noise). The methodology adopted here is first to 141 embedding a 1-dimension time series say, $Y_T = (y_1, ..., y_T)$ into a multi-dimensional time series 142 $X_{1,...,X_{K}}$ having vectors $X_{i} = (y_{i},...,y_{i+L-1}) \in \mathbb{R}^{L}$ (Golyandina et al., 2001, 2001). Here the value 143 of K = T - L + I. The X_i Vectors are called L-lagged vectors. The embedding depends on the 144 145 window length L, such that $2 \le L \le T$ which results for trajectory matrix (Hankel matrix: diagonal elements i + j = const. are equal) $X = [X_1, ..., X_K] = (X_{ij})_{i,j=1}^{L,K}$. Secondly the Singular 146 Value Decomposition (SVD) of the trajectory matrix is performed to represent it as an addition 147 of bi-orthogonal elementary matrices having rank one. Represented by $\lambda_{I,...,\lambda_{L}}$ which are the 148 Eigen-values of XX' in a descending order of magnitude ($\lambda_l \ge ... \lambda_L \ge 0$) and by $U_l, ..., U_L$ which 149





150	are the orthonormal system (i.e. Ui , $Uj = 0$ for $i \neq j$) is the orthogonality property) and $ U_i =1$,
151	of the eigenvectors of the matrix XX' corresponding to these eigenvalues. (U_i, U_j) is the inner
152	product of the vectors U_i and U_j and $ U_i $ is the norm of the vector U_i . The Set
153	$d = \max(i, \operatorname{such} \operatorname{that} \lambda_i > 0) = \operatorname{rank} X$ (I)
154	If we represent $V_i = X' U_i / \sqrt{\lambda_i}$, then SVD of the trajectory matrix can be represented
155	as:
156	$X = X_I + \dots + X_d \tag{II}$
157	Here $X_i = \sqrt{\lambda_i U_i V_i}$ $(i = 1,,d)$.
157 158	Here $X_i = \sqrt{\lambda_i U_i V_i}$ ($i = 1,,d$). Thirdly the series reconstruction is accomplished by grouping, to split the elementary
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158 159	Thirdly the series reconstruction is accomplished by grouping, to split the elementary matrices (X_i) into various groups and addition of the matrices within each and every group. Say

$$X = X_{l_1} + \dots + X_{l_m} \tag{III}$$

The eigentriple grouping is the process of choosing the sets $I_1,...,I_m$. Finally diagonally averaging transfers each I matrix into a time series, which is an additive component of the initial series Y_T . If z_{ij} is an element of the matrix Z. Then k-th term of the produced series is acquired by averaging z_{ij} for i, j such that i + j = k + 2. This procedure is known as diagonal averaging (Hankelization) of the matrix Z. The Hankelization of a matrix Z is the Hankel matrix HZ which is the trajectory matrix corresponding to the series obtained from diagonal averaging. Hankelization is a best logical technique that HZ matrix is the nearest to Z among all





171 corresponding size of Hankel matrices (Golyandina et al., 2001). Now applying the172 Hankelization technique to each and every matrix components in equation III, we get

173
$$X = \tilde{X}_{l_1} + \ldots + \tilde{X}_{l_m}$$
(IV)

174 Here $X_{II} = HX$ corresponding to initial series $Y_T = (y_1, \ldots, y_T)$ decomposition into a 175 sum of m series as

176
$$y_t = \sum_{k=1}^m y_t^{(k)}$$
 (V)

