# Joint Approximate multifractal analysis : further developments correlation and implementation on products of Universal Multifractal fields, with application to rainfall data

Auguste Gires<sup>1</sup>, Ioulia Tchiguirinskaia<sup>1</sup>, and Daniel Schertzer<sup>1</sup> <sup>1</sup>Hydrologie Météorologie et Complexité, Ecole des Ponts ParisTech, Champs-sur-Marne, France **Correspondence:** Auguste Gires (auguste.gires@enpc.fr)

Abstract. Universal Multifractals (UM) have been widely used to simulate and characterize, with the help of only two physically meaningful parameters, geophysical fields extremely variable across <u>a</u> wide range of scales. Such framework relies on the assumption that the underlying field is generated through a multiplicative cascade process. Derived analysis techniques have been extended to study correlations between two fields not only at a single scale and for a single statistical moment as with the

5 covariance, but across scales and for all moments. Such framework of joint multifractal analysis is used here as a starting point to develop and test an approach enabling to analyse and simulate correlation between (approx.) UM fields.

First, the behaviour of two fields consisting of renormalized multiplicative power law combinations of two UM fields is studied. It appears that in the general case the resulting fields can be well approximated by UM fields with known parameters. Limits of this approximation will be quantified and discussed. Techniques to retrieved retrieve the UM parameters of the

10 underlying fields as well as the exponents of the combination have been developed and successfully tested on numerical simulations. In a second step tentative correlation indicators are suggested.

Finally the suggested approach is implemented to study correlation across scales of detailed rainfall data collected with the help of disdrometers of the Fresnel Platform of Ecole des Ponts (see available data at https://hmco.enpc.fr/portfolio-archive/taranis-observatory/). More precisely, four quantities are used : the rain rate (R), the liquid water content (LWC),

15 and the total drop concentration  $(N_t)$  along with the mass weighed diameter  $(D_m)$  which are commonly used to characterize the drop size distribution. Correlations across scales are quantified. Their relative strength (very strong between R and LWC, strong between DSD features and R or LWC, almost null between  $N_t$  and  $D_m$ ) is discussed.

Copyright statement. TEXT

#### 1 Introduction

20 Numerous geophysical fields exhibit intermittent features with sharp fluctuations across all scales, skewed probability distribution and long range correlations. A common framework to analyse and simulate such fields is multifractals. The underlying idea of this framework is that these fields are the result of an underlying multiplicative cascade process. It is physically based in the sense that it is assumed the fields inherit the scale invariant properties of the governing Navier-Stokes equations and hence should exhibit scale invariant features as well. Reader The reader is referred to a review the reviews by Schertzer and

- 25 Lovejoy (2011) and Schertzer and Tchiguirinskaia (2017) for more details. In the specific framework large class of Universal Multifractals (UM) which is a limit behaviour toward which all multifractal processes converge are the stable and attractive limits of non-linearly interacting multifractal processes and correspond to a broad, multiplicative generalization of the central limit theorem (Schertzer and Lovejoy, 1987; Schertzer and Lovejoy, 1997), a conservative field is fully described with the help of only two parameters with a physical interpretation. UM framework was initially developed to address wind fluctuations, and
- 30 has also been implemented on numerous other geophysical fields ranging from rainfall, discharge, temperature or humidity to soil properties and phytoplankton concentration for example.

Much less work has been devoted to the analysis of the correlations / couplings between two fields exhibiting multifractal properties. A framework was originally presented by Meneveau et al. (1990), who suggested to study across scales the properties of joint moments of two multifractal fields, i.e. the product of the two fields raised to two different powers. The behaviour

- of the scaling exponent as a function of the two moments provides information on the correlations between the two fields. They tested their framework on velocity and temperature as well as velocity and vorticity. Such framework has been implemented in many other contexts. Bertol et al. (2017) used it to extract information on the tillage technique by joint analysis of water and soil losses. Siqueira et al. (2018) studied the correlations between soil properties (pH, organic carbon, exchangeable cations and acidity...) and altitude. Wang et al. (2011) focused on joint properties of soil water retention parameters and soil texture;
- 40 while Jiménez-Hornero et al. (2011) focused on the links between wind patterns and surface temperature. Xie et al. (2015) used this framework in a non geophysical domain to better understand the cross correlation between stock market indexes and index of volatilities.

Seuront and Schmitt (2005a, 2005b) suggested a refinement of this framework and introduced a re-normalization of these joint moments of to define an exponent called "generalized correlation function", and used the properties of this function to

- 45 better understand the coupling between fluorescence (which is related to phytoplankton concentration) and temperature for various levels of turbulence. Similar A similar formalism is used by Calif and Schmitt (2014) to study the coupling between wind fluctuations and the aggregate power output from a wind farm. The generalized correlation function is found to be symmetrical with regards to the chosen moments for the two studied fields suggesting a simple relation of proportionality between between the two quantities.
- 50 Actually the previously discussed frameworks have only been implemented for log-normal cascades, for which computations basically boil down to a single parameter and correlation functions are represented by linear ones. Furthermore only two specific cases have been primarily studied, either a proportional or a power law relation between the two studied fields. In this paper, we suggest relying on this theoretical framework and extending its use to Universal Multifractal and to relations between fields consisting of a multiplicative power law combinations.
- 55 In section 2, the theoretical framework of UM and joint multifractal analysis is presented. Its theoretical consequences on the analysis of multiplicative power law combination of UM fields are explored in section 3. Numerical simulations are used to confirm the validity of the suggested analysis techniques. A new indicator of correlation is presented in section 4 and its

limitations discussed. Finally the framework is implemented on rainfall data to study the correlation between rain rate, liquid water content and quantities characterizing the drop size distribution.

#### 60 2 Theoretical framework

65

#### 2.1 Universal Multifractals

The goal is to represent the behaviour of a field  $\epsilon_{\lambda}$  across scales. The resolution  $\lambda$  is defined as the ratio between the outer scale L (i.e. the duration or size of studied event) and the observation scale l ( $\lambda = L/l$ ). In practice, the field at resolution  $\lambda$  is computed by averaging over adjacent times steps or pixels the field measured or simulated at a maximum resolution ( $\lambda_{max}$ ). Multifractal fields exhibit a power law relation between their statistical moment of order q and the resolution  $\lambda$ :

$$\langle \epsilon_{\lambda}^{q} \rangle \approx \lambda^{K(q)} \tag{1}$$

where K(q) is the scaling moment function that fully characterizes the variability across scales of the field. Universal Multifractals (UM) are a specific case towards which multiplicative cascades processes converge (Schertzer and Lovejoy,1987; 1997). Only two parameters with physical interpretation are needed to define K(q) for conservative fields :

- 70  $C_1$ , the mean intermittency co-dimension, which measures the clustering of the (average) intensity at smaller and smaller scales.  $C_1 = 0$  for an a homogeneous field;
  - $\alpha$ , the multifractality index ( $0 \le \alpha \le 2$ ), which measures the clustering variability with regards to the intensity level.

For UM, we have :

$$K(q) = \frac{C_1}{\alpha - 1} (q^\alpha - q) \tag{2}$$

75 K(q) is computed through Trace Moment (TM) analysis which basically consists in plotting Eq. 1 in log-log and estimating the slope of the retrieved straight line. Double Trace Moment (DTM), specifically designed for UM fields, is commonly used to estimate UM parameters (Lavallée et al.,1993). One can also note that UM parameters characterize the first and second derivatives of K(q) near q = 1:

$$\begin{aligned}
K'(1) &= C_1 \\
K''(1) &= C_1 \alpha
\end{aligned}$$
(3)

80 When doing a multifractal analysis, one should keep in mind that such fields can be affected by phase transitions multifractal phase transitions (Schertzer and Lovejoy, 1992). One is associated with sampling limitations. It results from the fact that due to the limited size of studied samples, estimates of statistical moments greater than a given moment  $q_s$  are not be reliable (see Hubert et al., 1993; Douglas and Barros, 2003 for some examples of implementation). In practice, the empirical curve of K(q) will become linear from  $q_s$  and hence depart (being below) from the theoretical curve. The second one is trickier

and associated with the divergence of moments (Schertzer and Lovejoy, 1987). The issue was also mentionned in Mandelbrot (1974) and Kahane (1985) but they did not address the quantification of the spurious statistical estimates on finite samples and their dependence on their size (Schertzer and Lovejoy 1992). It is due to the fact the field generated by a cascade process can become so concentrated that its average over a given area can diverge. This results in  $K(q) \approx +\infty$  for  $q > q_D$ . In practice the K(q) will obviously be computed but its value will be an overestimation of the theoretical K(q) (hence it will be greater).

#### 90 2.2 Joint Multifractal Analysis

Let us consider two fields  $\phi_{\lambda}$  and  $\epsilon_{\lambda}$  exhibiting multifractal properties. In order to study the correlation across scales Seuront and Schmitt (2005a) refined the initial framework of Meneveau et al. (1990) and suggested to perform a joint multifractal analysis as follow follows :

95 where r(h,q)r(q,h) is a "generalized correlation exponent". If  $\phi_{\lambda}$  and  $\epsilon_{\lambda}$  are lognormal multifractal processes (i.e.  $\alpha = 2$ ), then r(h,q)r(q,h) is linear with regards to both *h* and *q*. r(h,q) = 0 r(q,h) = 0 for independent fields. If they are power law related with  $\phi_{\lambda} = c\epsilon_{\lambda}^{d}$ , then r(h,q)r(q,h) is symmetric in the dh - q dp - q plane.

#### 3 Multiplicative combinations of two fields

100

Let us consider two independent UM fields  $X_{\lambda}$  and  $Y_{\lambda}$ , with their respective characteristic parameters  $\alpha_X$ ,  $C_{1,X}$ ,  $\alpha_Y$ ,  $C_{1,Y}$ . The goal of this section is to understand the behaviour of two fields ( $\phi_{\lambda}$  and a field  $\epsilon_{\lambda}$ ) consisting of consisting of a renormalized multiplicative power law combinations of  $X_{\lambda}$  and  $Y_{\lambda}$ . Without any loss of generality, we can assume that  $\phi_{\lambda} = X_{\lambda}$  which is already normalized.  $\epsilon_{\lambda}$  is then defined by :

$$\epsilon_{\lambda} = \frac{X_{\lambda}^{a} Y_{\lambda}^{b}}{\left\langle X_{\lambda}^{a} Y_{\lambda}^{b} \right\rangle} \tag{5}$$

where a and b are exponents characterizing the relative weight of  $X_{\lambda}$  and  $Y_{\lambda}$  in the combination.

#### 105 **3.1** Intuitive understanding of *a* and *b*

Let us first discuss intuitively the influence of the parameters a and b. Fig. 1 displays the fields  $\epsilon_{\lambda}$  (in red) and  $\phi_{\lambda} X_{\lambda}$  (in blue) for a realization of  $X_{\lambda}$  and  $Y_{\lambda}$  with  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$  (Eq. 5 is used). Values of a ranging from 1 to 0 are shown. b was tuned to ensure the same  $C_1$  is retrieved on all the fields. For a = 1 and b = 0 (upper left), the two fields are obviously equal and hence superposed. The opposite case is a = 0 and b = 1 (lower right), for which  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$  are respectively

110 is simply equal to  $Y_{\lambda}$ ,  $X_{\lambda}$ , and hence fully independent with no correlation between themof  $X_{\lambda}$ . In the intermediate cases, the progressive decorrelation between the two fields is visible with decreasing values of *a*. In that sense the parameters *a* and *b* characterize the level of correlation between the two fields.



**Figure 1.**  $\epsilon_{\lambda}$  (in red) and  $\phi_{X} X_{\lambda}$  (in blue) for a realization of  $X_{\lambda}$  and  $Y_{\lambda}$  with  $\alpha_{X} = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_{Y} = 0.8$ ,  $C_{1,Y} = 0.3$ . Definition of Eq. 5 is used. Various values of a = a are shown. b is tuned to ensure the same  $C_{1}$  is retrieved on all the fields.

#### 3.2 Theoretical expectations

In order to evaluate the expected multifractal behaviour of  $\epsilon_{\lambda}$ , its statistical moments of order q are computed to evaluate  $K_{\epsilon}(q)$ . 115 Given that  $X_{\lambda}$  and  $Y_{\lambda}$  are independent, it yields:

$$\langle \epsilon_{\lambda}^{q} \rangle = \lambda^{K_{\epsilon}(q)} = \frac{\langle X_{\lambda}^{qa} \rangle \langle Y_{\lambda}^{qb} \rangle}{\langle X_{\lambda}^{a} \rangle^{q} \langle Y_{\lambda}^{b} \rangle^{q}}$$

$$= \lambda^{K_{X}(qa) - qK_{X}(a) + K_{Y}(qb) - qK_{Y}(b)}$$
(6)

which means we have :

$$K_{\epsilon}(q) = a^{\alpha_{X}} K_{X}(q) + b^{\alpha_{Y}} K_{Y}(q)$$
  
$$= a^{\alpha_{X}} \frac{C_{1,X}}{\alpha_{X}-1} (q^{\alpha_{X}} - q) + b^{\alpha_{Y}} \frac{C_{1,Y}}{\alpha_{Y}-1} (q^{\alpha_{Y}} - q)$$
  
$$\approx \frac{C_{1,\epsilon}}{\alpha_{\epsilon}-1} (q^{\alpha_{\epsilon}} - q)$$
(7)

The exact computation of  $K_{\epsilon}(q)$  is written in the second line of Eq. 7. The third line is not exact and corresponds the form 120  $K_{\epsilon}(q)$  would have if  $\epsilon_{\lambda}$  was actually UM. It is not true in the general case. In order to assess pseudo UM parameters  $C_{1,\epsilon}$  and  $\alpha_{\epsilon}$ , we suggest to use the properties of Eq. 3 and equalize the first and second derivatives of the two last lines of Eq. 7 for q = 1. This yields :

$$C_{1,\epsilon} = C_{1,X}a^{\alpha_X} + C_{1,Y}b^{\alpha_Y}$$

$$\alpha_{\epsilon} = \frac{C_{1,X}a^{\alpha_X}\alpha_X + C_{1,Y}b^{\alpha_Y}\alpha_Y}{C_{1,X}a^{\alpha_X} + C_{1,Y}b^{\alpha_Y}}$$
(8)

It should be noted that in the specific case of  $\alpha_X = \alpha_Y$ , then  $\alpha_\epsilon$  is also equal to this value and  $\epsilon_\lambda$  is actually an exact UM 125 field.

