Nonlinear Processes
in Geophysics
Discussions

1 **Multistable Slip of a One-degree-of-freedom Spring-slider**

2 **Model in the Presence of Thermal-pressurized**

3 **Slip-weakening Friction and Viscosity**

4

5 Jeen-Hwa Wang

6 Institute of Earth Sciences, Academia Sinica

7 P.O. Box 1-55, Nangang, Taipei, TAIWAN

8 e-mail: jhwang@earth.sinica.edu.tw

9

10

11

12 **Abstract** This study is focused on multistable slip of earthquakes based on a

13 one-degree-of-freedom slider-slider model in the presence of thermal-pressurized

14 slip-weakening friction and viscosity by using the normalized equation of motion of

15 the model. The major model parameters are the normalized characteristic

16 displacement, U_c, of the friction law and the normalized viscosity coeficient, η,

17 between the slider and background plate. Analytic results at small slip suggest that

18 there is a solution regime for η and γ $(=1/U_c)$ to make the slider slip steadily.

19 Numerical simulations exhibit that the time variation in normalized velocity, V/V_{max}

20 (V_{max} is the maximum velocity), obviously depends on U_c and η. The effect on the

21 amplitude is stronger due to η than due to U_c. In the phase portrait of V/V_{max} versus

22 the normalized displacement, U/U_{max} (U_{max} is the maximum displacement), there are

23 two fixed points. The one at large V/V_{max} and large U/U_{max} is not an attractor; while

24 that at small V/V_{max} and small U/U_{max} can be an attractor for some values of η and U_c.

25 When $U_c<0.55$, unstable slip does not exist. When $U_c\geq0.55$, U_c and η divide the

26 solution domain into three regimes: stable, intermittent, and unstable (or chaotic)

27 regimes. For a certain U_c, the three regimes are controlled by a lower bound, η_l, and

28 an upper bound, η_u, of η. The values of η_l, η_u, and $\eta_u-\eta_l$ all decrease with increasing

29 U_c, thus suggesting that it is easier to yield unstable slip for larger U_c than for smaller

30 U_c or for larger η than for smaller η. When $U_c<1$, the Fourier spectra calculated from

31 simulation velocity waveforms exhibit several peaks, thus suggesting the existence of

32 nonlinear behavior of the system. When $U_c>1$, the related Fourier spectra show only

33   one peak, thus suggesting linear behavior of the system.

34

35   **Key Words**: Multistable slip, one-degree-of-freedom spring-slider model,

36   displacement, velocity, thermal-pressurized slip-weakening friction, viscosity

37

38   **1. Introduction**

39   The earthquake ruptures consist of three steps: nucleation, dynamical

40   propagation, and arrest. Due to the lack of a comprehensive model, a set of equations

41   to completely describe fault dynamics has not yet been established, because

42   earthquake ruptures are very complicated. Nevertheless, some models, for instance

43   the crack model and dynamical lattice model, have been developed to approach fault

44   dynamics. Several factors will control earthquake ruptures (see Wang, 2016b; and

45   cited references herein), including at least brittle-ductile fracture rheology, normal

46   stress, re-distribution of stresses after fracture, fault geometry, friction, seismic

47   coupling, pore fluid pressure, elastohydromechanic lubrication, thermal effect,

48   thermal pressurization, and metamorphic dehydration. A general review can be seen in

49   Bizzarri (2009). Among the factors, friction and viscosity are two important ones in

50   controlling faulting.

51   Burridge and Knopoff (1967) proposed a one-dimensional spring-slider model

52   (abbreviated as the 1-D BK model henceforth) to approach fault dynamics. Wang

53   (2000, 2012) extended this model to a two-dimensional version. The two models and

54   their modified versions have been long and widely applied to simulate the occurrences

55   of earthquakes (see Wang, 2008, 2012; and cited references therein). In the followings,

56   the one-, two-, three-, few-, and many-body models are used to represent the one-,

57   two-, three-, few-, and many-degree-of-freedom spring-slider models, respectively.

58   The few-body models have been long and widely used to approach faults (Turcotte,

59   1992)

60   Since the commonly-used friction laws are nonlinear, the dynamical model itself

61   could behave nonlinearly. A nonlinear dynamical system can exhibit chaotic

62   behaviour under some conditions (Thompson and Stewart, 1986; Turcotte, 1992).

63   This means that the system is highly sensitive to initial conditions (SIC) and thus a

64   small difference in initial conditions, including those caused by rounding errors in

65   numerical computation, yields widely diverging outcomes. This indicates that

66   long-term prediction is impossible in general, even though the system is deterministic,

67   meaning that its future behavior is fully determined by their initial conditions, without

68   random elements. This behavior is known as (deterministic) chaos (Lorenz, 1963).

69   An interesting question is: Can a simple few-body model with total symmetry

70   make significant predictions for fault behavior? Gu et al. (1984) first found some

71   chaotically bounded oscillations based on a one-body model with rate- and state-

72   dependent friction. Perez Pascual and Lomnitz-Adler (1988) studied the chaotic

73   motions of coupled relaxation oscillators. Related studies have been made based on

74   different spring-slider models: (1) a one-body model with rate- and state-dependent

75   friction (e.g., Gu et al., 1984; Belardinelli and Belardinelli, 1996; Ryabov and Ito,

76   2001; Erickson et al., 2008, 2011; Kostić et al., 2013); (2) a one-body model with

77   velocity-weakening friction (e.g., Brun and Gomez, 1994); (3) a one-body model with

78   slip-weakening friction (e.g., Wang, 2016a,b); (4) a two-slider model with simple