177 Here
$$y_t^{(k)} = y_1^{(k)} + \ldots + y_T^{(k)}$$
 corresponds matrix X_{lj} .

178 4 Results

The average value of radon for a period of six month from April 2017-September 2017 179 180 was found to be in the range 55-117 kBq/m³. The average emanation of soil-gas radon at 181 MPGO for April, May, June, July, August and September 2017 is reported to be 55.94, 93.11, 109.12, 117.69, 101.45, 92.34 (kBq/m3) respectively with standard deviation (Std.) of 21.3, 182 183 28.53, 19.07, 28.09, 25.86, 18.65 (kBq /m3) respectively. Simultaneously variation of soil temperature and pressure with radon emanation was observed. Usually, radon shows positive 184 correlation with temperature i.e. the soil radon concentration increases with increase in 185 temperature and decreases as temperature decrease. The correlation coefficient (Pearson 186 correlation) between radon and temperature is found to be 0.5 signifying positive correlation, 187 while the correlation coefficient between radon and pressure is found to be -0.5 signifying 188 negative correlation with an average temperature and pressure of 28.60 °C and 991.03 mbar 189 respectively during afore mentioned period. The positive correlation of radon with temperature 190 might be due to the rise in diffusion rate with temperature (Sharma et al. 2000, Singh et al. 191 192 2008). The negative correlation coefficient was found for soil radon and pressure which signify, 193 with the increase in pressure, the radon emanation decreases while with the decrease in pressure





the radon emanation increases. In general, the negative correlation is due to the diffusion 194 process which slows down with increase in pressure, which in turn decreases the radon 195 concentration in the soil. The average value of pressure, temperature, standard deviation, 196 197 percentage (%) correlation coefficient for the observation period is detailed in Table 1. The maximum and minimum temperature observed was 31.24 °C and 23.78 °C i.e. a change of 7.46 198 199 0 C during the period of observation. Simultaneously, the maximum and minimum pressure during the period of observation was 999.32 mbar and 980.72 mbar i.e. a change of 18.6 mbar. 200 Digital filter (Butterworth) is applied to eliminate the high frequency quasi periodic 201 202 components form the soil radon time series for better discernibility of seismic induced 203 anomalies and is represented in Figure 2.

The covariance matrix of the first 9 group of soil radon (Rn-222) time series is 204 205 represented in Figure 3. The singular value decomposition (SVD) to Rn-222 data evinced that first 9 eigenfunctions (Figure 4) when grouped resulted for 99.90 % of the total variance in the 206 individual time series. The eigenfuction group 1 and 2 represents the aperiodic component and 207 group 3 to 9 represented periodic components. The periodic and aperiodic component mostly 208 corresponds to diurnal and semidiurnal variation (Kumar et al., 2015). The decomposed 209 eigenvectors in soil radon time series is grouped into two classes as diurnal and semidiurnal 210 variation. The sum of eigenfuction group 1 and 2 accounts for 98.62 % and group 3 to 9 211 212 accounts for 0.48 % of the variance. Radon variations is governed by daily variations, which accounts to 99.90 % of the total variance in soil gas radon at the MPGO, Tezpur. The Principal 213 214 Component of soil radon (Rn-222) related to the first 9 grouping of eigentriples is represented in Figure 5 and w-correlation matrix for the 9 reconstructed components is represented in 215 Figure 6. 216

The covariance matrix of the first 9 group of soil temperature and pressure time seriesis represented in Figure 7 and Figure 11 respectively. The decomposed eigenfunctions for soil





temperature and pressure time series by applying SSA is represented in Figure 8 and Figure 12 219 respectively. The first 9 eigenfunction group from SVD to atmospheric temperature records 220 accounts nearly about 99.99 % and first 9 eigenfunction group of soil pressure 100 % of the 221 222 variation respectively. It is discernible that first eigenfunction 1 alone itself is capable of 223 producing 100 % variation but the time series is well modeled when first 9 eigenfunction is 224 grouped. This evinces SVD to radon, soil temperature and soil pressure data evince that first 9 eigenfunction accounts >98.00 % of the variance of the individual data series. This is 225 fascinating to observe that sum of first 9 eigenfunction is fairly sufficient to reproduce the 226 227 prominent features of the overall variation. This also suggests that the most significantly 228 produced variations are mostly free from naturally induced variations. Soil pressure variations are dominated by semidiurnal variations by 100 % of the total variation in atmospheric pressure 229 at MPGO, Tezpur. On the other hand soil temperature variations is dominated by daily 230 variations like radon variation which account to 97.99 % whereas soil pressure accounts for 231 100 % of the total variance. This suggest that daily variations of soil radon (Rn-222) emanation 232 233 are controlled by soil pressure in MPGO, Tezpur during the investigation period followed by 234 soil temperature. The Principal Component of soil temperature related to the first 9 grouping of eigentriples is represented in Figure 9 and w-correlation matrix for the 9 reconstructed 235 236 components is represented in Figure 10. The Principal Component of soil pressure related to the first 9 grouping of eigentriples is represented in Figure 13 and w-correlation matrix for the 237 9 reconstructed components is represented in Figure 14. The reconstructed time series for soil 238 239 radon (Rn-222), temperature and pressure is represented is Figure 15 and it's residual in Figure 240 16. It is also discernible that the during the investigation time period the pressure change was more than temperature change which also evinces the variation of soil radon at MPGO, Tezpur 241 242 was more governed by pressure followed by temperature change. The quasi periodic, diurnal 243 (periodic mostly due to temperature) and semidiurnal (aperiodic mostly due to pressure) were