Fig. 2 displays the scaling moment functions of the previously discussed fields for various sets of parameters. Similar results are found for other sets of UM parameters and combinations of *a* and *b* exponents. In Fig. 2.a, the same α is used for both X<sub>λ</sub> and Y<sub>λ</sub>, and the expected exact UM behaviour is correctly retrieved. When α<sub>X</sub> ≠ α<sub>Y</sub>, ε<sub>λ</sub> is not exactly UM. As it is illustrated on Fig. 2.b and c, the smaller the differences, the better is the UM approximation for ε<sub>λ</sub>. In the extreme case when α<sub>Y</sub> = 0 (Fig. 2.c), the approximation remains valid only for *q* ranging from ~ .6 to 1.6. This range is much wider when the αs are closer. It should be noted that for great moments, some discrepancies are visible with the exact value of K<sub>ε</sub>(*q*) always being greater than its UM approximation. This could wrongly be interpreted as a hint suggesting that a multifractal phase transition associated with the divergence of moments is occurring (in the case of Fig. 2.c, we have *q<sub>D</sub>* = 91.1 for the UM parameters of ε<sub>λ</sub>) whereas it is merely an illustration of the limits of validity of the approximation of ε<sub>λ</sub> as a UM field. Indeed, the values of *q<sub>D</sub>* are much

135 greater than the moment for which the discrepancies start to be visible. In the cases of Fig. 2, we have  $q_D = 5.96$  for panel (a),  $q_D = 4.58$  for panel (b) and  $q_D = 119$  for panel (c) for which the approximation as a UM field is the less valid. These values are obtained by looking for the solution > 1 to the equation  $K(q_D) = (q_D - 1)D$  using the pseudo UM parameters of  $\epsilon_{\lambda}$  (D is the dimension of the embedding space, and is equal to 1 for time series). When confronted to such behaviour, keeping in mind this sort of interpretation could be interesting.



**Figure 2.** Illustration of the scaling moment functions K(q) of  $X_{\lambda}$ ,  $Y_{\lambda}$  and  $\epsilon_{\lambda}$ , along with the UM approximation for  $\epsilon_{\lambda}$  (fitted around q = 1). Three possible sets of parameters are displayed

In this sub-section an empirical technique to estimate the UM parameters of  $X_{\lambda}$ ,  $Y_{\lambda}$  and the exponents *a* and *b* from a joint multifractal analysis of  $\phi_{\lambda}$ ,  $X_{\lambda}$  and  $\epsilon_{\lambda}$  is presented. The following steps should be implemented :

(Step 1) Performing a UM analysis of each field φ<sub>X</sub> X<sub>λ</sub> and ε<sub>λ</sub> independently. This enables to confirm the quality of the scaling behaviour and to estimate α<sub>φ</sub> = α<sub>X</sub>, C<sub>1,φ</sub> = C<sub>1,X</sub> α<sub>X</sub>, C<sub>1,X</sub>, α<sub>ε</sub> and C<sub>1,ε</sub>. Without any loss of generality, we can
145 assume that C<sub>1,Y</sub> = C<sub>1,X</sub>. Indeed C<sub>1,Y</sub> is a rather arbitrary quantity that can be changed while the one that actually matters is C<sub>1,Y</sub> b<sup>α<sub>Y</sub></sup>.

(Step 2) Estimating *a*. It is actually the trickiest portion of the process and requires a joint multifractal analysis. More precisely Eq. 4 is implemented with  $\phi_{\lambda} X_{\lambda}$  and  $\epsilon_{\lambda}$ . In that case, it turns out that the ratio does not depend any more on  $Y_{\lambda}$  and only on  $X_{\lambda}$ . One obtains:

150 
$$r(q,h) = K_X(ha+q) - K(ha) - K(q) = \frac{C_{1,X}}{\alpha_X - 1} ((ha+q)^{\alpha_X} - (ha)^{\alpha_X} - (q)^{\alpha_X})$$
(9)

Hence, for a given value of h and q, r(h,q) r(q,h) is an increasing function of a. This property is used to compute an estimate of a. The simplest approach is to set h and q, compute an empirical value of  $r_{emp}(h,q)$   $r_{emp}(q,h)$  and find the a that yields this value. When implementing this technique, one should keep in mind that empirical fields are subject to multifractal phase transitions affecting their scaling behaviour. It means that ha + q, ha and q should remain within the range of values for which the estimations of the scaling moment functions remain reliable, i.e. smaller that the corresponding  $q_s$  and  $q_D$ .

155

(Step 3) Estimating  $\alpha_Y$ . Using Eq. 8, one can easily obtain  $\div$  (noting that  $\alpha_{\epsilon}C_{1,\epsilon} = C_{1,X}a^{\alpha_X}\alpha_X + C_{1,Y}b^{\alpha_Y}\alpha_Y$ , and that the term  $C_{1,X}b^{\alpha_Y}$  is simply equal to  $C_{1,\epsilon} - C_{1,X}a^{\alpha_X}$ , which enables to remove the non linear part of the equation):

$$\alpha_Y = \frac{\frac{C_{1,\epsilon}}{C_{1,\phi}}\alpha_\epsilon - a^{\alpha_\phi}\alpha_\phi}{\frac{C_{1,\epsilon}}{C_{1,\chi}} - a^{\alpha_\phi}} \frac{\frac{C_{1,\epsilon}}{C_{1,\chi}}\alpha_\epsilon - a^{\alpha_X}\alpha_X}{\frac{C_{1,\epsilon}}{C_{1,\chi}} - a^{\alpha_X}}$$
(10)

(Step 4) Computing b. Once  $\alpha_Y$  is known, Eq. 8 (top) can be used to estimate b as  $\div$  (noting that  $C_{1,Y}b^{\alpha_Y} = C_{1,\epsilon} - C_{1,X}a^{\alpha_X}$ 160 and that we have  $C_{1,Y} = C_{1,X}$ ):

$$b = \left(\frac{C_{1,\epsilon}}{C_{1,\phi}}\frac{C_{1,\epsilon}}{C_{1,\chi}} - a^{\alpha_{\phi}\alpha_{\chi}}_{-\infty}\right)^{1/\alpha_{Y}}$$
(11)

#### 3.4 Implementation on numerical simulations (discrete UM)

The approach presented above is tested on numerical simulations - obtained with discrete in scale cascades. It consists in iteratively repeating a cascade step with a non infinetisimal scale ratio in which a 'parent' structure is divided into 'daughter'

165 structures whose affected value is the one of the 'parent' structure multiplied by a random factor ensuring that Eqs. 1 and 2 remain valid. Such simple field generation process is sufficient for the purposes of this paper. The recent introduction of

multifractal operators and vectors paves the way for physically-based, continuous (in scale) multivariate analysis of multifractal fields or measures (Schertzer and Tchiguirinskaia, 2015; Schertzer and Tchiguirinskaia, 2019)

A set of 10 000 realizations of 512 long 1D discrete cascades is used, and analysis are carried out on ensemble average.

170 Before starting, let us clarify the objective of this section.  $X_{\lambda}$  and  $Y_{\lambda}$  are first simulated and then  $\epsilon_{\lambda}$  is build with some values of a and b. The purpose is after to retrieve the values of a, b and  $\alpha_Y$  by simply analysing  $X_{\lambda}$  and  $\epsilon_{\lambda}$  which are assumed to be known.

The parameters used for these simulations are  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$ , a = 0.6 and b = 0.2. As a consequence we expect to find  $\alpha_{\epsilon} = 1.39$ ,  $C_{1,\epsilon} = 0.20$ . Other sets of parameters have been tested and yield similar results.

- 175 Results of this analysis are displayed in Fig. 3. As expected, the scaling behaviour observed on both  $\phi_{\lambda} X_{\lambda}$  and  $\epsilon_{\lambda}$  is excellent. TM-Trace Moment (TM) analysis, i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$  is shown in 3.a and all the coefficients of determination of the straight lines used to compute K(q) are greater than 0.99. With regards to the estimates of UM parameters retrieved via the DTM-Double Trace Moment (DTM) technique, for  $\phi_{\chi} X_{\lambda}$  they are equal to 1.79 and 0.27 for respectively  $\alpha$ and  $C_1$ , which is close to the values inputted input in the simulations. The small discrepancy in  $C_1$  has already been noticed
- with such discrete simulations. The respective estimates for  $\epsilon_{\lambda}$  are 1.35 and 0.18, which are in agreement with the theoretical 180 expectations. These small differences are visible on Fig. 3.b which displays the empirical and theoretical fitting of K(q). For  $\phi_X X_\lambda$ , it can be noted that the empirical estimate of K(q) is smaller that its theoretical value (using UM estimates retrieved from the DTM analysis) for q greater than  $\sim 1.7$ . This is consistent with a behaviour affected by the multifractal phase transition associated with sampling limitation ( $q_s = 1.95$  for the inputted input UM parameters). It can be noted that for  $\epsilon_{\lambda}$  we have a
- 185 greater  $q_s$  equal to 1.95, while it is even greater for  $Y_{\lambda}$  (=4.5). The values of  $q_D$  are greater in all cases, meaning that the multifractal phase transition associated with divergence of moment will not bias our analysis.

In order to estimate a (step 2 of the process described in the previous sub-section), we consider the two moments q = h = 0.7. Note that with these values we have ha + q = 1.12, which is much smaller than the minimum  $q_s$  for the chosen values of UM parameters. It means that the estimates should not be affected by expected biases associated with multifractal phase transitions.

- Fig. 3.c shows the output of joint multifractal (Eq. 4 in log-log plot). It appears that the scaling is excellent and the slope gives 190 an estimate of r(0.7, 0.7). It is then used to estimate a by adjusting the value of a so that r(0.7, 0.7)(a) equals the computed empirical value (3.d). This yields a = 0.59. Finally (Eq. 10 and 11) we obtain an estimate of b equal to 0.20 and an estimate of  $\alpha_Y$  equal to 0.77. These values are very close to the ones inputted input in the simulations. In summary, there is a very good agreement between theoretical expectations and numerical simulations, which confirms the validity of the framework 195 presented in this section.

Finally, let us discuss the uncertainties in the estimates of a. Fig. 4 displays the estimates of a on the simulated fields (see Fig. 3) as a function of the moment orders q and h used in the joint multifractal analysis. It appears that as long as the studied moments remain within the range of reliability of the multifractal analysis (i.e.  $ha + q < q_s$  as previously discussed), the estimates are rather stable. For greater values, there is an underestimation of a.



**Figure 3.** Results of numerical analysis with  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$ , a = 0.6 and b = 0.2 as input parameters. (a) TM analysis i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$ . (b) Scaling moment functions K(q) for  $\epsilon_{\lambda}$  and  $\phi_{\chi} X_{\lambda}$ . (c) Joint multifractal analysis (Eq. 4 in log-log plot) for q = h = 0.7. (d) Illustration of the estimation of a with the values r(0.7, 0.7) computed in (c).

#### 200 4 Toward an indicator of correlation

Let us consider two fields  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ . It is assumed that they both exhibit UM properties, with known UM parameters. The purpose of this section is to present a framework to study the correlations across scales between the two fields. It relies on the joint multifractal analysis presented in section 2.1, with the suggestion of a simplified indicator. It furthermore opens the path to numerical simulations of one field from the other.

205

More precisely, the consequences of describing each field as a multiplicative power law combination of the other and an independent one will be explored. The notations are:

$$\epsilon_{\lambda} = \frac{\phi_{\lambda}^{a} \chi_{\lambda}^{b}}{\left\langle \phi_{\lambda}^{a} \chi_{\lambda}^{b} \right\rangle}$$

$$\phi_{\lambda} = \frac{\epsilon_{\lambda}^{a'} Z_{\lambda}^{b'}}{\left\langle \epsilon_{\lambda}^{a'} Z_{\lambda}^{b'} \right\rangle}$$
(12)



Figure 4. Estimate of a on the simulated fields (see Fig. 3) as a function of the moment orders q and h used in the joint multifractal analysis. The blue grid at the constant value of 0.6 corresponds to the value of a inputted in the simulations

where a, b, a' and b' characterize the level of correlation between the two fields, while  $Y_{\lambda}$  and  $Z_{\lambda}$  are independent random UM fields. As shown in the previous section, without any loss of generality it can be assumed that  $C_{1,Y} = C_{1,\phi}$  and  $C_{1,Z} = C_{1,\epsilon}$ . This enables to simplify the following calculations.

#### 4.1 Limitations of this symmetric framework

If both lines of Eq. 12 were to be correct, then the joint multifractal correlation of  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$  could be computed in two equivalent ways :

215 leading to :

$$\forall h, q \; \frac{C_{1,\phi}}{\alpha_{\phi} - 1} [(\underline{haqa} + \underline{q}\underline{h})^{\alpha_{\phi}} - (\underline{haqa})^{\alpha_{\phi}} - (\underline{q}\underline{h})^{\alpha_{\phi}}] = \frac{C_{1,\epsilon}}{\alpha_{\epsilon} - 1} [(\underline{h}q + a'\underline{q}\underline{h})^{\alpha_{\epsilon}} - (\underline{h}q)^{\alpha_{\epsilon}} - (a'\underline{q}\underline{h})^{\alpha_{\epsilon}}]$$
(14)

In the general case, Eq. 14 is not valid for any q and h. To better understand this, let us consider a given level of correlation by setting the parameters a and b. The goal is to compute a' and b' from the available parameters. The left part of Eq. 14 is known, and after setting given values of q and h it is possible to implement the same process as in section 3.3 to determine a', b' and α<sub>Z</sub>. Fig. 5 displays the outcome of this analysis, according to the values of h and q used, for a = 0.2 in the case α<sub>ε</sub> = 0.8, C<sub>1,ε</sub> = 0.4, α<sub>φ</sub> = 0.8, C<sub>1,φ</sub> = 0.2 (meaning that b = 0.30 and α<sub>Y</sub> = 0.68). As it can be seen, the estimates of a' exhibit a dependency on q and h. The dependency is stronger on h than on q q than on h and estimates remain rather stable as long as h < 0.8q < 0.8. Both sides of Eq. 14 are plotted in Fig. 6 for this set of UM parameters with a = 0.2 and estimates of a' = 0.19 obtained with h = q = 0.7. Expected differences are visible for the larger values of q and h and q. It should be mentioned that these results are presented for a bad case with strong differences between α<sub>ε</sub> and α<sub>φ</sub>. They are actually much smaller if both values are closer to 2. For the specific case, α<sub>ε</sub> = α<sub>φ</sub> = 2, Eq. 14 becomes:

$$\forall h, q \ C_{1,\phi}ahq = C_{1,\epsilon}a'hq \tag{15}$$

meaning that once hq has been removed, a' is deterministically obtained once a is set and  $r_{\epsilon\phi}(q,h) = r_{\epsilon\phi}(q,h) = C_{1,\phi}ahq$  is linear with regards to h and q.



Figure 5. Estimates of a' as a function of h and q using Eq. 14 and the process described in section 3.3. Computation are carried out with  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$  and a = 0.2. The blue horizontal grid corresponds to the value obtained with Eq. 17.

230

Fig. 7 illustrates the relation between the parameters retrieved by setting different values of a in the same case  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$ . First it should be mentioned that for a given set of UM parameters, not all values of a are possible. Indeed the inequality  $0 \le \alpha_Y \le 2$  must be respected leading to  $a \le \min[(\frac{C_{1,\epsilon}\alpha_{\epsilon}}{C_{1,\phi}\alpha_{\phi}})^{1/\alpha_{\phi}}, (\frac{C_{1,\epsilon}(2-\alpha_{\epsilon})}{C_{1,\phi}(2-\alpha_{\phi})})^{1/\alpha_{\phi}}]$ . In this case we must have  $a \le 0.43$ . We retrieved the expected behaviour and are able to quantify it : b decreases with increasing a



Figure 6. Both sides of Eq. 14 for  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$  in the case a = 0.2 and a' = 0.19.  $r_{\epsilon\phi}(q,h)$  is in red and corresponds to the left part of Eq. 14 while  $r_{\phi\epsilon}(q,h)$  is the right part and is in blue. Two views of the same figure are provided to improve visualization.

