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5	Role of nonlinear interaction between water and plant in
6	stability analysis of nonspatial plants
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8	Guodong Sun ^{1, 3,*} , Xiaodong Zeng ²
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11	¹ State Key Laboratory of Numerical Modeling for Atmospheric Sciences and Geophysical Fluid Dynamics
12	(LASG), Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China
13	² International Center for Climate and Environment Sciences (ICCES), Institute of Atmospheric Physics,
14	Chinese Academy of Sciences, Beijing 100029, China
15	³ University of Chinese Academy of Sciences, Beijing 100049, China
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21	*Corresponding author: Guodong Sun (sungd@mail.iap.ac.cn)
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Abstract: In this study, a theoretical ecosystem model is applied to discuss a 27 stability of plant using linear and nonlinear methods. Two common linear methods are 28 29 employed to analyze a linear stability of plant through judging the positive or negative 30 of eigenvalues (Lyapunov method), and solving a linear singular vector (LSV). To explore the nonlinear stability of plant, a conditional nonlinear optimal perturbation 31 (CNOP) approach is used. The CNOP, which is a type of initial perturbation, could 32 cause the nonlinearly most unstable for an equilibrium state. The CNOP is a nonlinear 33 34 development of the LSV which is the rapidest initial perturbation with a linear framework. The numerical results show that two linear stable equilibrium states (plant 35 and desert) with the linear methods are nonlinear unstable with the CNOP method. 36 When there is large enough magnitude of initial perturbation, the linear stable plant 37 (desert) equilibrium state will be evolved into desert (plant) equilibrium state using 38 the CNOP-type initial perturbation. This character disappears using the LSV-type 39 40 initial perturbation. The above results are effective for two types of plant, namely grasslands and trees. Through analyzing the nonlinear dynamics of the theoretical 41 model, it is found that the nonlinear interaction between plant and water play more 42 important role to a transitions between two equilibriums states than the evaporation 43 and the plant losses expressed by linear terms in the theoretical model. The findings 44 could be exhibited by using the nonlinear method (the CNOP method), but fail by 45 using the linear methods. 46

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56 1. Introduction

Arid and semiarid regions cover more than 40% of the globe, and are extremely 57 sensitive to environmental condition and human activities (Charney, 1975; Fu and An, 58 2002; Sankey et al., 2012; Huang et al., 2017). There are already plenty of evidences 59 that plant degeneration often occurs in arid and semi-arid regions (Ni, 2004; Okin et 60 al., 2009; Sun et al., 2017). The instability of plant impacts not only animal 61 biodiversity, soil productivity, and so on, but largescale climate change, the transfer of 62 radiation, water and energy, etc (Xue and Shukla, 1993; Xue, 1996; Eklundh and 63 Olsson, 2003; Lu and Ji, 2006; Notaro et al., 2006; Piao et al., 2007; Xu et al., 2012). 64 Hence, understanding the drivers and dynamics of plant stability in arid and semi-arid 65 region is motivated by analyzing the effect of disturbances regimes on plant 66 67 degradation.

Abrupt changes of plant in arid and semi-arid regions may be considered as the 68 69 transitions between two stable equilibrium states (Mauchamp et al., 1994). Precipitation was thought to the key factors to sharp transitions between different 70 71 vegetation states (Hardenberg et al., 2001; Motchell and Csillag, 2001; Sun and Mu, 72 2013, 2014). It had been reproduced to the transitions from bare soil at limited rainfall 73 to homogeneous vegetation at high rainfall observed in arid and semi-arid regions 74 using a theoretical model. For same rainfall, the transitions between different stable 75 equilibrium states are also discovered. Zeng et al. (2004, 2005, 2006) build a theoretical model, which could reconstruct different plant pattern along a moisture 76 index in North China, to reveal the coexistence of the grassland and the desert 77 78 equilibrium states at the same climate condition. The stabilities of the grassland and the desert equilibrium states were demonstrated by using the Lyapunov method. The 79 shading mechanism of the wilted biomass was announced as the key mechanism of 80 the maintenance of the grassland equilibrium state. The transition between two plant 81 equilibrium states is also investigated. Okin (2009) proposed a theoretical model to 82 examine the grassland-shrubland dynamics. Their finding suggested that a feedback 83 between grass biomass and soil erosion may cause an abrupt transition from grassland 84 to a shrubland state observed throughout the southwestern U.S. in the past 150 years. 85