eliminated by Singular Spectrum Analysis in the reconstructed time series by grouping and analyzing the eigenfunctions and principle component of individual time series of Rn-222, temperature and pressure respectively. The soil pressure and temperature were found to be negatively correlated to each other (-0.62) which produces a pseudo effect in the soil radon time series. The grouping and reconstruction of the time series also eliminates these pseudo effect arising due soil pressure and temperature.

250 5 Discussion

251 The reconstructed soil radon time series along with the seismic activity during the 252 investigation period is shown in Figure 17. The hypocentral parameters of the earthquake 253 events found to have correlation with soil radon emanation is listed in Table 2. It was observed 254 that there were certain positive amplitude rise anomalies in radon emanation prior to six out of 9 earthquakes which occurred in the vicinity (100 km radially from MPGO, Tezpur). Increase 255 256 in the soil radon concentration is generally assigned to developments of microcracks, fractures and fissure caused by dilatancy prior to earthquake. This process enhances the transportation 257 258 of radon from its original enclosure following the cracks. The rise in soil radon concentration prior to an earthquake may be due to the strain buildup processes in the area. During this 259 process, very small fractures are developed in the rocks which enhances the contribution of 260 261 radon gas to the soil near the surface of earth (Miklavčić et al., 2008). Three earthquake events were preceded by negative anomalies. The negative anomalies might be due to the 262 circumstance that during the final stage of dilatency model prior to an earthquake the Rn-222 263 emanation can be stable or it can decrease (Tomer, 2016). This is because, during the final 264 stage prior an earthquake, rupture occurs and fluid pressure and stress on rocks is released 265 (Bakhmutov and Groza, 2008). The fluid pressure increases resulting in water level rise and 266 this does not allow the soil gas Rn-222 to escape from the surface which in turn reduces or 267 stabilizes the emanation of Rn-222. Further a decreasing radon anomaly as observed in this 268





269 study may be the result of squeezing effect of compressional stress built up in the rock, which 270 in turn result in soil porosity changes at micro scale. There were certain events which occurred on the same day or just a very short seismic gap of 1 or 2 days. Here in the case earthquake 271 272 with higher magnitude might also be the probable reason for the anomalous behavior of the 273 soil gas radon emanation, as for spatio-temporal clustered earthquakes, the largest magnitude 274 earthquake is presumed to precede the anomalies in radon emanation (Hartmann and Leavy, 2005). Positive as well as negative anomalies were observed prior to 9 events which occurred 275 in the vicinity (100 km radially from MPGO) with in a short span of time. 276