(Fig. 7.a); a' increases with increasing a (Fig. 7.b),  $\alpha_Y$  decreases with increasing a (Fig. 7.c), and similar behaviour are found in terms of dependency in a' for the symmetric case.

#### 4.2 A simplified indicator

In section 4.1, limitations of this fully symmetric framework are highlighted. However, it is possible to suggest a rather intuitive indicator enabling to extract most of the information obtained from the joint multifractal correlation analysis (i.e. the computation of r(q, h)). It corresponds to the portion of intermittency  $C_1$  of one field explained by the other :

$$240 \qquad IC_{\epsilon\phi} = \frac{C_{1,\phi}a^{\alpha_{\phi}}}{C_{1,\epsilon}} \\ IC_{\phi\epsilon} = \frac{C_{1,\epsilon}a'^{\alpha_{\epsilon}}}{C_{1,\phi}}$$

$$(16)$$

Both "Indicators of Correlation" (IC) are displayed Fig. 8 for the data corresponding to Fig. 7. Both curves are close, and this symmetric behaviour is what is wanted for such an indicator of correlation. Again, much closer curves are obtained with



Figure 7. Illustration of the relations between the various parameters characterizing the correlation across scale between two UM fields in the case  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$ . The dash line in (b) corresponds to the relation obtained by implementing Eq. 17

greater values of  $\alpha$  and identical ones for both  $\alpha$  equal to 2. Forcing  $IC_{\epsilon\phi} = IC_{\phi\epsilon}$  can actually be a way to find an estimate of a' once a is known without having to implement the process described above. It yields :

245 
$$a' = \left(\frac{C_{1,\phi}}{C_{1,\epsilon}}\right)^{2/\alpha_{\epsilon}} a^{\alpha_{\phi}/\alpha_{\epsilon}}$$
(17)

Eq. 17 is actually plotted in dash line on Fig. 7.b, and provides very good estimates. Hence this IC appears as a good candidate for characterizing in a simple way the correlations across scales between two fields. One should keep in mind that it is mainly relevant in the case where the studied fields do not exhibit values of  $\alpha$  too small (typically < 0.8).

#### 5 Implementation on rainfall data

#### 250 5.1 Presentation of the data

The rainfall data used in this paper was collected by a OTT Parsivel<sup>2</sup> disdrometer (Battaglia et al., 2010; OTT, 2014) located on the roof of the Carnot building of the Ecole des Ponts ParisTech campus near Paris between 15 January 2018 and 9 December 2018. It is part of the TARANIS observatory of the Fresnel Platform of École des Ponts ParisTech (https://hmco.enpc.fr/portfolio-archive/fresnel-platform/). Data is collected with 30 seconds time steps. Data will only be briefly presented in this paper and



**Figure 8.** Plot of  $IC_{\epsilon\phi}$  and  $IC_{\phi\epsilon}$  as a function of *a* (Eq. 17) for the same data that is presented in Fig. 7

255 interested readers are referred to Gires et al. (2018b) which discusses available data base in detail along with some data samples for a similar measurement campaign.

In this paper four quantities are studied:

- R, the rain rate in  $mm.h^{-1}$
- LWC, the liquid water content in  $g.m^{-3}$
- 260  $N_t$ , the total drop concentration in  $m^{-3}$

265

-  $D_m$ , the mass weight diameter in mm

 $N_t$  and  $D_m$  are used to characterize the Drop Size Distribution (DSD, N(D), in  $m^{-3}.mm^{-1}$ ) of the rainfall. N(D)dDis the number of drops per unit volume (in  $m^{-3}$ ) with an equivolumic diameter between D and D + dD (in mm). DSD are commonly written in the form  $N(D) = N_t f(D_m)$ , with  $D_m$  being an indicator of the shape of the DSD and  $N_t$  an indicator of the total intensity. They can be computed from the DSD as (Leinonen et al., 2012; Jaffrain and Berne, 2012):

$$N_t = \int_{D_{min}}^{D_{max}} N(D) dD \tag{18}$$

$$D_m = \frac{\int_{D_{min}}^{D_{max}} N(D) D^4 dD}{\int_{D_{min}}^{D_{max}} N(D) D^3 dD}$$
(19)

It should be noted that the disdrometer provides data binned per class of equivolumic diameter and fall velocity, from which

270

a discrete DSD is computed and then used to evaluate the integrals of Eq. 18 and 19 (see Gires et al., 2018b for more details). Multifractal analysis are carried out on ensemble analysis, i.e. on average over various samples. Once rainfall events (an event is defined as a rainy period during which more than 1 mm is collected and that is separated by more than 15 min of dry conditions before and after) have been selected within the longer time series, a process similar as in Gires et al. (2016) and Gires et al. (2018a) is implemented to extract the various samples of data: "for each event (i) a sample size is chosen (a power of two, if possible); (ii) the maximum number of samples for this event is computed; (iii) the portion of the event of length equal to the sample size multiplied by the number of samples found in (ii) with the greatest cumulative depth is extracted; (iv) the extracted series is cut into various samples." Since  $D_m$  is not defined when there is no rain, only samples with no zeros are used.

Dyadic sample size are simpler to use for multifractal analysis, which results in some data not used. With the process

275

described above, 63, 52, 38 and 22 % of the data is actually not used for sample sizes of respectively 32, 64, 128 and 256. 280 A size of 32 time steps, corresponding to 16 min is used, to maximize the amount of data used while keeping an acceptable length for the studied time series. An example of sample for the 4 studied quantities during a rainfall event that occurred on 15 January 2018. 491 such samples are used in the analysis.



Figure 9. Illustration of the four studied rainfall quantifies for 64 long sample corresponding to 32 minutes that occurred on 15 January 2018.

#### 5.2 Joint analysis and discussion

Let us first discuss the results of the joint multifractal analysis carried out between  $N_t$  and R. The purpose is to check if the

285 scale invariant analysis of correlations is relevant for these fields and then to quantify their correlations in this framework (i.e. write the fields as in Eq. 12-top- and estimate a, b and  $\alpha_X$  from the two fields only).

Main curves are shown in Fig. 10 with  $\epsilon_{\lambda}$  being the fluctuations of  $N_t$  and  $\phi_{\lambda}$  being the fluctuations of R. The analysis directly on the field showed that they were non conservative, meaning that the TM and DTM analysis would be biased. Hence multifractal analysis was carried out on an approximation of the underlying conservative fields consisting of their fluctuations

- 290 (Lavallée et al., 1993). Numerical values of the various parameters of the analysis are in Table 1 and 2. R exhibits a very good scaling behaviour on the whole range of scales taken into account as shown by the TM analysis where the coefficients of correlation  $r^2$  of the linear regressions for q around 1 are all greater than 0.98 (Fig. 10.b). Similar scaling behaviour were found on a previous campaign with the same devices (Gires et al., 2016). The scaling for  $N_t$  is worse, with  $r^2$ s only slightly greater than 0.9, but it remains acceptable (Fig. 10.a). We find  $\alpha_R = 1.86$  and  $C_{1,R} = 0.14$  and  $\alpha_{N_t} = 1.78$  and  $C_{1,N_t} = 0.10$ .
- 295 The values of UM parameters observed mean that we are in the domain of highest relevance of the framework developed in the previous section. For R, and to a lesser extent  $N_t$ , there is a clear departure of the fitted K(q) from the empirical one with much greater values for the fitted curve. Furthermore the empirical ones exhibit a linear behaviour from for q approx. greater than 1.5 (Fig. 10.c). Such behaviour is consistent with the expected one when a multifractal phase transition associated with sampling limitations occurs.
- The joint multifractal analysis (Eq. 4 in log-log) for q = h = 0.7 of the two studied fields is displayed in Fig. 10.d. The scaling is good with a value of r<sup>2</sup> = 0.97 for the linear fit. It enables to estimate the exponents a and b at respectively 0.33 and 0.75 (Fig. 10.e). The corresponding *IC* is equal to 0.18. In addition to quantifying the level of correlations between the two fields, it suggests how to simulate one from the other. More precisely, once a time series of fluctuations of *R* is available, it is possible to simulate a realistic corresponding time series of fluctuations of N<sub>t</sub>, by multiplying the *R* series raising to the power a = 0.33 the *R* series and multiplying it with an independent random fields *Y* with α = 1.76 and C<sub>1</sub> = 0.14 raised to the power b = 0.75, and renormalizing the ensemble. Formally it suggests the fluctuations of N<sub>t</sub> can be written as *R*<sup>0.33</sup>/<sub>(Ruetuations Y0.75)</sub>. Such relations opens the path for techniques to simulate fluctuations of N<sub>t</sub> knowing only the temporal evolution of the rain rate.
- Similar qualitative results are found for the other combinations, and numerical values are reported in Tab. 1. Both LWC and  $D_m$  exhibit a good scaling behaviour and their UM parameters are in Tab. 1. As expected given the observed values of  $\alpha$ , the ICs computed in one way or the other (i.e. inverting the role of  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ ) are very similar. Furthermore the values of a' found using Eq. 17 (not shown) are very close to the <u>one ones</u> obtain by inverting the role of the two fields. This confirms the relevancy of the framework of section 4 in this case. It appears that the correlation found between R and LWC is much stronger than between R and N<sub>t</sub> or D<sub>m</sub>. There is no correlation between N<sub>t</sub> or D<sub>m</sub> which is a hint for independence but not
- a proof (if would be for Gaussian variables). Note that the very bad scaling for the joint analysis of these two quantities is partially due to the very small values found for r(q, h) which is basically equal to zero. R exhibits a slightly greater correlation

Field	$\alpha$	$C_1$	$r^2$ for $q = 1.5$
R	1.86	0.14	0.99
LWC	1.82	0.12	0.98
$N_t$	1.78	0.10	0.91
$D_m$	1.87	0.12	0.97

**Table 2.** Numerical output of the joint multifractal analysis of the four studied fields. For each box, using the notations of Eq. 12  $\epsilon_{\lambda}$  corresponds to the field of the column and  $\phi_{\lambda}$  to the linerow.

	R	LWC	$N_t$	$D_m$	
R		0.98	0.97	0.97	$r^2$
		0.82	0.33	0.45	a
		0.38	0.75	0.80	b
		0.78	0.18	0.26	IC
LWC	0.98		0.95	0.97	$r^2$
	0.93		0.44	0.36	a
	0.50		0.75	0.92	b
	0.77		0.27	0.15	IC
N <sub>t</sub>	0.97	0.95		0.50	$r^2$
	0.44	0.53		0.00	a
	1.08	0.94		1.11	b
	0.17	0.27		0.00	IC
$D_m$	0.97	0.97	0.50		$r^2$
	0.51	0.37	0.00		a
	0.91	0.91	0.89		b
	0.25	0.16	0.00		IC

with  $D_m$  (IC = 0.26) than with  $N_t$  (IC = 0.18). It is the inverse for LWC with values of IC respectively equal to 0.15 and 0.27.

#### 6 Conclusions

320 In this paper, we used the framework of joint multifractal analysis to characterize the correlation across scales between two multifractal fields. We extended existing framework to Universal Multifractal and also to analyse the correlations between two fields consisting of renormalized multiplicative power law combinations of two known UM fields. In general, the resulting fields can be well approximated by UM fields. Estimates of the corresponding pseudo UM parameters can be theoretically



**Figure 10.** Results of joint multifractal analysis for  $\epsilon_{\lambda}$  being the fluctuations of  $N_t$  and  $\phi_{\lambda}$  being the fluctuations of R. (a) TM analysis i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$ . (b) Same as in (a) for  $\phi_{\lambda}$  (c) Scaling moment functions K(q) for  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ . (d) Joint multifractal analysis (Eq. 4 in log-log) for q = h = 0.7. (e) Illustration of the estimation of a with the values r(0.7, 0.7) computed in (d).

computed by focusing on the behaviour for moments close to one. These estimates remain valid for a range of moments 325 between  $\sim 0.6$  and  $\sim 1.6$  in the worst case. The closer the two the  $\alpha$  of the initial fields are, the better is the approximation. When both  $\alpha$  are equal, the approximation is exact. An analysis technique to estimate the properties of the underlying fields (UM parameters and power law exponents used in the combination) was developed and validated with the help of numerical simulations.

- In a second step, this analysis was used to develop an innovative framework to investigate the correlations between two 330 UM fields. It basically consists in looking at the best parameters enabling to write one field as a power law multiplicative combination of the other field and a random one. In this context, a good candidate for a simple indicator of the strength of the correlation (called *IC*) is the proportion of intermittency of a field explained by the other one. In the general case, this framework is not symmetric, which is a limitation. However when the  $\alpha$  are typically greater than  $\sim 0.8$ , it is approximately symmetric; meaning that it is relevant to extract some information on the correlations between two fields.
- Finally it was implemented on rainfall data collected by a disdrometer installed on the roof the Ecole des Ponts ParisTech. More precisely the correlations between R and LWC, and DSD features ( $N_t$  and  $D_m$ ) are investigated. First it should be mentioned that the scaling behaviour of both R and LWC is excellent, while the one of the DSD features is only good. The  $\alpha$ are rather similar and greater than 1.7 meaning that it is a favourable context to use the newly developed approach. It appears that the correlation between R and LWC is as expected very strong, the one between R or LWC and the DSD features is
- medium, and the one between  $N_t$  and  $D_m$  is basically null. Besides quantifying these correlations, the developed framework suggests a simple technique to simulate one field from the other. Indeed, it is sufficient to compute a power law multiplicative

combination between one field and a random one to obtain the other. The characteristic parameters of the random field as long well as the power law exponents of the relation can the obtained through a joint multifractal analysis of the two studied fields.

345

Further investigations on other fields in various context should be carried out to confirm the interest of this framework to both characterize and simulate correlations across scales between two multifractal fields. In future work, this framework should also be extended to more than two fields.

Acknowledgements. Authors greatly acknowledge partial financial support from the Chair "Hydrology for Resilient Cities" (endowed by Veolia) of Ecole des Ponts ParisTech, and the Île-de-France region RadX@IdF Project.