86 The bistability character was explained by the theoretical model and the Lyapunov method. However, the above models are the nonlinear model. It is inappropriate that 87 the linear method (Lyapunov method) is applied to explore the stability of equilibrium 88 states. A nonlinear stability analysis method (the condition nonlinear optimal 89 perturbation method, the CNOP method, Mu et al., 2003) was employed to illuminate 90 a nonlinear stability of the grassland and the desert equilibrium states (Mu and Wang, 91 92 2007; Sun and Mu, 2009, 2011; Sun and Xie, 2017). The CNOP is a type initial perturbation, which could cause the most unstable state compared to the linear stable 93 grassland and the desert equilibrium states. If the CNOP could bring to the transition 94 from one linear stable equilibrium state to another, the linear stable equilibrium state 95 was considered as the nonlinear stable. Mu and Wang (2007), and Sun and Mu (2009, 96 2011) found the CNOP-type initial perturbation, which brought to the transition 97 between two equilibrium states. The nonlinear mechanism played the key role in the 98 99 transition between two equilibrium states opened out by their studies.

The purpose of this report is to investigate the nonlinear stability of equilibrium state, and reveal which dynamics mechanism is important to transition between two equilibrium states by using the CNOP method and a theoretical model. We argue further the roles of the linear terms and the nonlinear terms in the transition processes.

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105 2. Model and method

106 2.1. The model

A simple theoretical model, which could descript the plant and water dynamics,
and simulate the different types of plant patterns, is employed to explore the stability
of plant (Klausmeier, 1999).

110 The model is presented as follows:

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$$\begin{cases}
\frac{\partial W}{\partial T} = A - LW - RWN^2 + V \frac{\partial W}{\partial X} \\
\frac{\partial N}{\partial T} = RJWN^2 - MN + D(\frac{\partial^2}{\partial X} + \frac{\partial^2}{\partial Y})N
\end{cases}$$
(1)

112 W and N represent water and plant biomass in the two-dimensional domain indexed





- by X and Y. A is the rate of water input, L is the rate of evaporation. The expression of 113
- RWN^2 means the rate of plants taking up water, and is a nonlinear term. V is speed of 114
- water flowing downhill. J is the yield of plant biomass per unit water consumed. Plant 115
- biomass misses according to the density-independent mortality and maintenance at 116
- rate MN. Plant dispersal is simulated by a diffusion term with diffusion coefficient D. 117
- The above theoretical model could be nondimensionalized (Klausmeier, 1999) to 118

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$$\begin{cases} \frac{\partial w}{\partial t} = a - W - wn^2 + v \frac{\partial w}{\partial x} \\ \frac{\partial n}{\partial t} = wn^2 - mn + (\frac{\partial^2}{\partial x} + \frac{\partial^2}{\partial y})n \end{cases}$$
(2)

In this study, to explore the linear and nonlinear stability of nonspatial plant, the 120 121 space derivatives are set to zero. So, the below theoretical model is analyzed

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$$\begin{cases} \frac{\partial w}{\partial t} = a - w - wn^2 \\ \frac{\partial n}{\partial t} = wn^2 - mn \end{cases}$$
 (3)

- 124 2.2 Conditional nonlinear optimal perturbation method (CNOP)
- To determine whether the nonlinearly stability or instability of the plant or not, 125
- the conditional nonlinear optimal perturbation approaches related to initial errors 126
- (CNOP) are applied (Mu et al., 2003, 2004). The CNOP method is introduced below 127
- 128 for the convenience of the reader.