277 6 Conclusion

278 The investigation concludes that digital filter assists in eliminating the high frequency 279 quasi periodic components from the time series. The SSA method helps eliminating the diurnal and semidiurnal fluctuations from soil radon time series for improved correlating and detection 280 281 of the soil radon emanation with seismic activity. The investigation also evinced that radon is dominated by daily variation at MPGO, Tezpur and is controlled by soil pressure followed by 282 283 temperature. It is also concluded that principle component analysis helps in removing the pseudo effect pertaining to simultaneous soil pressure and temperature effect. It was observed 284 that there were certain positive amplitude rise anomalies in radon emanation prior to six events 285 286 out of 9 earthquakes which occurred in the vicinity (100 km radially from MPGO, Tezpur) within a short span of time. The increase of radon emanation with temperature might be the 287 result of increasing diffusion rate with temperature. Three earthquake events were preceded by 288 negative anomalies. The negative anomalies might be due to the circumstance that during the 289 final stage of dilatency model prior to an earthquake the soil gas radon emanation can be stable 290 or it can decrease. This is because, during the final stage prior an earthquake, rupture occurs 291 and fluid pressure and stress on rocks is released. Further a decreasing radon anomaly as 292





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- rock, which in turn changes porosity of soil at micro scale.
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- 299 **References**
- Adams, R. D.: The Haicheng, China, earthquake of 4 February 1975; the first successfully
 predicted major earthquake, Earthquake Engineering & Structural Dynamics, 4, 423 437, 1976.
- Alexandrov, T., and Golyandina, N.: The automatic extraction of time series trend and
 periodical components with the help of the Caterpillar-SSA approach, Exponenta Pro,
 3, 54-61, 2004.
- Bakhmutov, V., and Groza, A.: The dilatancy-diffusion model: new prospects, Proceedings,
 International Conference: Problems of Geocosmos, 7th, St. Petersburg, Russia, May,
 2008, 26-30.
- Bhattacharya, P. M., Kayal, J., Baruah, S., and Arefiev, S.: Earthquake source zones in northeast India: seismic tomography, fractal dimension and b value mapping, in: Seismogenesis and Earthquake Forecasting: The Frank Evison Volume II, Springer, 145-158, 2010.
- Bhattacharya, P. M., Mukhopadhyay, S., Majumdar, R., and Kayal, J.: 3-D seismic structure
 of the northeast India region and its implications for local and regional tectonics,
 Journal of Asian Earth Sciences, 33, 25-41, 2008.
- Demmel, J., and Veselić, K.: Jacobi's method is more accurate than QR, SIAM Journal on
 Matrix Analysis and Applications, 13, 1204-1245, 1992.
- Fraedrich, K.: Estimating the dimensions of weather and climate attractors, Journal of the
 atmospheric sciences, 43, 419-432, 1986.
- Golyandina, N., Nekrutkin, V., and Solntsev, V.: "Caterpillar"-SSA Technique for Analysis of
 Time Series in Economics, New Models of Business: Managerial Aspects and Enabling
 Technology, 198, 2001.
- Golyandina, N., Nekrutkin, V., and Zhigljavsky, A. A.: Analysis of time series structure: SSA
 and related techniques, Chapman and Hall/CRC, 2001.
- Golyandina, N.: On the choice of parameters in singular spectrum analysis and related
 subspace-based methods, arXiv preprint arXiv:1005.4374, 2010.
- Gupta, H.: Medium-term earthquake prediction, Eos, Transactions American Geophysical
 Union, 69, 1620-1630, 1988.