#### References

- 350 Battaglia, A., Rustemeier, E., Tokay, A., Blahak, U., and Simmer, C.: PARSIVEL Snow Observations: A Critical Assessment, Journal of Atmospheric and Oceanic Technology, 27, 333–344, http://dx.doi.org/10.1175/2009JTECHA1332.1, 2010.
- Bertol, I., Schick, J., Bandeira, D. H., Paz-Ferreiro, J., and Vázquez, E. V.: Multifractal and joint multifractal analysis of water and soil losses from erosion plots: A case study under subtropical conditions in Santa Catarina highlands, Brazil, Geoderma, 287, 116 125, https://doi.org/https://doi.org/10.1016/j.geoderma.2016.08.008, http://www.sciencedirect.com/science/article/pii/
   S0016706116303470, structure and function of soil and soil cover in a changing worl: characterization and scaling, 2017.
- Calif, R. and Schmitt, F. G.: Multiscaling and joint multiscaling description of the atmospheric wind speed and the aggregate power output from a wind farm, Nonlinear Processes in Geophysics, 21, 379–392, https://doi.org/10.5194/npg-21-379-2014, http://www.nonlin-processes-geophys.net/21/379/2014/, 2014.
- Douglas, E. M. and Barros, A. P.: Probable maximum precipitation estimation using multifractals: Application in the eastern United States,
   Journal of Hydrometeorology, 4, 1012–1024, <GotoISI>://WOS:000187534900003, j. Hydrometeorol., 2003.
  - Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Multifractal comparison of the outputs of two optical disdrometers, Hydrological Sciences Journal, 61, 1641–1651, https://doi.org/10.1080/02626667.2015.1055270, https://doi.org/10.1080/02626667.2015.1055270, 2016.
    - Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Pseudo-radar algorithms with two extremely wet months of disdrometer data in the Paris area, Atmospheric Research, 203, 216 230, https://doi.org/10.1016/j.atmosres.2017.12.011, http://www.sciencedirect. com/science/article/pii/S0169809517307111, 2018a.
    - Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Two months of disdrometer data in the Paris area, Earth System Science Data, 10, 941–950, https://doi.org/10.5194/essd-10-941-2018, https://www.earth-syst-sci-data.net/10/941/2018/, 2018b.
    - Hubert, P., Tessier, Y., Ladoy, P., Lovejoy, S., Schertzer, D., Carbonnel, J. P., Violette, S., Desurosne, I., and Schmitt, F.: Multifractals and extreme rainfall events, Geophys. Res. Lett., 20, 931–934, 1993.
- 370 Jaffrain, J. and Berne, A.: Influence of the Subgrid Variability of the Raindrop Size Distribution on Radar Rainfall Estimators, Journal of Applied Meteorology and Climatology, 51, 780–785, http://dx.doi.org/10.1175/JAMC-D-11-0185.1, 2012.
  - Jiménez-Hornero, F., Pavón-Domínguez, P., de Ravé, E. G., and Ariza-Villaverde, A.: Joint multifractal description of the relationship between wind patterns and land surface air temperature, Atmospheric Research, 99, 366 376, https://doi.org/https://doi.org/10.1016/j.atmosres.2010.11.009, http://www.sciencedirect.com/science/article/pii/S0169809510003121,
- 375 2011.

365

- Kahane, J.: Sur le Chaos Multiplicatif, Ann. Sci. Math. Que., 9, 435–444, 1985.
- Lavallée, D., Lovejoy, S., and Ladoy, P.: Nonlinear variability and landscape topography: analysis and simulation, in: Fractas in geography, edited by de Cola, L. and Lam, N., pp. 171–205, Prentice-Hall, 1993.

Leinonen, J., Moisseev, D., Leskinen, M., and Petersen, W. A.: A Climatology of Disdrometer Measurements of Rainfall in Finland over

- 380 Five Years with Implications for Global Radar Observations, Journal of Applied Meteorology and Climatology, 51, 392–404, http://dx. doi.org/10.1175/JAMC-D-11-056.1, 2012.
  - Mandelbrot, B. B.: Intermittent turbulence in self-similar cascades: divergence of high moments and dimension of the carrier, Journal of Fluid Mechanics, 62, 331–358, https://doi.org/10.1017/S0022112074000711, 1974.

Meneveau, C., Sreenivasan, K. R., Kailasnath, P., and Fan, M. S.: Joint multifractal measures: Theory and applications to turbulence, Phys.

385 Rev. A, 41, 894–913, https://doi.org/10.1103/PhysRevA.41.894, https://link.aps.org/doi/10.1103/PhysRevA.41.894, 1990.

Schertzer, D. and Lovejoy, S.: Physical modelling and analysis of rain and clouds by anisotropic scaling and multiplicative processes, J. Geophys. Res., 92, 9693–9714, 1987.

Schertzer, D. and Lovejoy, S.: Hard and soft multifractal processes, Physica A, 185, 187-194, 1992.

- Schertzer, D. and Lovejoy, S.: Universal multifractals do exist!: Comments, Journal of Applied Meteorology, 36, 1296–1303, <GotoISI>:
   //WOS:A1997XU62400016, j. Appl. Meteorol., 1997.
  - Schertzer, D. and Lovejoy, S.: Multifractals, generalized scale invariance and complexity in geophysics, International Journal of Bifurcation and Chaos, 21, 3417–3456, <GotoISI>://WOS:000300016000003, 2011.
    - Schertzer, D. and Tchiguirinskaia, I.: Multifractal vector fields and stochastic Clifford algebra, Chaos, 25, 123127, https://doi.org/DOI:http://dx.doi.org/10.1063/1.4937364, 2015.
- 395 Schertzer, D. and Tchiguirinskaia, I.: An Introduction to Multifractals and Scale Symmetry Groups, in: Fractals: Concepts and Applications in Geosciences., CRC Press, edited by Ghanbarian, B. and Hunt, A., pp. 1–28, 2017.
  - Schertzer, D. and Tchiguirinskaia, I.: A century of turbulent cascades and the emergence of multifractal operators (invited paper under review), Earth and Space Science, 2019.
- Seuront, L. and Schmitt, F. G.: Multiscaling statistical procedures for the exploration of biophysical couplings in inter mittent turbulence. Part I. Theory, Deep Sea Research Part II: Topical Studies in Oceanography, 52, 1308 1324, https://doi.org/https://doi.org/10.1016/j.dsr2.2005.01.006, http://www.sciencedirect.com/science/article/pii/S0967064505000470, observations and modelling of mixed layer turbulence: Do they represent the same statistical quantities?, 2005a.
- Seuront, L. and Schmitt, F. G.: Multiscaling statistical procedures for the exploration of biophysical couplings in intermittent turbulence. Part II. Applications, Deep Sea Research Part II: Topical Studies in Oceanography, 52, 1325 1343, https://doi.org/https://doi.org/10.1016/j.dsr2.2005.01.005, http://www.sciencedirect.com/science/article/pii/S0967064505000482, obser-
- vations and modelling of mixed layer turbulence: Do they represent the same statistical quantities?, 2005b.
  - Siqueira, G. M., Ênio F.F. Silva, Vidal-Vázquez, E., and Paz-González, A.: Multifractal and joint multifractal analysis of general soil properties and altitude along a transect, Biosystems Engineering, 168, 105 – 120, https://doi.org/https://doi.org/10.1016/j.biosystemseng.2017.08.024, http://www.sciencedirect.com/science/article/pii/
- 410 S1537511016305943, computational Tools to Support Soil Management Decisions, 2018.
  - Wang, Z.-Y., Shu, Q.-S., Xie, L.-Y., Liu, Z.-X., and Si, B.: Joint Multifractal Analysis of Scaling Relationships Between Soil Water-Retention Parameters and Soil Texture, Pedosphere, 21, 373 – 379, https://doi.org/https://doi.org/10.1016/S1002-0160(11)60138-0, http://www.sciencedirect.com/science/article/pii/S1002016011601380, 2011.

Xie, W.-J., Jiang, Z.-Q., Gu, G.-F., Xiong, X., and Zhou, W.-X.: Joint multifractal analysis based on the partition function approach: analytical

415 analysis, numerical simulation and empirical application, New Journal of Physics, 17, 103 020, http://stacks.iop.org/1367-2630/17/i=10/ a=103020, 2015.

21

## Approximate multifractal correlation and products of Universal Multifractal fields, with application to rainfall data

Auguste Gires<sup>1</sup>, Ioulia Tchiguirinskaia<sup>1</sup>, and Daniel Schertzer<sup>1</sup> <sup>1</sup>Hydrologie Météorologie et Complexité, Ecole des Ponts ParisTech, Champs-sur-Marne, France **Correspondence:** Auguste Gires (auguste.gires@enpc.fr)

Abstract. Universal Multifractals (UM) have been widely used to simulate and characterize, with the help of only two physically meaningful parameters, geophysical fields extremely variable across a wide range of scales. Such framework relies on the assumption that the underlying field is generated through a multiplicative cascade process. Derived analysis techniques have been extended to study correlations between two fields not only at a single scale and for a single statistical moment as with the covariance, but across scales and for all moments. Such framework of joint multifractal analysis is used here as a starting point

to develop and test an approach enabling to analyse and simulate correlation between (approx.) UM fields.

First, the behaviour of two fields consisting of renormalized multiplicative power law combinations of two UM fields is studied. It appears that in the general case the resulting fields can be well approximated by UM fields with known parameters. Limits of this approximation will be quantified and discussed. Techniques to retrieve the UM parameters of the underlying

10 fields as well as the exponents of the combination have been developed and successfully tested on numerical simulations. In a second step tentative correlation indicators are suggested.

Finally the suggested approach is implemented to study correlation across scales of detailed rainfall data collected with the help of disdrometers of the Fresnel Platform of Ecole des Ponts (see available data at https://hmco.enpc.fr/portfolio-archive/taranis-observatory/). More precisely, four quantities are used : the rain rate (R), the liquid water content (LWC),

and the total drop concentration  $(N_t)$  along with the mass weighed diameter  $(D_m)$  which are commonly used to characterize the drop size distribution. Correlations across scales are quantified. Their relative strength (very strong between R and LWC, strong between DSD features and R or LWC, almost null between  $N_t$  and  $D_m$ ) is discussed.

Copyright statement. TEXT

#### 1 Introduction

5

20 Numerous geophysical fields exhibit intermittent features with sharp fluctuations across all scales, skewed probability distribution and long range correlations. A common framework to analyse and simulate such fields is multifractals. The underlying idea of this framework is that these fields are the result of an underlying multiplicative cascade process. It is physically based in the sense that it is assumed the fields inherit the scale invariant properties of the governing Navier-Stokes equations and hence should exhibit scale invariant features as well. The reader is referred to a the reviews by Schertzer and Lovejoy (2011) and

25 Schertzer and Tchiguirinskaia (2017) for more details. In the large class of Universal Multifractals (UM) which are the stable and attractive limits of non-linearly interacting multifractal processes and correspond to a broad, multiplicative generalization of the central limit theorem (Schertzer and Lovejoy, 1987; Schertzer and Lovejoy, 1997), a conservative field is fully described with the help of only two parameters with a physical interpretation. UM framework was initially developed to address wind fluctuations, and has also been implemented on numerous other geophysical fields ranging from rainfall, discharge, temperature or humidity to soil properties and phytoplankton concentration for example.

Much less work has been devoted to the analysis of the correlations / couplings between two fields exhibiting multifractal properties. A framework was originally presented by Meneveau et al. (1990), who suggested to study across scales the properties of joint moments of two multifractal fields, i.e. the product of the two fields raised to two different powers. The behaviour of the scaling exponent as a function of the two moments provides information on the correlations between the two fields. They

- 35 tested their framework on velocity and temperature as well as velocity and vorticity. Such framework has been implemented in many other contexts. Bertol et al. (2017) used it to extract information on the tillage technique by joint analysis of water and soil losses. Siqueira et al. (2018) studied the correlations between soil properties (pH, organic carbon, exchangeable cations and acidity...) and altitude. Wang et al. (2011) focused on joint properties of soil water retention parameters and soil texture; while Jiménez-Hornero et al. (2011) focused on the links between wind patterns and surface temperature. Xie et al. (2015)
- 40 used this framework in a non geophysical domain to better understand the cross correlation between stock market indexes and index of volatilities.

Seuront and Schmitt (2005a, 2005b) suggested a refinement of this framework and introduced a re-normalization of these joint moments to define an exponent called "generalized correlation function", and used the properties of this function to better understand the coupling between fluorescence (which is related to phytoplankton concentration) and temperature for various

45 levels of turbulence. A similar formalism is used by Calif and Schmitt (2014) to study the coupling between wind fluctuations and the aggregate power output from a wind farm. The generalized correlation function is found to be symmetrical with regards to the chosen moments for the two studied fields suggesting a simple relation of proportionality between the two quantities.

Actually the previously discussed frameworks have only been implemented for log-normal cascades, for which computations basically boil down to a single parameter and correlation functions are represented by linear ones. Furthermore only two specific cases have been primarily studied, either a proportional or a power law relation between the two studied fields. In this paper,

50 cases have been primarily studied, either a proportional or a power law relation between the two studied fields. In this paper, we suggest relying on this theoretical framework and extending its use to Universal Multifractal and to relations between fields consisting of a multiplicative power law combinations.

In section 2, the theoretical framework of UM and joint multifractal analysis is presented. Its theoretical consequences on the analysis of multiplicative power law combination of UM fields are explored in section 3. Numerical simulations are used

55 to confirm the validity of the suggested analysis techniques. A new indicator of correlation is presented in section 4 and its limitations discussed. Finally the framework is implemented on rainfall data to study the correlation between rain rate, liquid water content and quantities characterizing the drop size distribution.

### 2 Theoretical framework

#### 2.1 Universal Multifractals

60 The goal is to represent the behaviour of a field ε<sub>λ</sub> across scales. The resolution λ is defined as the ratio between the outer scale L (i.e. the duration or size of studied event) and the observation scale l (λ = L/l). In practice, the field at resolution λ is computed by averaging over adjacent times steps or pixels the field measured or simulated at a maximum resolution (λ<sub>max</sub>). Multifractal fields exhibit a power law relation between their statistical moment of order q and the resolution λ:

$$\langle \epsilon^q_\lambda \rangle \approx \lambda^{K(q)} \tag{1}$$

- 65 where K(q) is the scaling moment function that fully characterizes the variability across scales of the field. Universal Multifractals (UM) are a specific case towards which multiplicative cascades processes converge (Schertzer and Lovejoy,1987; 1997). Only two parameters with physical interpretation are needed to define K(q) for conservative fields :
  - $C_1$ , the mean intermittency co-dimension, which measures the clustering of the (average) intensity at smaller and smaller scales.  $C_1 = 0$  for a homogeneous field;
  - 70  $\alpha$ , the multifractality index ( $0 \le \alpha \le 2$ ), which measures the clustering variability with regards to the intensity level.