The target problem can be represented in following ordinary or partial 129

differential equation: 130

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$$\begin{cases} \frac{\partial U}{\partial t} = F(U, P) & U \in \mathbb{R}^n, t \in [0, T] \\ U\Big|_{t=0} = U_0 \end{cases}$$
(4)

where F is a nonlinear operator; P represents the model parameters; U_0 132





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 $U(\tau) = M_{\tau}(U_0).$

- 133 contains the initial values of the state variables; M_{τ} represents the propagator of the 134 ordinary or partial differential equations from the initial time 0 to τ ; and U_{τ} is a 135 solution of the ordinary or partial differential equations at time τ that satisfies
- Let $U(T;U_0)$ and $U(T;U_0)+u(T;u_0)$ be the solutions of the ordinary or partial differential equations (2) with the initial and model vectors U_0 and U_0+u_0 . u_0 indicates the errors and perturbations related to the initial values and model parameter values. $u(T;u_0)$ describes the variations in the reference state $U(T;U_0)$ caused by the initial errors and the model parameter errors u_0 . $u(T;U_0)$ satisfies the following conditions:
- 143 $\begin{cases} U(T;U_0) = M_T(U_0) \\ U(T;U_0) + u(T;u_0) = M_T(U_0 + u_0) \end{cases}$ (5)

144 For the chosen norm $\| \|$, a perturbation u_{σ} is the CNOP if and only if

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$$J(u_{\sigma}) = \max_{u_{\sigma} \in \sigma} J(u_0), \quad (6)$$

146 where

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$$J(u_0) = \left\| M_T(U_0 + u_0) - M_T(U_0) \right\|$$
(7)

Here, U_0 is the reference state; u_0 is the error of the initial conditions; $u_0 \in \sigma$ is the constraint condition. So, the CNOP represents a type of initial errors, which could cause the most unstable state. To obtain the maximum value of (2), the sequential quadratic programming (SQP) algorithm (Barclay et al., 1997) is employed. The gradients of the cost function are computed by the definition of the gradient.

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- 154 2.2. Experimental design





155 A control factor in the theoretical model is water input (a). There are different equilibriums for different water input. According the climate character of arid and 156 semi-arid (rainfall is about from 250 Kg H_2O m⁻² year⁻¹ ~ 750 Kg H_2O m⁻² year⁻¹). If 157 the water input a is from 0.077 to 0.23, the plant is tree, and m is 0.045 (see appendix 158 for details). For the grassland state, a is from 0.94 to 2.81, and m is 0.45. In our 159 160 reports, two water inputs (a=1.2, 0.2) and two optimization times (T=20 and 30, 20 and 30 years) are considered in order to determine whether or not the numerical 161 results are dependent upon the choices of the reference state and the optimization time. 162 The model is discretized based on the fourth-order Runge-Kutta method with a time 163 step of dt=1/24 (representing half of a month). L2 norm is chosen, and the constrained 164 condition about the initial perturbation is $||u_0|| \le \delta$ 165

To analyze the linear stability of plant, a traditional method is used to judge the 166 positive and negative of eigenvalues. If the eigenvalues are positive (negative), the 167 plant or desert is stable (unstable). To analyze the nonlinear stability of plant or desert, 168 the CNOP is calculated to determine the nonlinear evolution of the initial perturbation. 169 If the nonlinear evolution of the initial perturbation will be zero, the plant or desert is 170 nonlinear stable. On the contrary, an abrupt change occurs, and the plant or desert is 171 nonlinear unstable. In the same way, a linear singular vector (LSV) method is used to 172 compare to the CNOP method. 173

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175 3. Numerical results

176 3.1. Linear stability analysis

For the theoretical model, there are three equilibrium states. One is a desert state, and two is plant state. If the water input *a* is from 0.077 to 0.23, the plant is tree, and *m* is 0.045. For the grassland state, *a* is from 0.94 to 2.81, and *m* is 0.45. Table 1 shows examples about the stability analysis for different equilibrium states of grassland (a=1.2) and tree (a=0.2). It is found that there are three equilibrium states for grassland or tree. One is the linearly stable grassland or tree equilibrium state, and another is the linearly stable desert equilibrium state due to negative







184 eigenvalues. The final one is the linearly unstable grassland or tree equilibrium state. Figure 1 shows the equilibrium states and stability analysis for different 185 water inputs within grassland or tree. 186

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188 3.2. Nonlinear stability analysis