- Hartmann, J., and Levy, J. K.: Hydrogeological and gasgeochemical earthquake precursors–A
 review for application, Natural Hazards, 34, 279-304, 2005.
- Hassani, H.: Singular spectrum analysis: methodology and comparison, 2007.
- Kayal, J., Arefiev, S. S., Baruah, S., Tatevossian, R., Gogoi, N., Sanoujam, M., Gautam, J.,
 Hazarika, D., and Borah, D.: The 2009 Bhutan and Assam felt earthquakes (Mw 6.3
 and 5.1) at the Kopili fault in the northeast Himalaya region, Geomatics, Natural
 Hazards and Risk, 1, 273-281, 2010.
- Kayal, J., Arefiev, S., Barua, S., Hazarika, D., Gogoi, N., Kumar, A., Chowdhury, S., and
 Kalita, S.: Shillong plateau earthquakes in northeast India region: complex tectonic
 model, CURRENT SCIENCE-BANGALORE-, 91, 109, 2006.
- Kayal, J., Arefiev, S., Baruah, S., Hazarika, D., Gogoi, N., Gautam, J., Baruah, S., Dorbath, C.,
 and Tatevossian, R.: Large and great earthquakes in the Shillong plateau–Assam valley
 area of Northeast India Region: Pop-up and transverse tectonics, Tectonophysics, 532,
 186-192, 2012.
- Kayal, J.: Microearthquake seismology and seismotectonics of South Asia, Springer Science
 & Business Media, 2008.
- Khan, M. A. R., and Poskitt, D.: Window length selection and signal-noise separation and
 reconstruction in singular spectrum analysis, Monash Econometrics and Business
 Statistics Working Papers, 23, 2011-2023, 2011.
- Khattri, K., and Tyagi, A.: Seismicity patterns in the Himalayan plate boundary and
 identification of the areas of high seismic potential, Tectonophysics, 96, 281-297, 1983.
- Kumar, A., Walia, V., Arora, B. R., Yang, T. F., Lin, S.-J., Fu, C.-C., Chen, C.-H., and Wen,
 K.-L.: Identifications and removal of diurnal and semidiurnal variations in radon time
 series data of Hsinhua monitoring station in SW Taiwan using singular spectrum
 analysis, Natural Hazards, 79, 317-330, 2015.
- Miklavčić, I., Radolić, V., Vuković, B., Poje, M., Varga, M., Stanić, D., and Planinić, J.: Radon
 anomaly in soil gas as an earthquake precursor, Applied radiation and isotopes, 66,
 1459-1466, 2008.
- Nandy, D., and Dasgupta, S.: Seismotectonic domains of northeastern India and adjacent areas,
 Physics and Chemistry of the Earth, 18, 371-384, 1991.
- Nandy, D.: Geodynamics of Northeastern India and the adjoining region, ACB publications,
 2001.
- Ni, J., and Barazangi, M.: Seismotectonics of the Himalayan collision zone: Geometry of the
 underthrusting Indian plate beneath the Himalaya, Journal of Geophysical Research:
 Solid Earth, 89, 1147-1163, 1984.
- Serita, A., Hattori, K., Yoshino, C., Hayakawa, M., and Isezaki, N.: Principal component
 analysis and singular spectrum analysis of ULF geomagnetic data associated with
 earthquakes, Natural Hazards and Earth System Science, 5, 685-689, 2005.





367 368 369	Sharma, A. K., Walia, V., and Virk, H.: Effect of meteorological parameters on radon emanation at Palampur (HP), Journal of Association of Exploration Geophysicists, 21, 47-50, 2000.
370 371 372	Singh, H., Singh, J., Singh, S., and Bajwa, B.: Radon exhalation rate and uranium estimation study of some soil and rock samples from Tusham ring complex, India using SSNTD technique, Radiation Measurements, 43, S459-S462, 2008.
373 374 375	Singh, S., Kumar, A., Bajwa, B. S., Mahajan, S., Kumar, V., and Dhar, S.: Radon monitoring in soil gas and ground water for earthquake prediction studies in North West Himalayas, India, Terr. Atmos. Ocean. Sci, 21, 685-695, 2010.
376 377	Stranden, E., Kolstad, A., and Lind, B.: The influence of moisture and temperature on radon exhalation, Radiation Protection Dosimetry, 7, 55-58, 1984.
378 379	Stranden, E., Kolstad, A., and Lind, B.: The influence of moisture and temperature on radon exhalation, Radiation Protection Dosimetry, 7, 55-58, 1984.
380	Tomer, A.: Radon as a earthquake precursor: a review, Int J Sci Eng Technol, 4, 815-822, 2016.
381 382	Ulomov, V., and Mavashev, B.: The Tashkent Earthquake of 26 April, Tashkent: Acad. Nauk. Uzbeck, 1971.
383 384	Verma, M., and Bansal, B. K.: Earthquake precursory studies in India: Scenario and future perspectives, Journal of Asian Earth Sciences, 54, 1-8, 2012.
385 386	Virk, H., and Singh, B.: Radon recording of Uttarkashi earthquake, Geophysical Research Letters, 21, 737-740, 1994.
387 388 389	Yang, T., Walia, V., Chyi, L., Fu, C., Chen, CH., Liu, T., Song, S., Lee, C., and Lee, M.: Variations of soil radon and thoron concentrations in a fault zone and prospective earthquakes in SW Taiwan, Radiation Measurements, 40, 496-502, 2005.
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- 402 FIGURES AND TABLES