For UM, we have :

$$K(q) = \frac{C_1}{\alpha - 1}(q^\alpha - q) \tag{2}$$

K(q) is computed through Trace Moment (TM) analysis which basically consists in plotting Eq. 1 in log-log and estimating the slope of the retrieved straight line. Double Trace Moment (DTM), specifically designed for UM fields, is commonly used
to estimate UM parameters (Lavallée et al.,1993). One can also note that UM parameters characterize the first and second derivatives of K(q) near q = 1:

$$\begin{array}{rcl}
K'(1) &=& C_1 \\
K''(1) &=& C_1 \alpha
\end{array}$$
(3)

80

When doing a multifractal analysis, one should keep in mind that such fields can be affected by multifractal phase transitions (Schertzer and Lovejoy, 1992). One is associated with sampling limitations. It results from the fact that due to the limited size of studied samples, estimates of statistical moments greater than a given moment  $q_s$  are not be reliable (see Hubert et al., 1993; Douglas and Barros, 2003 for some examples of implementation). In practice, the empirical curve of K(q) will become linear from  $q_s$  and hence depart (being below) from the theoretical curve. The second one is trickier and associated with the divergence of moments (Schertzer and Lovejoy, 1987). The issue was also mentionned in Mandelbrot (1974) and Kahane (1985) but they did not address the quantification of the spurious statistical estimates on finite samples and their dependence on their size (Schertzer and Lovejoy 1992). It is due to the fact the field generated by a cascade process can become so concentrated that its average over a given area can diverge. This results in  $K(q) \approx +\infty$  for  $q > q_D$ . In practice the K(q) will obviously be computed but its value will be an overestimation of the theoretical K(q) (hence it will be greater).

#### 2.2 Joint Multifractal Analysis

Let us consider two fields φ<sub>λ</sub> and ε<sub>λ</sub> exhibiting multifractal properties. In order to study the correlation across scales Seuront
and Schmitt (2005a) refined the initial framework of Meneveau et al. (1990) and suggested to perform a joint multifractal analysis as follows :

$$\frac{\langle \epsilon_{\lambda}^{q} \phi_{\lambda}^{h} \rangle}{\langle \epsilon_{\lambda}^{q} \rangle \langle \phi_{\lambda}^{h} \rangle} \approx \lambda^{S(q,h)-K_{\epsilon}(q)-K_{\phi}(h)} \approx \lambda^{r(q,h)}$$

$$\tag{4}$$

where r(q,h) is a "generalized correlation exponent". If  $\phi_{\lambda}$  and  $\epsilon_{\lambda}$  are lognormal multifractal processes (i.e.  $\alpha = 2$ ), then r(q,h) is linear with regards to both h and q. r(q,h) = 0 for independent fields. If they are power law related with  $\phi_{\lambda} = c\epsilon_{\lambda}^{d}$ , 95 then r(q,h) is symmetric in the dp - q plane.

#### 3 Multiplicative combinations of two fields

Let us consider two independent UM fields  $X_{\lambda}$  and  $Y_{\lambda}$ , with their respective characteristic parameters  $\alpha_X$ ,  $C_{1,X}$ ,  $\alpha_Y$ ,  $C_{1,Y}$ . The goal of this section is to understand the behaviour of a field  $\epsilon_{\lambda}$  consisting of a renormalized multiplicative power law combinations of  $X_{\lambda}$  and  $Y_{\lambda}$ .  $\epsilon_{\lambda}$  is then defined by :

100 
$$\epsilon_{\lambda} = \frac{X_{\lambda}^{a} Y_{\lambda}^{b}}{\left\langle X_{\lambda}^{a} Y_{\lambda}^{b} \right\rangle} \tag{5}$$

where a and b are exponents characterizing the relative weight of  $X_{\lambda}$  and  $Y_{\lambda}$  in the combination.

#### 3.1 Intuitive understanding of a and b

Let us first discuss intuitively the influence of the parameters a and b. Fig. 1 displays the fields ε<sub>λ</sub> (in red) and X<sub>λ</sub> (in blue) for a realization of X<sub>λ</sub> and Y<sub>λ</sub> with α<sub>X</sub> = 1.8, C<sub>1,X</sub> = 0.3, α<sub>Y</sub> = 0.8, C<sub>1,Y</sub> = 0.3 (Eq. 5 is used). Values of a ranging from 1 to 0
are shown. b was tuned to ensure the same C<sub>1</sub> is retrieved on all the fields. For a = 1 and b = 0 (upper left), the two fields are obviously equal and hence superposed. The opposite case is a = 0 and b = 1 (lower right), for which ε<sub>λ</sub> is simply equal to Y<sub>λ</sub>, and hence fully independent of X<sub>λ</sub>. In the intermediate cases, the progressive decorrelation between the two fields is visible with decreasing values of a. In that sense the parameters a and b characterize the level of correlation between the two fields.



Figure 1.  $\epsilon_{\lambda}$  (in red) and  $X_{\lambda}$  (in blue) for a realization of  $X_{\lambda}$  and  $Y_{\lambda}$  with  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$ . Definition of Eq. 5 is used. Various values of *a* are shown. *b* is tuned to ensure the same  $C_1$  is retrieved on all the fields.

#### 3.2 Theoretical expectations

110 In order to evaluate the expected multifractal behaviour of  $\epsilon_{\lambda}$ , its statistical moments of order q are computed to evaluate  $K_{\epsilon}(q)$ . Given that  $X_{\lambda}$  and  $Y_{\lambda}$  are independent, it yields:

$$\langle \epsilon_{\lambda}^{q} \rangle = \lambda^{K_{\epsilon}(q)} = \frac{\langle X_{\lambda}^{qa} \rangle \langle Y_{\lambda}^{qb} \rangle}{\langle X_{\lambda}^{a} \rangle^{q} \langle Y_{\lambda}^{b} \rangle^{q}}$$

$$= \lambda^{K_{X}(qa) - qK_{X}(a) + K_{Y}(qb) - qK_{Y}(b)}$$

$$(6)$$

which means we have :

$$K_{\epsilon}(q) = a^{\alpha_{X}} K_{X}(q) + b^{\alpha_{Y}} K_{Y}(q)$$
  
$$= a^{\alpha_{X}} \frac{C_{1,X}}{\alpha_{X}-1} (q^{\alpha_{X}} - q) + b^{\alpha_{Y}} \frac{C_{1,Y}}{\alpha_{Y}-1} (q^{\alpha_{Y}} - q)$$
  
$$\approx \frac{C_{1,\epsilon}}{\alpha_{\epsilon}-1} (q^{\alpha_{\epsilon}} - q)$$
(7)

115

The exact computation of  $K_{\epsilon}(q)$  is written in the second line of Eq. 7. The third line is not exact and corresponds the form  $K_{\epsilon}(q)$  would have if  $\epsilon_{\lambda}$  was actually UM. It is not true in the general case. In order to assess pseudo UM parameters  $C_{1,\epsilon}$  and  $\alpha_{\epsilon}$ , we suggest to use the properties of Eq. 3 and equalize the first and second derivatives of the two last lines of Eq. 7 for q = 1. This yields :

$$C_{1,\epsilon} = C_{1,X}a^{\alpha_X} + C_{1,Y}b^{\alpha_Y}$$

$$\alpha_{\epsilon} = \frac{C_{1,X}a^{\alpha_X}\alpha_X + C_{1,Y}b^{\alpha_Y}\alpha_Y}{C_{1,X}a^{\alpha_X} + C_{1,Y}b^{\alpha_Y}}$$
(8)

120 It should be noted that in the specific case of  $\alpha_X = \alpha_Y$ , then  $\alpha_\epsilon$  is also equal to this value and  $\epsilon_\lambda$  is actually an exact UM field.

Fig. 2 displays the scaling moment functions of the previously discussed fields for various sets of parameters. Similar results are found for other sets of UM parameters and combinations of a and b exponents. In Fig. 2.a, the same  $\alpha$  is used for both  $X_{\lambda}$ 

125

and  $Y_{\lambda}$ , and the expected exact UM behaviour is correctly retrieved. When  $\alpha_X \neq \alpha_Y$ ,  $\epsilon_{\lambda}$  is not exactly UM. As it is illustrated on Fig. 2.b and c, the smaller the differences, the better is the UM approximation for  $\epsilon_{\lambda}$ . In the extreme case when  $\alpha_Y = 0$ (Fig. 2.c), the approximation remains valid only for q ranging from  $\sim$  .6 to 1.6. This range is much wider when the  $\alpha$ s are closer. It should be noted that for great moments, some discrepancies are visible with the exact value of  $K_{\epsilon}(q)$  always being greater than its UM approximation. This could wrongly be interpreted as a hint suggesting that a multifractal phase transition associated with the divergence of moments is occurring whereas it is merely an illustration of the limits of validity of the 130 approximation of  $\epsilon_{\lambda}$  as a UM field. Indeed, the values of  $q_D$  are much greater than the moment for which the discrepancies start to be visible. In the cases of Fig. 2, we have  $q_D = 5.96$  for panel (a),  $q_D = 4.58$  for panel (b) and  $q_D = 119$  for panel (c) for which the approximation as a UM field is the less valid. These values are obtained by looking for the solution > 1 to the equation  $K(q_D) = (q_D - 1)D$  using the pseudo UM parameters of  $\epsilon_{\lambda}$  (D is the dimension of the embedding space, and is equal to 1 for time series). When confronted to such behaviour, keeping in mind this sort of interpretation could be interesting.



Figure 2. Illustration of the scaling moment functions K(q) of  $X_{\lambda}$ ,  $Y_{\lambda}$  and  $\epsilon_{\lambda}$ , along with the UM approximation for  $\epsilon_{\lambda}$  (fitted around q = 1). Three possible sets of parameters are displayed

#### 3.3 **Techniques for retrieving parameters** 135

In this sub-section an empirical technique to estimate the UM parameters of  $Y_{\lambda}$  and the exponents a and b from a joint multifractal analysis of  $X_{\lambda}$  and  $\epsilon_{\lambda}$  is presented. The following steps should be implemented :

(Step 1) Performing a UM analysis of each field  $X_{\lambda}$  and  $\epsilon_{\lambda}$  independently. This enables to confirm the quality of the scaling behaviour and to estimate  $\alpha_X$ ,  $C_{1,X}$ ,  $\alpha_{\epsilon}$  and  $C_{1,\epsilon}$ . Without any loss of generality, we can assume that  $C_{1,Y} = C_{1,X}$ . Indeed  $C_{1,Y}$  is a rather arbitrary quantity that can be changed while the one that actually matters is  $C_{1,Y}b^{\alpha_Y}$ . 140

(Step 2) Estimating *a*. It is actually the trickiest portion of the process and requires a joint multifractal analysis. More precisely Eq. 4 is implemented with  $X_{\lambda}$  and  $\epsilon_{\lambda}$ . In that case, it turns out that the ratio does not depend any more on  $Y_{\lambda}$  and only on  $X_{\lambda}$ . One obtains:

$$r(q,h) = K_X(ha+q) - K(ha) - K(q) = \frac{C_{1,X}}{\alpha_X - 1}((ha+q)^{\alpha_X} - (ha)^{\alpha_X} - (q)^{\alpha_X})$$
(9)

Hence, for a given value of h and q, r(q,h) is an increasing function of a. This property is used to compute an estimate of a. The simplest approach is to set h and q, compute an empirical value of  $r_{emp}(q,h)$  and find the a that yields this value. When implementing this technique, one should keep in mind that empirical fields are subject to multifractal phase transitions affecting their scaling behaviour. It means that ha + q, ha and q should remain within the range of values for which the estimations of the scaling moment functions remain reliable, i.e. smaller that the corresponding  $q_s$  and  $q_D$ .

150

(Step 3) Estimating  $\alpha_Y$ . Using Eq. 8, one can easily obtain (noting that  $\alpha_{\epsilon}C_{1,\epsilon} = C_{1,X}a^{\alpha_X}\alpha_X + C_{1,Y}b^{\alpha_Y}\alpha_Y$ , and that the term  $C_{1,Y}b^{\alpha_Y}$  is simply equal to  $C_{1,\epsilon} - C_{1,X}a^{\alpha_X}$ , which enables to remove the non linear part of the equation) :

$$\alpha_Y = \frac{\frac{C_{1,\epsilon}}{C_{1,X}} \alpha_\epsilon - a^{\alpha_X} \alpha_X}{\frac{C_{1,\epsilon}}{C_{1,X}} - a^{\alpha_X}} \tag{10}$$

(Step 4) Computing b. Once  $\alpha_Y$  is known, Eq. 8 (top) can be used to estimate b as (noting that  $C_{1,Y}b^{\alpha_Y} = C_{1,\epsilon} - C_{1,X}a^{\alpha_X}$ and that we have  $C_{1,Y} = C_{1,X}$ ):

155 
$$b = \left(\frac{C_{1,\epsilon}}{C_{1,X}} - a^{\alpha_X}\right)^{1/\alpha_Y}$$
 (11)

#### 3.4 Implementation on numerical simulations (discrete UM)

The approach presented above is tested on numerical simulations obtained with discrete in scale cascades. It consists in iteratively repeating a cascade step with a non infinetisimal scale ratio in which a 'parent' structure is divided into 'daughter' structures whose affected value is the one of the 'parent' structure multiplied by a random factor ensuring that Eqs. 1 and 2 remain valid. Such simple field generation process is sufficient for the purposes of this paper. The recent introduction of multifractal operators and vectors paves the way for physically-based, continuous (in scale) multivariate analysis of multifractal fields or measures (Schertzer and Tchiguirinskaia, 2015; Schertzer and Tchiguirinskaia, 2019)

A set of 10 000 realizations of 512 long 1D discrete cascades is used, and analysis are carried out on ensemble average.

Before starting, let us clarify the objective of this section.  $X_{\lambda}$  and  $Y_{\lambda}$  are first simulated and then  $\epsilon_{\lambda}$  is build with some 165 values of *a* and *b*. The purpose is after to retrieve the values of *a*, *b* and  $\alpha_Y$  by simply analysing  $X_{\lambda}$  and  $\epsilon_{\lambda}$  which are assumed to be known.

The parameters used for these simulations are  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$ , a = 0.6 and b = 0.2. As a consequence we expect to find  $\alpha_{\epsilon} = 1.39$ ,  $C_{1,\epsilon} = 0.20$ . Other sets of parameters have been tested and yield similar results.

Results of this analysis are displayed in Fig. 3. As expected, the scaling behaviour observed on both  $X_{\lambda}$  and  $\epsilon_{\lambda}$  is excellent.

- 170 Trace Moment (TM) analysis, i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$  is shown in 3.a and all the coefficients of determination of the straight lines used to compute K(q) are greater than 0.99. With regards to the estimates of UM parameters retrieved via the Double Trace Moment (DTM) technique, for  $X_{\lambda}$  they are equal to 1.79 and 0.27 for respectively  $\alpha$  and  $C_1$ , which is close to the values input in the simulations. The small discrepancy in  $C_1$  has already been noticed with such discrete simulations. The respective estimates for  $\epsilon_{\lambda}$  are 1.35 and 0.18, which are in agreement with the theoretical expectations. These small differences
- are visible on Fig. 3.b which displays the empirical and theoretical fitting of K(q). For X<sub>λ</sub>, it can be noted that the empirical estimate of K(q) is smaller that its theoretical value (using UM estimates retrieved from the DTM analysis) for q greater than ~ 1.7. This is consistent with a behaviour affected by the multifractal phase transition associated with sampling limitation (q<sub>s</sub> = 1.95 for the input UM parameters). It can be noted that for ε<sub>λ</sub> we have a greater q<sub>s</sub> equal to 1.95, while it is even greater for Y<sub>λ</sub> (=4.5). The values of q<sub>D</sub> are greater in all cases, meaning that the multifractal phase transition associated with divergence
  of moment will not bias our analysis.
  - In order to estimate a (step 2 of the process described in the previous sub-section), we consider the two moments q = h = 0.7. Note that with these values we have ha + q = 1.12, which is much smaller than the minimum  $q_s$  for the chosen values of UM parameters. It means that the estimates should not be affected by expected biases associated with multifractal phase transitions.
- an estimate of r(0.7, 0.7). It is then used to estimate *a* by adjusting the value of *a* so that r(0.7, 0.7)(a) equals the computed empirical value (3.d). This yields a = 0.59. Finally (Eq. 10 and 11) we obtain an estimate of *b* equal to 0.20 and an estimate of  $\alpha_Y$  equal to 0.77. These values are very close to the ones input in the simulations. In summary, there is a very good agreement between theoretical expectations and numerical simulations, which confirms the validity of the framework presented in this section.