To analyze the nonlinear stabilities of equilibrium states, the CNOP method 189 190 and the LSV method are employed. If the grassland (desert, a=1.2) equilibrium state is transformed into the desert (grassland) equilibrium state due to the 191 CNOP-type initial perturbation. The linear stable grassland (desert) equilibrium 192 state is considered as nonlinearly unstable. When the constrained value (δ) is 1.0, 193 194 it is found that the grassland equilibrium state is transformed into the desert equilibrium state due to the CNOP-type initial perturbation (w'=-0.489, n'=-0.872) 195 (Figure 2a and 2b). However, the grassland equilibrium state fails to be 196 197 transformed into the desert equilibrium state due to the LSV-type initial perturbation (w'=-0.741, n'=-0.671) for the same constrained value $\delta = 1.0$. There 198 are similar results for the desert equilibrium state as the reference state. The above 199 200 analysis results imply that the linearly stable grassland (desert) equilibrium state is 201 nonlinearly unstable for grassland (a=1.2).

202 In addition, the tree and desert equilibrium states (a=0.2) are also analyzed. It 203 is shown that the desert equilibrium state will be transformed into the tree desert equilibrium state due to the CNOP-type initial perturbation (w'=0.0125, 204 n'=0.2397). However, the desert equilibrium state is kept due to the LSV-type 205 206 initial perturbation (w'=0.0721, n'=0.2290). Our findings also illuminated that the turnover time for the tree (about 200 years) as the reference state is longer than that 207 for the grassland (about 20 years) as the reference state. The variational ratios of 208 three terms about the plant and the water in the theoretical model for the grassland 209 210 are larger than those for the trees. The ratio of plant biomass losses (m) for the grassland is larger than that for the tree. The above results also hint that the tree is 211 more stable than the grassland. 212





- 4. Discussions
- 215 To analyze the dynamics of plant due to two types of initial perturbations, the
- 216 nonlinear model (Eq. 8) and linear model (Eq. 9) of initial perturbation are showed for
- 217 the Eq. 3.

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$$\begin{cases} \frac{\partial w'}{\partial t} = -w' - (\overline{w} + w')(\overline{n} + n')^2 + \overline{w}\overline{n}^2 \\ \frac{\partial n'}{\partial t} = (\overline{w} + w')(\overline{n} + n')^2 - \overline{w}\overline{n}^2 - mn' \\ \frac{\partial w'}{\partial t} = -w' - w'\overline{n}^2 - 2nwn' \end{cases}$$
(8)

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$$\begin{cases} \frac{\partial t}{\partial t} = -w - w n - 2nwn \\ \frac{\partial n'}{\partial t} = w' \overline{n}^2 + 2\overline{nw}n' - mn' \end{cases}$$
 (9)

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221 \overline{w} and \overline{n} are the basic state of the water and the plant. w' and n' are the perturbation of the water and the plant. $(\overline{w} + w')(\overline{n} + n')^2 - \overline{w}\overline{n}^2$ 222 represents the nolinear term about plants taking up water of initial perturbation. $w'\overline{n}^2 + 2\overline{nw}n'$ is 223 224 the linearization of the nonlinear ter. It is shown that the nonlinear evolutions of the 225 CNOP and the LSV with Eq. 8 are similar to those of the CNOP and the LSV being 226 superimposed on the basic state (Figure 2). In addition, the linear evolution of the 227 LSV is also discussed. We find that the character is similar between the nonlinear 228 evolution and linear evolution about the LSV. This suggests that the CNOP could 229 cause the nonlinear stability for the nonlinear model, however the LSV fails under the 230 same extent of constrained condition.

To analyze the difference of the dynamical mechanisms about the CNOP and the LSV, the right terms of Eq. 8 and 9 are computed. Figure 4 shows the evolutions of every term. For the grassland as the reference state, it is found that the variation extent (0.98) of the nonlinear term $(\overline{w} + w')(\overline{n} + n')^2 - \overline{w}\overline{n}^2$, which represents nonlinear interaction between water and plant in arid and semi-arid region, caused by the CNOP is greater than those of the linear terms w' (0.87) and mn' (0.10), which represent evaporation, and plant biomass loss through density-independent mortality and