Figure 1: Map illustrates the earthquake events of $Mw \ge 5$ during 1918 to 2018, in NE-India

408 and its border region $(20^{\circ}-30^{\circ} \text{ N} \text{ and } 86^{\circ}-98^{\circ} \text{ E})$ along with the major tectonic features of the

409 region.







Figure 2: The plot represents the removal of high frequency quasi periodic component for A)
filtered time series of soil radon, B) filtered time series of soil temperature and C) filtered time
series of soil pressure.







423 Figure 3: Covariance matrix of the first 9 group of soil radon (Rn-222) time series.



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426 **Figure 4:** Eigenfunctions of soil radon (Rn-222) first 9 group.







428 Figure 5: Principal Component of soil radon (Rn-222) related to the first 9 eigentriples.











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434 Figure 7: Covariance matrix of the first 9 group of soil temperature time series.

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437 **Figure 8:** Eigenfunctions of soil temperature first 9 group.







439 **Figure 9:** Principal Component of soil temperature related to the first 9 eigentriples.



441 Figure 10: w-Correlation matrix for the 9 reconstructed components of soil temperature time

442 series.







Figure 11: Covariance matrix of the first 9 group of soil pressure time series



Figure 12: Eigenfunctions of soil pressure first 9 group.







449 **Figure 13:** of soil temperature related to the first 9 Eigentriples.













456 Figure 15: Plot showing the reconstructed time series of A) filtered soil radon, B) filtered457 temperature and C) filtered pressure.









462 Figure 16: Residual of reconstructed time series of A) Filtered soil radon, B) temperature and

463 C) pressure respectively.

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Figure 17: Plot showing the reconstructed filtered time series of soil radon emanation along
with earthquake during the investigation period in the vicinity of MPGO, Tezpur (100 km
radially from MPGO) which occurred in a very short span of time.





- 487 Table 1: The correlation co-efficient of soil radon gas concentration with soil pressure and
- temperature at OH-MPGO during year 2017.

Parameters	Average (Avg.)	Standard Deviation (Std.)	% Coefficient /Avg.)	Variation (Std.	Correlation Coefficient
Radon (KBq/m ³)	94.94	23.58	24.84		
Temperature (⁰ C)	28.60	0.62	2.19		0.5
Pressure (mbar)	991.03	2.48	0.25		-0.5





Table 2: Hypocentral parameters of the earthquake events found to have correlation with radon

509 anomaly.

Date of Event	UTC TIME	Lat (°N)	Long (°E)	Place	Depth (km)	Mag (m _b)	Distance from MPGO (km)
09/05/2017	01:53:55	26.3	92.7	Assam	25	3.7	44
09/05/2017	03:26:54	26.6	93.2	Assam	28	3.4	67
20/06/2017	04:31:58	27.1	92.5	West Kameng,	10	3.5	67
				Arunachal Pradesh			
04/07/2017	10:05:47	27.0	92.1	West Kameng, Arunachal	10	3.5	78
				Pradesh			
10/07/2017	23:28:30	27.1	93.8	Papumpare, Arunachal Pradesh	10	3.1	78
25/07/2017	18:28:00	26.3	93.1	Karbi Anglong, Assam	28	3.2	67
05/08/2017	12:24:56	26.8	92.2	Darrang, Assam	10	3.3	44
07/08/2017	11:25:07	26.3	91.7	Kamrup, Assam	10	3.4	100
31/08/2017	17:57:26	26.6	92.7	Sonitpur Assam	10	3.4	67