Fig. 3.c shows the output of joint multifractal (Eq. 4 in log-log plot). It appears that the scaling is excellent and the slope gives

Finally, let us discuss the uncertainties in the estimates of a. Fig. 4 displays the estimates of a on the simulated fields (see Fig. 3) as a function of the moment orders q and h used in the joint multifractal analysis. It appears that as long as the studied moments remain within the range of reliability of the multifractal analysis (i.e.  $ha + q < q_s$  as previously discussed), the estimates are rather stable. For greater values, there is an underestimation of a.

#### 4 Toward an indicator of correlation

195 Let us consider two fields  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ . It is assumed that they both exhibit UM properties, with known UM parameters. The purpose of this section is to present a framework to study the correlations across scales between the two fields. It relies on the joint multifractal analysis presented in section 2.1, with the suggestion of a simplified indicator. It furthermore opens the path to numerical simulations of one field from the other.



**Figure 3.** Results of numerical analysis with  $\alpha_X = 1.8$ ,  $C_{1,X} = 0.3$ ,  $\alpha_Y = 0.8$ ,  $C_{1,Y} = 0.3$ , a = 0.6 and b = 0.2 as input parameters. (a) TM analysis i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$ . (b) Scaling moment functions K(q) for  $\epsilon_{\lambda}$  and  $X_{\lambda}$ . (c) Joint multifractal analysis (Eq. 4 in log-log plot) for q = h = 0.7. (d) Illustration of the estimation of a with the values r(0.7, 0.7) computed in (c).

More precisely, the consequences of describing each field as a multiplicative power law combination of the other and an 200 independent one will be explored. The notations are:

$$\epsilon_{\lambda} = \frac{\phi_{\lambda}^{a} Y_{\lambda}^{b}}{\langle \phi_{\lambda}^{a} Y_{\lambda}^{b} \rangle}$$
  

$$\phi_{\lambda} = \frac{\epsilon_{\lambda}^{a'} Z_{\lambda}^{b'}}{\langle \epsilon_{\lambda}^{a'} Z_{\lambda}^{b'} \rangle}$$
(12)

where a, b, a' and b' characterize the level of correlation between the two fields, while  $Y_{\lambda}$  and  $Z_{\lambda}$  are independent random UM fields. As shown in the previous section, without any loss of generality it can be assumed that  $C_{1,Y} = C_{1,\phi}$  and  $C_{1,Z} = C_{1,\epsilon}$ . This enables to simplify the following calculations.



Figure 4. Estimate of a on the simulated fields (see Fig. 3) as a function of the moment orders q and h used in the joint multifractal analysis. The blue grid at the constant value of 0.6 corresponds to the value of a inputted in the simulations

#### 205 4.1 Limitations of this symmetric framework

If both lines of Eq. 12 were to be correct, then the joint multifractal correlation of  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$  could be computed in two equivalent ways :

$$\frac{\langle \epsilon_{\lambda}^{a} \phi_{\lambda}^{h} \rangle}{\langle \epsilon_{\lambda}^{a} \rangle \langle \phi_{\lambda}^{h} \rangle} = \frac{\langle \phi_{\lambda}^{aq+h} Y_{\lambda}^{bq} \rangle}{\langle \phi_{\lambda}^{ah} \rangle \langle Y_{\lambda}^{bq} \rangle \langle \phi_{\lambda}^{d} \rangle} = \frac{\langle \phi_{\lambda}^{aq+h} \rangle}{\langle \phi_{\lambda}^{aq} \rangle \langle \phi_{\lambda}^{h} \rangle} 
= \lambda^{r_{\epsilon\phi}(q,h)} = \lambda^{\frac{c_{1,\phi}}{c_{\phi}-1}[(aq+h)^{\alpha_{\phi}} - (aq)^{\alpha_{\phi}} - (h)^{\alpha_{\phi}}]} 
\frac{\langle \epsilon_{\lambda}^{q} \phi_{\lambda}^{h} \rangle}{\langle \epsilon_{\lambda}^{q} \rangle \langle \phi_{\lambda}^{h} \rangle} = \frac{\langle \epsilon_{\lambda}^{q+a'h} Z_{\lambda}^{hb'} \rangle}{\langle \epsilon_{\lambda}^{q} \rangle \langle \epsilon_{\lambda}^{a'h} \rangle \langle Z_{\lambda}^{hb'} \rangle} = \frac{\langle \epsilon_{\lambda}^{q+a'h} \rangle}{\langle \epsilon_{\lambda}^{q} \rangle \langle \epsilon_{\lambda}^{a'h} \rangle}$$

$$= \lambda^{r_{\phi\epsilon}(q,h)} = \lambda^{\frac{c_{1,\epsilon}}{c_{\epsilon}-1}[(q+a'h)^{\alpha_{\epsilon}} - (q)^{\alpha_{\epsilon}} - (a'h)^{\alpha_{\epsilon}}]}$$
(13)

leading to :

210 
$$\forall h, q \; \frac{C_{1,\phi}}{\alpha_{\phi} - 1} [(qa+h)^{\alpha_{\phi}} - (qa)^{\alpha_{\phi}} - (h)^{\alpha_{\phi}}] = \frac{C_{1,\epsilon}}{\alpha_{\epsilon} - 1} [(q+a'h)^{\alpha_{\epsilon}} - (q)^{\alpha_{\epsilon}} - (a'h)^{\alpha_{\epsilon}}]$$
 (14)

In the general case, Eq. 14 is not valid for any q and h. To better understand this, let us consider a given level of correlation by setting the parameters a and b. The goal is to compute a' and b' from the available parameters. The left part of Eq. 14 is

known, and after setting given values of q and h it is possible to implement the same process as in section 3.3 to determine a', b' and  $\alpha_Z$ . Fig. 5 displays the outcome of this analysis, according to the values of h and q used, for a = 0.2 in the case  $\alpha_{\epsilon} = 0.8, C_{1,\epsilon} = 0.4, \alpha_{\phi} = 0.8, C_{1,\phi} = 0.2$  (meaning that b = 0.30 and  $\alpha_Y = 0.68$ ). As it can be seen, the estimates of a'215 exhibit a dependency on q and h. The dependency is stronger on q than on h and estimates remain rather stable as long as q < 0.8. Both sides of Eq. 14 are plotted in Fig. 6 for this set of UM parameters with a = 0.2 and estimates of a' = 0.19obtained with h = q = 0.7. Expected differences are visible for the larger values of h and q. It should be mentioned that these results are presented for a bad case with strong differences between  $\alpha_{\epsilon}$  and  $\alpha_{\phi}$ . They are actually much smaller if both values 220 are closer to 2. For the specific case,  $\alpha_{\epsilon} = \alpha_{\phi} = 2$ , Eq. 14 becomes:

$$\forall h, q \ C_{1,\phi}ahq = C_{1,\epsilon}a'hq \tag{15}$$

meaning that once hq has been removed, a' is deterministically obtained once a is set and  $r_{\epsilon\phi}(q,h) = r_{\epsilon\phi}(q,h) = C_{1,\phi}ahq$ is linear with regards to h and q.



Figure 5. Estimates of a' as a function of h and q using Eq. 14 and the process described in section 3.3. Computation are carried out with  $\alpha_{\epsilon} = 0.8, C_{1,\epsilon} = 0.4, \alpha_{\phi} = 0.8, C_{1,\phi} = 0.2$  and a = 0.2. The blue horizontal grid corresponds to the value obtained with Eq. 17.

Fig. 7 illustrates the relation between the parameters retrieved by setting different values of a in the same case  $\alpha_{\epsilon} = 0.8$ ,

225

 $C_{1,\epsilon} = 0.4, \alpha_{\phi} = 0.8, C_{1,\phi} = 0.2$ . First it should be mentioned that for a given set of UM parameters, not all values of a are possible. Indeed the inequality  $0 \le \alpha_Y \le 2$  must be respected leading to  $a \le min[(\frac{C_{1,\epsilon}\alpha_{\epsilon}}{C_{1,\phi}\alpha_{\phi}})^{1/\alpha_{\phi}}, (\frac{C_{1,\epsilon}(2-\alpha_{\epsilon})}{C_{1,\phi}(2-\alpha_{\phi})})^{1/\alpha_{\phi}}]$ . In this case we must have  $a \le 0.43$ . We retrieved the expected behaviour and are able to quantify it : b decreases with increasing a



Figure 6. Both sides of Eq. 14 for  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$  in the case a = 0.2 and a' = 0.19.  $r_{\epsilon\phi}(q,h)$  is in red and corresponds to the left part of Eq. 14 while  $r_{\phi\epsilon}(q,h)$  is the right part and is in blue. Two views of the same figure are provided to improve visualization.

(Fig. 7.a); a' increases with increasing a (Fig. 7.b),  $\alpha_Y$  decreases with increasing a (Fig. 7.c), and similar behaviour are found in terms of dependency in a' for the symmetric case.

#### 230 4.2 A simplified indicator

In section 4.1, limitations of this fully symmetric framework are highlighted. However, it is possible to suggest a rather intuitive indicator enabling to extract most of the information obtained from the joint multifractal correlation analysis (i.e. the computation of r(q, h)). It corresponds to the portion of intermittency  $C_1$  of one field explained by the other :

$$IC_{\epsilon\phi} = \frac{C_{1,\phi}a^{\alpha_{\phi}}}{C_{1,\epsilon}}$$

$$IC_{\phi\epsilon} = \frac{C_{1,\epsilon}a'^{\alpha_{\epsilon}}}{C_{1,\phi}}$$
(16)

235

Both "Indicators of Correlation" (IC) are displayed Fig. 8 for the data corresponding to Fig. 7. Both curves are close, and this symmetric behaviour is what is wanted for such an indicator of correlation. Again, much closer curves are obtained with



Figure 7. Illustration of the relations between the various parameters characterizing the correlation across scale between two UM fields in the case  $\alpha_{\epsilon} = 0.8$ ,  $C_{1,\epsilon} = 0.4$ ,  $\alpha_{\phi} = 0.8$ ,  $C_{1,\phi} = 0.2$ . The dash line in (b) corresponds to the relation obtained by implementing Eq. 17

greater values of  $\alpha$  and identical ones for both  $\alpha$  equal to 2. Forcing  $IC_{\epsilon\phi} = IC_{\phi\epsilon}$  can actually be a way to find an estimate of a' once a is known without having to implement the process described above. It yields :

$$a' = \left(\frac{C_{1,\phi}}{C_{1,\epsilon}}\right)^{2/\alpha_{\epsilon}} a^{\alpha_{\phi}/\alpha_{\epsilon}} \tag{17}$$

240

Eq. 17 is actually plotted in dash line on Fig. 7.b, and provides very good estimates. Hence this *IC* appears as a good candidate for characterizing in a simple way the correlations across scales between two fields. One should keep in mind that it is mainly relevant in the case where the studied fields do not exhibit values of  $\alpha$  too small (typically < 0.8).

#### 5 Implementation on rainfall data

#### 5.1 Presentation of the data

245 The rainfall data used in this paper was collected by a OTT Parsivel<sup>2</sup> disdrometer (Battaglia et al., 2010; OTT, 2014) located on the roof of the Carnot building of the Ecole des Ponts ParisTech campus near Paris between 15 January 2018 and 9 December 2018. It is part of the TARANIS observatory of the Fresnel Platform of École des Ponts ParisTech (https://hmco.enpc.fr/portfolioarchive/fresnel-platform/). Data is collected with 30 seconds time steps. Data will only be briefly presented in this paper and



**Figure 8.** Plot of  $IC_{\epsilon\phi}$  and  $IC_{\phi\epsilon}$  as a function of *a* (Eq. 17) for the same data that is presented in Fig. 7

interested readers are referred to Gires et al. (2018b) which discusses available data base in detail along with some data samples for a similar measurement campaign.

In this paper four quantities are studied:

- R, the rain rate in  $mm.h^{-1}$
- LWC, the liquid water content in  $g.m^{-3}$
- $N_t$ , the total drop concentration in  $m^{-3}$
- 255  $D_m$ , the mass weight diameter in mm

 $N_t$  and  $D_m$  are used to characterize the Drop Size Distribution (DSD, N(D), in  $m^{-3}.mm^{-1}$ ) of the rainfall. N(D)dD is the number of drops per unit volume (in m<sup>-3</sup>) with an equivolumic diameter between D and D + dD (in mm). DSD are commonly written in the form  $N(D) = N_t f(D_m)$ , with  $D_m$  being an indicator of the shape of the DSD and  $N_t$  an indicator of the total intensity. They can be computed from the DSD as (Leinonen et al., 2012; Jaffrain and Berne, 2012):

260 
$$N_t = \int_{D_{min}}^{D_{max}} N(D) dD$$
(18)

$$D_{m} = \frac{\int_{D_{min}}^{D_{max}} N(D) D^{4} dD}{\int_{D_{min}}^{D_{max}} N(D) D^{3} dD}$$
(19)

It should be noted that the disdrometer provides data binned per class of equivolumic diameter and fall velocity, from which a discrete DSD is computed and then used to evaluate the integrals of Eq. 18 and 19 (see Gires et al., 2018b for more details).

Multifractal analysis are carried out on ensemble analysis, i.e. on average over various samples. Once rainfall events (an event is defined as a rainy period during which more than 1 mm is collected and that is separated by more than 15 min of dry conditions before and after) have been selected within the longer time series, a process similar as in Gires et al. (2016) and Gires et al. (2018a) is implemented to extract the various samples of data: "for each event (i) a sample size is chosen (a power of two, if possible); (ii) the maximum number of samples for this event is computed; (iii) the portion of the event of length equal to the sample size multiplied by the number of samples found in (ii) with the greatest cumulative depth is extracted; (iv) the extracted series is cut into various samples." Since  $D_m$  is not defined when there is no rain, only samples with no zeros are used.

Dyadic sample size are simpler to use for multifractal analysis, which results in some data not used. With the process described above, 63, 52, 38 and 22 % of the data is actually not used for sample sizes of respectively 32, 64, 128 and 256. A size of 32 time steps, corresponding to 16 min is used, to maximize the amount of data used while keeping an acceptable length for the studied time series. An example of sample for the 4 studied quantities during a rainfall event that occurred on 15 January 2018. 491 such samples are used in the analysis.