238 maintenance (Figure 4a). However, although the variation extent (2.06) of the linear term $w'\overline{n}^2 + 2\overline{nwn'}$, which represents linear interaction between water and plant in 239 arid and semi-arid region, caused by the LSV is greater than those of the linear terms 240 241 w' (0.15) and mn'(0.87) at first period, the linear interaction between water and plant $w'\overline{n}^2 + 2\overline{n}\overline{w}n'$ rapidly decays to 0.15 with the developing of time, which is 242 243 lower than effect of evaporation (0.21) and plant biomass loss (0.90) (Figure 4b). So, 244 the nonlinear character may be not reflected by the linear model and the LSV. The 245 high effect of the nonlinear terms brings to the loss of the water due to the CNOP-type 246 initial perturbation, and the remaining water does not supply the plant. The desert equilibrium state is generated. However, the low effect between water and plant of the 247 248 linear term also leads to the loss of the water due to the LSV-type initial perturbation, but the rest of water could support the plant. The grassland equilibrium state is kept. 249

For the desert state as the reference state (a=0.12), we find that the variation 250 extent (0.26) of the nonlinear term $(\overline{w} + w')(\overline{n} + n')^2 - \overline{wn}^2$ caused by the CNOP is 251 smaller than that of the linear term w' (0.33), and is greater mn'(0.19) at first year 252 253 (Figure 4c). However, after one year, the nonlinear effect is shown, and the variation 254 extent (0.26) is greater than those of the linear terms evaporation (0.04) and plant 255 biomass loss (0.21) (Figure 4c). The variation extent of the linear term of interaction 256 between water and plant is lower than those of the linear terms of evaporation and plant biomass loss all the time caused by the LSV (Figure 4d). For the desert as the 257 reference state (a=0.2), the effects of the nonlinear term (0.012) and two linear terms 258 259 (0.012 and 0.011) due to the CNOP are identical. As the change of time, the variation extent of the nonlinear term (0.011) is greater those of two linear terms (0.003) and 260 0.010) (Figure 4e). The variation extent of the linear term of interaction between 261 water and plant is always smaller than those of the linear terms of evaporation and 262 plant biomass loss caused by the LSV (Figure 4f). For the two desert equilibrium 263 states, the effect of the nonlinear term results in enough water to plant growing, and 264 the desert equilibrium states are finally transformed into the grassland and tree 265 equilibrium states. However, the effect of the linear term results in deficient water to 266





267 plant growing, and the desert equilibrium states are kept.

For the tree as the reference state (a=0.2), it is found that the patterns of the CNOP (*w*'=-0.001380, *n*'=-0.9399) and the LSV (*w*'=-0.0002412, *n*'=-0.9999) are similar. So, the variations of the tree due to the CNOP and the LSV are equivalent. The linear stable tree equilibrium is also nonlinear stable. In addition, the plant in this study was chosen as nonspatial plant in Eq. 3. In fact, this model could be employed to explore the stability of spatial plant to consider the advective term ($v \frac{\partial w}{\partial x}$) and diffusion term

274 $\left(\left(\frac{\partial^2}{\partial x} + \frac{\partial^2}{\partial y}\right)n\right)$ (Klausmeier, 1999; Sherratt and Lord, 2007; Sherratt, 2016) (Figure 5).

It is interesting that the linear stability and nonlinear stability of spatial plant arediscussed. In future, the issues will be answered.

Beyond all doubt, evapotranspiration (ET) is an important indicator factor for the 277 transition between two equilibrium states in arid and semi-arid regions (Kurc and 278 Small, 2004). The relationships of the ET-soil moisture impact on the plant stability. 279 Consistent with the previous work, the evaporation was also an important factor in our 280 281 studies (Huang et al., 2017). The evaporation caused the decreasing water in the soil layer, and the resulting would bring to a lack of supply for the plant. Hence, the 282 transition of from the grassland to the desert easily occurred. Compared to the effect 283 of the evaporation, the nonlinear interaction between the plant and the water was more 284 important from our findings using the CNOP approach. And, this effect will directly 285 286 result in the transition.

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288 5. Conclusions

In this study, three common methods (the Lyapunov method, the LSV method, and the CNOP method) are employed to explore the stabilities of plant (including grassland and tree) and desert. The first two methods are used to analyze the linear stabilities of plant and desert. The last method is applied to discuss the nonlinear stabilities of plant and desert. It is found that the linear stable grassland and desert equilibrium states are nonlinear stable when there is enough larger variation of initial





295	perturbation. Through computing the variations of nonlinear terms					
296	$(\overline{w} + w')(\overline{n} + n')^2 - \overline{w}\overline{n}^2$, it is demonstrated that the nonlinear interaction between					
297	water and plant plays an important role in the stabilities of grassland and desert					
298	compared to the linear terms of evaporation and plant biomass losses. The CNOP					
299	approach could reflect this nonlinear character, but the Lyapunov method and the LSV					
300	method fail.					
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324 Appendix

325	In the section, the dimensionless of the Eq. 1 is introduced (Klausmeier, 1999). In
326	the studies of Klausmeier (1999), the dimensionless processes have been stated. Here,
327	this treatment is introduced for readers' convenience. Let $w = R^{1/2}L^{-1/2}JW$, $n = R^{1/2}L^{-1/2}N$,
328	$x=L^{1/2}D^{-1/2}X$, $y=L^{1/2}D^{-1/2}Y$, $t=LT$, $a=AR^{1/2}L^{-3/2}J$, $m=ML^{-1}$, and $v=VL^{-1/2}D^{-1/2}$.
329	Klausmeier (1999) indicated that the rainfall (A) was about from 250 to 750 Kg H_2O
330	in arid and semi-arid region. The evaporation rate was $L=4$ year ⁻¹ . According to the
331	researches of Mauchamp et al. (1994), Klausmeier (1999) confined the four
332	parameters values: J_{tree} =0.002 kg dry mass (kg H ₂ O) ⁻¹ , J_{grass} =0.003 kg dry mass (kg
333	$H_2O)^{-1}$, $M_{tree}=0.18$ year ⁻¹ , and $M_{grass}=1.8$ year ⁻¹ . And, the $R_{tree}=1.5$ kg H_2O m ⁻² year ⁻¹
334	(kg dry mass) ⁻² and R_{grass} =100 kg H ₂ O m ⁻² year ⁻¹ (kg dry mass) ⁻² were also determined.
335	According to the above the parameters values, the dimensionless a and m could be
336	obtained for grass and tree as follows: $a_{\text{tree}}=0.077$ to 0.23, $m_{\text{tree}}=0.045$, $a_{\text{grass}}=0.94$ to
337	$2.81, m_{\text{grass}}=0.45.$
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354	Reference						
355							
356	Barclay, A., Gill, P. E., and Rosen, J. B.: SQP methods and their application to numerica						
357	optimal control, Variational Calculus, Optimal Control and Applications, Schmidt W. H., Heier K.,						
358	Bittner L., and Bulirsch R., Eds, Birkhauser, Basel, 207-222, 1998.						
359	Charney, J. G.: Dynamics of deserts and drought in the Sahel, Q. J. R. Meteorol. Soc., 101,						
360	192–202, 1975.						
361	Eklundh, L., and Olsson, L., Vegetation index trends for the African Sahel 1982–1999,						
362	Geophys. Res. Lett., 30(8), 1430, doi:10.1029/2002GL016772, 2003.						
363	Fu, C., and An, Z.: Study of aridization in northern China-A global change issue facing						
364	directly the demand of nation, Earth Science Frontiers, 9(2), 271-275, 2002.						
365	Huang, J., et al.: Dryland climate change: Recent progress and challenges, Rev. Geophys.,						
366	55, 719–778, doi:10.1002/2016RG000550, 2017.						
367	Klausmeier, C. A.: Regular and irregular patterns in semiarid vegetation, Science, 284,						
368	1826-1828, 1999.						
369	Kurc, S. A., and Small, E. E.: Dynamics of evapotranspiration in semiarid grassland and						
370	shrubland ecosystems during the summer monsoon season, central New Mexico, Water Resour.						
371	Res., 40, W09305, doi:10.1029/2004WR003068, 2004.						
372	Lu, J., and Ji, J.: A simulation and mechanism analysis of long-term variations at land surface						
373	over arid/semi-arid area in north China, J. Geophys. Res., 111, D09306,						
374	doi:10.1029/2005JD006252, 2006.						
375	Mauchamp A., Rambal, S., and Lepart, J.: Simulating the dynamics of a vegetation mosaic: a						
376	spatialized functional model, Ecological Modelling, 71(1-3): 107-130, 1994.						
377	Mitchell, S. W., and Csillag, F.: Assessing the stability and uncertainty of predicted						
378	vegetation growth under climatic variability: northern mixed grass prairie, Ecological Modelling,						
379	139(2-3), 101-121, 2001.						
380	Mu, M., Duan, W. S., and Wang, B.: Conditional nonlinear optimal perturbation and its						
381	applications, Nonlinear Processes in Geophysics, 10, 493-501, 2003.						
382	Mu, M. and Wang, B.: Nonlinear instability and sensitivity of a theoretical grassland						
383	ecosystem to finite-amplitude perturbations, Nonlin. Processes Geophys., 14, 409-423,						