Figure 9. Illustration of the four studied rainfall quantifies for 64 long sample corresponding to 32 minutes that occurred on 15 January 2018.

#### 5.2 Joint analysis and discussion

Let us first discuss the results of the joint multifractal analysis carried out between  $N_t$  and R. The purpose is to check if the scale invariant analysis of correlations is relevant for these fields and then to quantify their correlations in this framework (i.e. write the fields as in Eq. 12-top- and estimate a, b and  $\alpha_Y$  from the two fields only).

280

Main curves are shown in Fig. 10 with  $\epsilon_{\lambda}$  being the fluctuations of  $N_t$  and  $\phi_{\lambda}$  being the fluctuations of R. The analysis directly on the field showed that they were non conservative, meaning that the TM and DTM analysis would be biased. Hence multifractal analysis was carried out on an approximation of the underlying conservative fields consisting of their fluctuations (Lavallée et al., 1993). Numerical values of the various parameters of the analysis are in Table 1 and 2. R exhibits a very

- good scaling behaviour on the whole range of scales taken into account as shown by the TM analysis where the coefficients of correlation  $r^2$  of the linear regressions for q around 1 are all greater than 0.98 (Fig. 10.b). Similar scaling behaviour were found on a previous campaign with the same devices (Gires et al., 2016). The scaling for  $N_t$  is worse, with  $r^2$ s only slightly greater than 0.9, but it remains acceptable (Fig. 10.a). We find  $\alpha_R = 1.86$  and  $C_{1,R} = 0.14$  and  $\alpha_{N_t} = 1.78$  and  $C_{1,N_t} = 0.10$ . The values of UM parameters observed mean that we are in the domain of highest relevance of the framework developed in
- 290 the previous section. For R, and to a lesser extent  $N_t$ , there is a clear departure of the fitted K(q) from the empirical one with much greater values for the fitted curve. Furthermore the empirical ones exhibit a linear behaviour from for q approx. greater than 1.5 (Fig. 10.c). Such behaviour is consistent with the expected one when a multifractal phase transition associated with sampling limitations occurs.
- The joint multifractal analysis (Eq. 4 in log-log) for q = h = 0.7 of the two studied fields is displayed in Fig. 10.d. The scaling is good with a value of  $r^2 = 0.97$  for the linear fit. It enables to estimate the exponents a and b at respectively 0.33 and 0.75 (Fig. 10.e). The corresponding *IC* is equal to 0.18. In addition to quantifying the level of correlations between the two fields, it suggests how to simulate one from the other. More precisely, once a time series of fluctuations of R is available, it is possible to simulate a realistic corresponding time series of fluctuations of  $N_t$ , by raising to the power a = 0.33 the Rseries and multiplying it with an independent random fields Y with  $\alpha = 1.76$  and  $C_1 = 0.14$  raised to the power b = 0.75, and renormalizing the ensemble. Formally it suggests the fluctuations of  $N_t$  can be written as  $\frac{R_{fluctuations}^{0.33}Y^{0.75}}{(R_{luctuations}^{0.33}Y^{0.75})}$ . Such relations opens the path for techniques to simulate fluctuations of  $N_t$  knowing only the temporal evolution of the rain rate.

Similar qualitative results are found for the other combinations, and numerical values are reported in Tab. 1. Both *LWC* and  $D_m$  exhibit a good scaling behaviour and their UM parameters are in Tab. 1. As expected given the observed values of  $\alpha$ , the *ICs* computed in one way or the other (i.e. inverting the role of  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ ) are very similar. Furthermore the values of a' found

- 305 using Eq. 17 (not shown) are very close to the ones obtain by inverting the role of the two fields. This confirms the relevancy of the framework of section 4 in this case. It appears that the correlation found between R and LWC is much stronger than between R and  $N_t$  or  $D_m$ . There is no correlation between  $N_t$  or  $D_m$  which is a hint for independence but not a proof (if would be for Gaussian variables). Note that the very bad scaling for the joint analysis of these two quantities is partially due to the very small values found for r(q, h) which is basically equal to zero. R exhibits a slightly greater correlation with  $D_m$
- 310 (IC = 0.26) than with  $N_t$  (IC = 0.18). It is the inverse for LWC with values of IC respectively equal to 0.15 and 0.27.

Field	$\alpha$	$C_1$	$r^2$ for $q = 1.5$
R	1.86	0.14	0.99
LWC	1.82	0.12	0.98
$N_t$	1.78	0.10	0.91
$D_m$	1.87	0.12	0.97

**Table 2.** Numerical output of the joint multifractal analysis of the four studied fields. For each box, using the notations of Eq. 12  $\epsilon_{\lambda}$  corresponds to the field of the column and  $\phi_{\lambda}$  to the row.

	R	LWC	$N_t$	$D_m$	
R		0.98	0.97	0.97	$r^2$
		0.82	0.33	0.45	a
		0.38	0.75	0.80	b
		0.78	0.18	0.26	IC
LWC	0.98		0.95	0.97	$r^2$
	0.93		0.44	0.36	a
	0.50		0.75	0.92	b
	0.77		0.27	0.15	IC
$N_t$	0.97	0.95		0.50	$r^2$
	0.44	0.53		0.00	a
	1.08	0.94		1.11	b
	0.17	0.27		0.00	IC
$D_m$	0.97	0.97	0.50		$r^2$
	0.51	0.37	0.00		a
	0.91	0.91	0.89		b
	0.25	0.16	0.00		IC

#### 6 Conclusions

315

In this paper, we used the framework of joint multifractal analysis to characterize the correlation across scales between two multifractal fields. We extended existing framework to Universal Multifractal and also to analyse the correlations between two fields consisting of renormalized multiplicative power law combinations of two known UM fields. In general, the resulting fields can be well approximated by UM fields. Estimates of the corresponding pseudo UM parameters can be theoretically computed by focusing on the behaviour for moments close to one. These estimates remain valid for a range of moments between  $\sim 0.6$  and  $\sim 1.6$  in the worst case. The closer the two  $\alpha$  of the initial fields are, the better is the approximation. When both  $\alpha$  are equal, the approximation is exact. An analysis technique to estimate the properties of the underlying fields



**Figure 10.** Results of joint multifractal analysis for  $\epsilon_{\lambda}$  being the fluctuations of  $N_t$  and  $\phi_{\lambda}$  being the fluctuations of R. (a) TM analysis i.e. Eq. 1 in log-log plot, for  $\epsilon_{\lambda}$ . (b) Same as in (a) for  $\phi_{\lambda}$  (c) Scaling moment functions K(q) for  $\epsilon_{\lambda}$  and  $\phi_{\lambda}$ . (d) Joint multifractal analysis (Eq. 4 in log-log) for q = h = 0.7. (e) Illustration of the estimation of a with the values r(0.7, 0.7) computed in (d).

(UM parameters and power law exponents used in the combination) was developed and validated with the help of numerical simulations.

In a second step, this analysis was used to develop an innovative framework to investigate the correlations between two UM fields. It basically consists in looking at the best parameters enabling to write one field as a power law multiplicative combination of the other field and a random one. In this context, a good candidate for a simple indicator of the strength of the correlation (called *IC*) is the proportion of intermittency of a field explained by the other one. In the general case, this framework is not symmetric, which is a limitation. However when the  $\alpha$  are typically greater than  $\sim 0.8$ , it is approximately symmetric; meaning that it is relevant to extract some information on the correlations between two fields.

325

Finally it was implemented on rainfall data collected by a disdrometer installed on the roof the Ecole des Ponts ParisTech. More precisely the correlations between R and LWC, and DSD features ( $N_t$  and  $D_m$ ) are investigated. First it should be mentioned that the scaling behaviour of both R and LWC is excellent, while the one of the DSD features is only good. The  $\alpha$ are rather similar and greater than 1.7 meaning that it is a favourable context to use the newly developed approach. It appears that the correlation between R and LWC is as expected very strong, the one between R or LWC and the DSD features is medium, and the one between  $N_t$  and  $D_m$  is basically null. Besides quantifying these correlations, the developed framework suggests a simple technique to simulate one field from the other. Indeed, it is sufficient to compute a power law multiplicative combination between one field and a random one to obtain the other. The characteristic parameters of the random field as well as the power law exponents of the relation can the obtained through a joint multifractal analysis of the two studied fields.

18

Further investigations on other fields in various context should be carried out to confirm the interest of this framework to both characterize and simulate correlations across scales between two multifractal fields. In future work, this framework should also be extended to more than two fields.

Acknowledgements. Authors greatly acknowledge partial financial support from the Chair "Hydrology for Resilient Cities" (endowed by 340 Veolia) of Ecole des Ponts ParisTech, and the Île-de-France region RadX@IdF Project.

#### References

- Battaglia, A., Rustemeier, E., Tokay, A., Blahak, U., and Simmer, C.: PARSIVEL Snow Observations: A Critical Assessment, Journal of Atmospheric and Oceanic Technology, 27, 333–344, http://dx.doi.org/10.1175/2009JTECHA1332.1, 2010.
- Bertol, I., Schick, J., Bandeira, D. H., Paz-Ferreiro, J., and Vázquez, E. V.: Multifractal and joint multifractal analysis of wa ter and soil losses from erosion plots: A case study under subtropical conditions in Santa Catarina highlands, Brazil, Geo derma, 287, 116 125, https://doi.org/https://doi.org/10.1016/j.geoderma.2016.08.008, http://www.sciencedirect.com/science/article/pii/
   S0016706116303470, structure and function of soil and soil cover in a changing worl: characterization and scaling, 2017.
  - Calif, R. and Schmitt, F. G.: Multiscaling and joint multiscaling description of the atmospheric wind speed and the aggregate power output from a wind farm, Nonlinear Processes in Geophysics, 21, 379–392, https://doi.org/10.5194/npg-21-379-2014, http://www.
- nonlin-processes-geophys.net/21/379/2014/, 2014.
  - Douglas, E. M. and Barros, A. P.: Probable maximum precipitation estimation using multifractals: Application in the eastern United States, Journal of Hydrometeorology, 4, 1012–1024, <GotoISI>://WOS:000187534900003, j. Hydrometeorol., 2003.
    - Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Multifractal comparison of the outputs of two optical disdrometers, Hydrological Sciences Journal, 61, 1641–1651, https://doi.org/10.1080/02626667.2015.1055270, https://doi.org/10.1080/02626667.2015.1055270, 2016.
- 355 Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Pseudo-radar algorithms with two extremely wet months of disdrometer data in the Paris area, Atmospheric Research, 203, 216 – 230, https://doi.org/https://doi.org/10.1016/j.atmosres.2017.12.011, http://www.sciencedirect. com/science/article/pii/S0169809517307111, 2018a.
  - Gires, A., Tchiguirinskaia, I., and Schertzer, D.: Two months of disdrometer data in the Paris area, Earth System Science Data, 10, 941–950, https://doi.org/10.5194/essd-10-941-2018, https://www.earth-syst-sci-data.net/10/941/2018/, 2018b.
- 360 Hubert, P., Tessier, Y., Ladoy, P., Lovejoy, S., Schertzer, D., Carbonnel, J. P., Violette, S., Desurosne, I., and Schmitt, F.: Multifractals and extreme rainfall events, Geophys. Res. Lett., 20, 931–934, 1993.
  - Jaffrain, J. and Berne, A.: Influence of the Subgrid Variability of the Raindrop Size Distribution on Radar Rainfall Estimators, Journal of Applied Meteorology and Climatology, 51, 780–785, http://dx.doi.org/10.1175/JAMC-D-11-0185.1, 2012.
- Jiménez-Hornero, F., Pavón-Domínguez, P., de Ravé, E. G., and Ariza-Villaverde, A.: Joint multifractal description of the relationship between wind patterns and land surface air temperature, Atmospheric Research, 99, 366 – 376, https://doi.org/https://doi.org/10.1016/j.atmosres.2010.11.009, http://www.sciencedirect.com/science/article/pii/S0169809510003121, 2011.
  - Kahane, J.: Sur le Chaos Multiplicatif, Ann. Sci. Math. Que., 9, 435–444, 1985.

Lavallée, D., Lovejoy, S., and Ladoy, P.: Nonlinear variability and landscape topography: analysis and simulation, in: Fractas in geography,

- aro edited by de Cola, L. and Lam, N., pp. 171–205, Prentice-Hall, 1993.
  - Leinonen, J., Moisseev, D., Leskinen, M., and Petersen, W. A.: A Climatology of Disdrometer Measurements of Rainfall in Finland over Five Years with Implications for Global Radar Observations, Journal of Applied Meteorology and Climatology, 51, 392–404, http://dx. doi.org/10.1175/JAMC-D-11-056.1, 2012.
- Mandelbrot, B. B.: Intermittent turbulence in self-similar cascades: divergence of high moments and dimension of the carrier, Journal of Fluid Mechanics, 62, 331–358, https://doi.org/10.1017/S0022112074000711, 1974.
  - Meneveau, C., Sreenivasan, K. R., Kailasnath, P., and Fan, M. S.: Joint multifractal measures: Theory and applications to turbulence, Phys. Rev. A, 41, 894–913, https://doi.org/10.1103/PhysRevA.41.894, https://link.aps.org/doi/10.1103/PhysRevA.41.894, 1990.

Schertzer, D. and Lovejoy, S.: Physical modelling and analysis of rain and clouds by anisotropic scaling and multiplicative processes, J. Geophys. Res., 92, 9693–9714, 1987.