- 384 doi:10.5194/npg-14-409-2007, 2007.
- 385 Mu M., Sun, L. and Dijkstra, H.A.: The sensitivity and stability of the ocean's thermohaline
- 386 circulation to finite amplitude perturbations, J. Physical Oceanography, 34, 2305-2315, 2004.
- 387 Ni, J.: Estimating grassland net primary productivity from field biomass measurements in
- temperate northern China, Plant Ecology, 174(2), 217-234, 2004.
- 389 Notaro, M., Liu, Z., and Williams, J. W.: Observed Vegetation-Climate Feedbacks in the
- 390 United States, J. Climate, 19, 763-786, 2006.
- 391 Piao, S. L., Fang, J. Y., Zhou, L., Tan, K., and Tao, S.: Changes in biomass carbon stocks in
- 392 China's grasslands between 1982 and 1999, Global Biogeochemical Cycles, 21, GB2002, doi:
- 393 10.1029/2005GB002634, 2007.
- Okin, G. S., D'Odorico, P., and Archer, S. R.: Impact of feedbacks on Chihuahuan desert
 grasslands: Transience and metastability, J. Geophys. Res., 114, G01004,
 doi:10.1029/2008JG000833, 2009.
- Sankey, J. B., Ravi, S., Wallace, C. S. A., Webb, R. H., and Huxman, T. E.: Quantifying soil
 surface change in degraded drylands: Shrub encroachment and effects of fire and vegetation
 removal in a desert grassland, J. Geophys. Res., 117, G02025, doi:10.1029/2012JG002002, 2012.
- Sherratt, J. A., and Lord, G. J.: Nonlinear dynamics and pattern bifurcations in a model for
 vegetation stripes in semi-arid environments, Theor. Popul. Biol., 71(1), 1-11, 2007.
- 402 Sherratt J.A.: When does colonisation of a semi-arid hillslope generate vegetation patterns? J.
 403 Math. Biol. 73, 199-226, 2016
- 404 Sun, G. D., and Mu, M.: Nonlinear feature of the abrupt transitions between multiple
 405 equilibria states of an ecosystem model, Adv. Atmos. Sci., 26, 293-304, doi:
 406 10.1007/s00376-009-0293-8, 2009.
- Sun, G. D., and Mu, M.: Nonlinearly combined impacts of initial perturbation from human
 activities and parameter perturbation from climate change on the grassland ecosystem, Nonlin.
 Processes Geophys., 18, 883-893, doi:10.5194/npg-18-883-2011, 2011.
- 410 Sun, G. D., and Mu, M.: The analyses of the net primary production due to regional and
- 411 seasonal temperature differences in eastern China using the LPJ model, Ecological Modelling, 289,
- 412 66–76, DOI: 10.1016/j.ecolmodel.2014.06.021, 2014
- 413 Sun, G. D., and Mu, M.: Understanding variations and seasonal characteristics of net primary