380 Schertzer, D. and Lovejoy, S.: Hard and soft multifractal processes, Physica A, 185, 187–194, 1992.

- Schertzer, D. and Lovejoy, S.: Universal multifractals do exist!: Comments, Journal of Applied Meteorology, 36, 1296–1303, <GotoISI>: //WOS:A1997XU62400016, j. Appl. Meteorol., 1997.
  - Schertzer, D. and Lovejoy, S.: Multifractals, generalized scale invariance and complexity in geophysics, International Journal of Bifurcation and Chaos, 21, 3417–3456, <GotoISI>://WOS:000300016000003, 2011.
- 385 Schertzer, D. and Tchiguirinskaia, I.: Multifractal vector fields and stochastic Clifford algebra, Chaos, 25, 123127, https://doi.org/DOI:http://dx.doi.org/10.1063/1.4937364, 2015.
  - Schertzer, D. and Tchiguirinskaia, I.: An Introduction to Multifractals and Scale Symmetry Groups, in: Fractals: Concepts and Applications in Geosciences., CRC Press, edited by Ghanbarian, B. and Hunt, A., pp. 1–28, 2017.
- Schertzer, D. and Tchiguirinskaia, I.: A century of turbulent cascades and the emergence of multifractal operators (invited paper under review), Earth and Space Science, 2019.
- Seuront, L. and Schmitt, F. G.: Multiscaling statistical procedures for the exploration of biophysical couplings in intermittent turbulence. Part I. Theory, Deep Sea Research Part II: Topical Studies in Oceanography, 52, 1308 – 1324, https://doi.org/https://doi.org/10.1016/j.dsr2.2005.01.006, http://www.sciencedirect.com/science/article/pii/S0967064505000470, observations and modelling of mixed layer turbulence: Do they represent the same statistical quantities?, 2005a.
- 395 Seuront, L. and Schmitt, F. G.: Multiscaling statistical procedures for the exploration of biophysical couplings in intermittent turbulence. Part II. Applications, Deep Sea Research Part II: Topical Studies in Oceanography, 52, 1325 – 1343, https://doi.org/https://doi.org/10.1016/j.dsr2.2005.01.005, http://www.sciencedirect.com/science/article/pii/S0967064505000482, observations and modelling of mixed layer turbulence: Do they represent the same statistical quantities?, 2005b.
- Siqueira, G. M., Ênio F.F. Silva, Vidal-Vázquez, E., and Paz-González, A.: Multifractal and joint multifrac tal analysis of general soil properties and altitude along a transect, Biosystems Engineering, 168, 105 120, https://doi.org/https://doi.org/10.1016/j.biosystemseng.2017.08.024, http://www.sciencedirect.com/science/article/pii/
   S1537511016305943, computational Tools to Support Soil Management Decisions, 2018.
- Wang, Z.-Y., Shu, Q.-S., Xie, L.-Y., Liu, Z.-X., and Si, B.: Joint Multifractal Analysis of Scaling Relationships Between Soil Water-Retention Parameters and Soil Texture, Pedosphere, 21, 373 – 379, https://doi.org/https://doi.org/10.1016/S1002-0160(11)60138-0, http://www.sciencedirect.com/science/article/pii/S1002016011601380, 2011.
- Xie, W.-J., Jiang, Z.-Q., Gu, G.-F., Xiong, X., and Zhou, W.-X.: Joint multifractal analysis based on the partition function approach: analytical analysis, numerical simulation and empirical application, New Journal of Physics, 17, 103 020, http://stacks.iop.org/1367-2630/17/i=10/a=103020, 2015.

### Answer to referee #1

Authors would like to thank the anonymous reviewer for his/her very careful reading of the paper and suggestions to improve it. Hopefully the modifications implemented will satisfy him/her.

This paper about joint multifractal analysis applied to rainfall data, has a theoretical and technical part, and an application on some rainfall data. I have several comments below. Some parts present confusions on the notations.

Line 25. I suggest to modify the sentence. It is not proved that all multifractal processes converge to UM (universal multifractals). There are many multifractal models that do not belong to UM. Following your comment, the sentence was updated to "In the large class of Universal Multifractals (UM) which are the stable and attractive limits of non-linearly interacting multifractal processes and correspond to a broad, multiplicative generalization of the central limit theorem; Schertzer and Lovejoy, 1987, 1997)"

### Line 81. For divergence of moments cite also Mandelbrot (1974) and Kahane (1985).

I guess that you are referring to these two papers:

- Kahane, J.P., Sur le Chaos Multiplicatif, Ann. Sci. Math. Que., 9, 435-444, 1985.
- Mandelbrot, B. Intermittent turbulence in self-similar cascades: Divergence of high moments and dimension of the carrier, J. Fluid 1036-1038, 1987.

These papers are cited in Schertzer and Lovejoy (1987), which specifically describes this effect in the framework of UM. Anyway, we included citations of these papers. In addition, we will mention that they did not address the quantification of the spurious statistical estimates on finite samples and their dependence on their size (Schertzer and Lovejoy 1992).

:Equation (4). There is a mixture between p, q and h. Please double check this, and also in other parts of the manuscript, to have consistent notations.

This was updated, thank you for your careful reading

Equation (5). Do the authors restrict to a > 0 and b > 0? Yes and this was clarified in the text.

Figure 1. I recommend to plot X  $\lambda$  and epsilon  $\lambda$  since one does not understand what is the blue field.

Since we assumed that phi\_lambda = X\_lamdba, this is actually what is plotted. It was clarified in the caption and reference to phi was removed as well (see answer to previous comment).

Section 3. Why not indicate from the beginning that the aim is to study the relation between X  $\lambda$  and epsilon  $\lambda$ . I do not understand the use of  $\varphi$  here, and also I do not believe in the sentence line 96 "without loss of generality". It is here an hypothesis, it is not a general situation.

As suggested by the reviewer, to improve clarity, references to phi were removed throughout the section.

Equation (7). Second line, the prefactor of the second term is not correct (b  $\alpha$  y and not a  $\alpha$  x ) Indeed, this was corrected.

Line 128. Where the strange value q D = 91 comes from? This is much too large.

Values of  $q_D$  are obtained by solving this equation  $K(q_D)=(q_D-1)D$  using the pseudo UM parameters of epsilon\_lambda. This was clarified in text and values were computed for each panel of Fig. 2. We find values equal to 5.96, 4.68 and 119 (meaning that a wrong value was written in the first version of the paper).

Equation (9). Some mistakes: insert two minus signs and last term is K(q) and not K(a) Thank you for your careful reading, this was corrected.

Equation (10). Equation (8) is given for the field epsilon, not for Y. Explain better how this equation is used to obtain  $\alpha$  y. Indeed in equation (8)  $\alpha$  y is nonlinearly related to other variables and it does not seem easy to isolate its expression. Same for equation (11). Where does this come from?

Indeed some clarifications were missing and they have been added:

- For Eq. 10 this parenthesis has been added : (noting that  $\lambda = C_{1,X} a^{\lambda} = C_{1,X} a^{\lambda} = C_{1,Y} b^{\lambda} + C_{1,Y} b^{\lambda$ 

- For Eq. 11 this parenthesis has been added : (noting that  $C_{1,Y} b^{\Delta_Y}=C_{1,\nabla_Y}=C_{1,Y}=C_{1,X}$  and that we have  $C_{1,Y}=C_{1,X}$ )

## Section 3.4:

Line 151. Why the use of discrete cascades? The term is not explained. Why not continuous cascades?

The following sentences were added to explain discrete cascades.

"The approach presented above is tested on numerical simulations obtained with discrete in scale cascades.

It consists in iteratively repeating a cascade step with a non infinetisimal scale ratio in which a 'parent' structure is divided into 'daughter' structures whose affected value is the one of the 'parent' structure multiplied by a random factor ensuring that Eqs. 1 and 2 remain valid. Such simple field generation process is sufficient for the purposes of this paper. The recent introduction of multifractal operators and vectors paves the way for physically-based, continuous (in scale) multivariate analysis of multifractal fields or measures (Schertzer and Tchiguirinskaia 2015, 2019)"

- Schertzer, D. and Tchiguirinskaia, I. (2015) 'Multifractal vector fields and stochastic Clifford algebra', *Chaos*, 25(12). doi: 10.1063/1.4937364.
- Schertzer, D. and Tchiguirinskaia, I. (2019) 'A century of turbulent cascades and the emergence of multifractal operators', Earth and Space Science (invited paper under review)

### Line 157. What is DTM analysis?

It is actually defined in section 1: "Double Trace Moment (DTM), specifically designed for UM fields, is commonly used to estimate UM parameters (Lavallée et al.,1993)". Authors thinks that this description is sufficient for the purposes of this paper, but if the reviewer still thinks it should be completed, it can obviously be done.

Line 166 and further. Explain better the objectives and hypotheses of the numerical work. I understand that X and Y are simulated, epsilon is built with some values of a and b. Then the exercise is (i) to find the approximate values of  $\alpha$ \_epsilon and C1\_epsilon and (ii) to assume that epsilon and X are known, and try to find a, b and  $\alpha$  y. Is this correct? If yes it should be clearly stated in the text.

You are indeed correct. Following your comment, clarifications were added:

Lines 175-179. A quantification of the error is needed.

It is actually displayed by Fig. 4.

Section 4.1. This is very technical and of poor interest. It could be moved to an appendix. It is indeed quite technical. But authors believe it might be better to keep it in the main part of the paper because it highlights the limitations of the developed framework and enables to introduce the simplified indicator.

## Section 5:

Line 276 and further. Explain better the hypothesis of joint multifractal analysis. What is assumed to be known, what is the objective of the work, what is assumed, what is known and unknown.

The following sentence was added to clarify the study "The purpose is to check if the scale invariant analysis of correlations is relevant for these fields and then to quantify their correlations in this framework (i.e. write the fields as in Eq. 13-top- and estimate \$a\$, \$b\$ and \$\alpha\_Y\$ from simply the two fields)."

## Line 280 is not "multiplying" but "raising to the power"

The sentence was re-written to insert this correction.

Line 280 and further. Do you obtain  $N_t = R^{1/3} X^{3/4}$ ? Where X is an unknown field? If yes the equation should be written down and more interpretation should be given to this proposed relation.

Yes it is indeed correct, and it is now written down. And comments added on the implications, notably in terms of numerical simulations.

References: why some references have a web reference, some have two web references, and some have none. There is a text in capital letters in the second reference, that should be removed. The capital letters were removed. With regards to the web references, it was done automatically from my bibtext library which could need to the updated for some web references. This can be done at the editorial stage.

Authors would like to thank the anonymous reviewer for his/her very careful reading of the paper and suggestions to improve it. Hopefully the modifications implemented will satisfy him/her.

Interactive comment on "Joint multifractal analysis: further developments and implementation on rainfall data" by Auguste Gires et al.

Anonymous Referee #2

Received and published: 24 November 2019

This work studies the behaviour of fields which are composed of a product of two universal multifractal (UM) fields. First, the properties of UM fields are briefly reviewed. Then the properties of multiplicative combinations of UM are discussed and it is shown how approximate UM parameters can derived from products of UM fields. The authors warn for the possible confusion between the phase transition causing diverging scaling moment functions K(q) and the combined nature of the field, both of which give rise to K(q) which are higher than predicted by UM theory. The authors then perform a numerical experiment with the discussed set-up of one UM field  $\varphi$  and one combined field epsilon. They estimate the parameters of the underlying fields using their newly developed methodology, and demonstrate the use of a simplified correlation indicator. The validity of the approach seems to be constrained to UM fields with sufficiently similar values of  $\alpha$  in this symmetric case.

The technique is then applied to observational rainfall data from a disdrometer to infer correlations between different properties such as rain rate, liquid water content, drop concentration and mass weighed diameter. For these fields the validity ranges of the parameters seem to be well respected. The result of such an analysis can be used to simulate one of these quantities, based on another known quantity and a random one.

General comments:

This paper shows a new technique to infer the properties of multiplicative fields, which could be useful to investigate correlations between UM fields and simulate a field based on a given one, if the correlation is known. The application to rainfall data nicely highlights the potential of this method.

Thank you for your positive feedback.

The title does not capture the subject of the paper, that is the analysis of correlation between approximate UM fields. "Further developments" is very vague for a title. I would also say "application to" instead of "implementation on".

Following your comment, the title was changed to : "Approximate multifractal correlation and products of Universal Multifractal fields, with application to rainfall data"

The structure of the manuscript is fine, the formalism is explained clearly and the results are shown in a logical way. The figures could be improved somewhat (see specific comments below). The equations, however, contain errors. I hope these are merely typographical, but to remove any doubts on the correctness of the results I suggest the authors provide their code and/or data as supplementary material or through a citable repository (e.g. Zenodo). This would also be in accordance with the best practices of this journal.

As suggested by the referee, code and/or data will be made available on a citable repository.

Finally, there are many grammatical and spelling errors throughout the manuscript (e.g. "betwen", "dash line", ...). Articles seem to be missing, e.g. p.2 l.44: Similar formalism -> A similar formalism. Please check the whole manuscript carefully for spelling and grammar; the list below is not complete.

These corrections as well as the ones below were implemented. The manuscript was also carefully checked for spelling and grammar.

Specific comments: p.1 l.2: across wide -> across a wide p.1 l.9: to retrieved -> to retrieve p.2 l.24: Reader is -> The reader is p.2 l.42: of define -> to define p.2 l.50: relying this -> relying on this p.3 l.68: an homogeneous -> a homogeneous p.3 l.59: Please specify the "outer scale" more clearly. p.4 l.88: as follow -> as follows This was corrected, thank you for your careful reading

p.4 Eq. (40): I think the RHS should read

 $\lambda^{S(h,q)-K_{\epsilon}(q)-K_{\phi}(h)} \approx \lambda^{r(h,q)}$ 

This was corrected (it was simply a typographical error)

Fig. 1: Spurious "=" in the caption.

p.5 Eq. (7): in the second line, the second term should start with b  $\alpha$  Y, not a  $\alpha$  Y. p.7 l.156 Please mention the meaning of TM again here for clarity l.157 Please mention the meaning of DTM again here for clarity This was corrected

p.7 l.161: The fact that the empirical K(q) in section 3.4 are lower than expected seems in contradiction with earlier remarks that the empirical K(q) would in both cases be higher than expected: please clarify this or clearly disentangle the two kinds of phase transition that can occur. Indeed the two kind of multifractal phase transitions discussed result in different behaviour of the empirical K(q) with regards to the theoretical one. Following the reviewer's comment, this was clarified in the section 2.1.

p.7 l.158, 163 and 172: "inputted" does not exist This was corrected and changed to "input".

Fig. 4: It would be helpful to visualize the line ha + q = q s on the surface (mentioned in p.8 l.177) ha $+q=q_s$  would actually be another surface. So authors have the feeling it would not improve visualization to add another surface on the figure.

Fig. 5: It would be helpful to visualize the intersection between the two planes. The orientation of the figure has been changed to improve visualisation.

p.10. Eqns. (12) and (13) are not consistent with each other. For the first line of Eq. (13), for example, I obtain:

$$\frac{\langle \phi_{\lambda}^{aq+h} \rangle}{\langle \phi_{\lambda}^{aq} \rangle \langle \phi_{\lambda}^{h} \rangle}.$$

 $\frac{\langle \epsilon_{\lambda}^{a'h+q} \rangle}{\langle \epsilon_{\lambda}^{a'h} \rangle \langle \epsilon_{\lambda}^{q} \rangle}$ 

For the third line I obtain

and likewise for Eq. (14) and what follows. Please check carefully whether this affects the presented results. Also verify whether a and a' are not swapped in the rest of the manuscript (e.g. Eq. (16))

Indeed q and h were reversed in the mentioned equations. This was corrected. It means the the axis legend were reversed for Fig. 5 and 6 which are updated. It does not affect the other results since estimates are obtained with q=h=0.7

Fig. 6: It would be helpful to visualize the intersection between the two planes. Also it seems that the blue plane is covering the red plane where I would expect the red plane to be visible. Please improve this figure and mention the meaning of the different colours in the caption. The figure was updated (see previous comment) and two views are now provided to improve visualization. Indeed the meaning of the colours was missing and is now in the caption.

Table 2 caption: "using the notations of 12" -> "using the notations of Eq. (12), "; "line" -> "row" p.16 l.286: the one obtain -> the ones obtained p.17 l.298: "the two the" -> "the two" This was corrected.

p.18 l.315: "The characteristic parameters [...] as long as the power law exponents [...] can the obtained through [...] of the studied fields." I don't understand this sentence, please correct. It should have been "as well as" and not "as long as". This has been corrected.