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- 414 production under two types of climate change scenarios in China using the LPJ model, Climatic
- 415 Change,120:755 769, DOI 11.1007/s10584-013-0833-1, 2013
- 416 Sun, G. D., and Xie, D. D.: A study of parameter uncertainties causing uncertainties in
- 417 modeling a grassland ecosystem using the conditional nonlinear optimal perturbation method,
- 418 Science China Earth Sciences, 60(9): 1674-1684, doi:10.1007/s11430-016-9065-9, 2017
- 419 Sun, G. D., Peng, F., and Mu, M.: Variations in soil moisture over the 'Huang-Huai-Hai Plain'
- 420 in China due to temperature change using the CNOP-P method and outputs from CMIP5, Science
- 421 China Earth Sciences, 60(10): 1838-1853 doi:10.1007/s11430-016-9061-3, 2017
- 422 Von Hardenberg, J., Meron, E., Shachak, M., and Zarmi, Y.: Diversity of vegetation patterns
- 423 and desertification, Phys. Rev. Lett., 87(19), 198101-1-4, 2001.
- 424 Xu, J., Ren, L. L., Ruan, X. H., Liu, X. F., and Yuan, F.: Development of a physically based
- 425 PDSI and its application for assessing the vegetation response to drought in northern China, J.
- 426 Geophys. Res., 117, D08106, doi:10.1029/2011JD016807, 2012.
- 427 Xue, Y. K., and Shukla, J.: The influence of land surface properties on Sahel climate. Part I:
- 428 desertification, J. Climate, 6, 2232-2245, 1993.
- 429 Xue, Y.: The Impact of Desertification in the Mongolian and the Inner Mongolian Grassland
- 430 on the Regional Climate, J. Climate, 9, 2173-2189, 1996.
- 431 Zeng, X. D., Shen, S. S. P., Zeng, X. B., and Dickinson, R. E.: Multiple equilibrium states
- 432 and the abrupt transitions in a dynamical system of soil water interacting with vegetation,
- 433 Geophys. Res. Lett., 31, 5501, doi: 10.1029/2003GL018910, 2004.
- 434 Zeng, X. D., Zeng, X. B., Shen, S. S. P., Dickinson, R. E., and Zeng, Q. C.: Vegetation-soil
- 435 water interaction within a dynamical ecosystem model of grassland in semi-arid areas, Tellus, 57B,
- 436 189-202, 2005.
- 437 Zeng, X. D., Wang, A. H., Zeng, Q. C., Dickinson, R. E., Zeng, X. B., and Shen, S. S. H.:
- 438 Intermediately complex models for the hydrological interactions in the atmosphere-vegetation-soil
- 439 system, Adv. Atmos. Sci., 23, 127-140, 2006.
- 440
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Figure 1. The equilibrium state of the vegetation for the theoretical model. Linearly stable equilibrium state is denoted by solid line corresponding to bare soil and vegetation. Linearly unstable equilibrium state is denoted by dash line corresponding to vegetation. (a): grass; (b) tree.

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514	Table 1 Th	Table 1 The linear stability analysis for different equilibrium states						
	Types	Equilibrium	Eigenvalues	Stability				
		w: water						
		<i>n</i> : plant biomass						
	Grassland	Grassland (<i>w</i> =0.203, <i>n</i> =2.21)	-0.344, -5.093	Linear stable				
	(<i>a</i> =1.2)	Grassland (w=0.997, n=0.451)	0.330, -1.084	Linear unstable				
		Desert (<i>w</i> =1.2, <i>n</i> =0)	-1.000, -0.450	Linear stable				
	Tree (<i>a</i> =0.2)	Tree (w=0.011, n=4.207)	-0.040, -18.613	Linear stable				
		Tree (w=0.189, n=0.238)	0.040, -1.052	Linear unstable				
		Desert (<i>w</i> =0.2, <i>n</i> =0)	-1.000, -0.045	Linear stable				
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534	Table 2 The	CNOP and I	LSV initial	perturbations	for	different	grassland	or	tree
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535 equilibrium state δ CNOP LSV Types Reference state (w' = -0.489,(w'=-0.741,Grassland Grassland (w=0.203, 1.0 (*a*=1.2) *n*=2.21) n' = -0.872)*n*'=-0.671) Desert (*w*=1.2, *n*=0) 0.53 (w'=0.332, (*w*'=0.365, n'=0.413)n'=0.384)Tree (w=0.01069, 1.0 Tree (*w*'=-0.001380, (w'=-0.0002412, (*a*=0.2) *n*=4.207) n' = -0.9399)n' = -0.9999)Desert (w=0.2, n=0) (w'=0.0125, (w'=0.0721, 0.24 *n*'=0.2397) *n*'=0.2290)

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