79   static/dynamic friction (e.g., Nussbaum and Ruina, 1987; Huang and Turcotte, 1990);

80   (5) a two-body model with velocity-dependent friction (e.g., Huang and Turcotte,

81   1992; de Sousa Vieira, 1999; Galvanetto, 2002); (6) a two-body model with rate- and

82   state-dependent friction (e.g., Abe and Kato, 2013); (7) a two-body model with

83   velocity-weakening friction (Brun and Gomez, 1994); (8) a two-body model with

84   slip-weakening friction (e.g., Wang, 2017); (9) many-body model with velocity-

85   weakening friction (e.g., Carlson and Langer, 1989; Wang, 1995, 1996); and (10)

86   one-body quasi-static model with rate- and state-dependent friction (e.g., Shkoller and

87   Minster, 1997). Results suggest that predictions for fault behaviour are questionable

88   due to the possible presence of chaotic slip.

89   The frictional effect on earthquake ruptures has been widely studied as

90   mentioned above. However, the studies of viscous effect on earthquake ruptures are

91   rare. The viscous effect mentioned in Rice et al. (2001) was just an implicit factor

92   which is included in the evolution effect of friction law. In this work, I will investigate

93   the effects of thermal pressurized slip-weakening friction and viscosity on earthquake

94   ruptures and the generation of unstable (or chaotic) slip based on a one-body model.

95

96   **2. MODEL**

97   **2.1 One-body Model**

98   Fig. 1 shows the one-body model whose equation of motion is:

99

100       $md^2u/dt^2 = -K(u-u_o) - F(u,v) - \Phi(v),$ (1)

101

102 where m is the mass of the slider, u and v (=du/dt) are, respectively, the displacement

103 and velocity of the slider, u_o is the equilibrium location of the slider, K is the spring

104 constant, F is the frictional force between the slider and the background and a

105 function of u or v, and Φ is the viscous force between the slider and the background

106 and a function of v. The slider is pulled by a driving force F_D due to the moving plate

107 with a constant driving velocity, v_p, through a leaf spring of strength, K. Hence, the

108 driving force is $F_D=Kv_pt$ and thus $u_o=v_pt$. When F_D is slightly larger than the static

109 frictional force, F_o, friction changes from static friction strength to dynamic one and

110 thus the slider moves.

111 **2.2 Viscosity**

112   Jeffreys (1942) first emphasized the importance of viscosity on faulting.

113 Frictional melts in faults depend on temperature, pressure, water content, and etc.

114 (Turcotte and Schubert, 1982) and can yield viscosity on the fault plane (Byerlee,

115 1968). Rice et al. (2001) discussed that rate- and state-dependent friction in thermally

116 activated processes allows creep slippage at asperity contacts on the fault plane.

117 Scholz (1990) suggested that the friction melts would present significant viscous

118 resistance to shear and thus inhibit continued slip. However, Spray (1993, 1995, 2005)

119 stressed that the frictional melts possessing low viscosity could generate a sufficient

120 melt volume to reduce the effective normal stress and thus act as fault lubricants

121 during co-seismic slip. His results show that viscosity remarkably decreases with

122 increasing temperature. For example, Wang (2011) assumed that quartz plasticity

123 could be formed in the fault zone when T>300 ^{o}C after faulting and it would lubricate

124 the fault plane at higher T and yield viscous stresses to resist slip at lower T. From

125 numerical simulations, Wang (2007, 2016b, 2017) stressed the viscous effect on

126 faulting. Noted that several researchers (Knopoff et al., 1973; Cohen, 1979; Xu and

127 Knopoff, 1994; Knopoff and Ni, 2001; Dragoni and Santini, 2015) took viscosity as a

128 factor in causing seismic radiation to reduce energy during faulting.

129   The viscosity coefficient, υ, of rocks is mainly controlled by temperature, T. An

130 increase in T will yield partial melting of rocks and thus the viscosity coefficient, υ,

131 first is increased, then reaches the largest value at a particular T, and finally decreases

132 with increasing T The relation between υ and T can be described by the following

133 equation (e.g., Turcotte and Schubert, 1982): $\upsilon=\upsilon_o\exp[(E_o+pV_a/RT)]$ where υ_o is the

Nonlinear Processes
in Geophysics
Discussions

Open Access

EGU

134   largest viscosity at low ambient T of an area, E_o is the activation energy per mole, p is

135   the pressure, V_a is the activation volume per mole, and R is the universal gas constant

136   ($E_o/R \approx 3 \times 10^4$ K). Obviously, υ decreases with increasing T. This is particularly

137   remarkable in regions of high confining pressure. On the other hand, Diniega et al.

138   (2013) assume that υ exponentially depends on temperature: $\upsilon \sim e^{\beta(1-T^*)}$, where β is a

139   constant and $T^* = (T-T_C)/(T_H-T_C)$ is a dimensionless temperature within a

140   temperature range of T_C to T_H. The value of υ increases with T^* when $T^* < 1$ and

141   decreases with increasing T^* when $T^* > 1$. Wang (2011) inferred that in the major slip

142   zone<0.01 m of the 1999 M_s7.6 Chi-Chi, Taiwan, earthquake, T(t) in the fault zone at

143   a depth of 1111 m increased from ambient temperature $T_a \approx 45$ ^{o}C at t=0 s to peak

144   temperature T_{peak}=1135.1 ^{o}C at t=~2.5 s. T(t) began to decrease after t=2.5 s and

145   dropped to 160 ^{o}C at t=195 s. This yields a change of viscosity in the fault zone.

146   The description about the physical models of viscosity can be found in several

147   articles (Jaeger and Cook, 1977; Cohen, 1979; Hudson, 1980; Wang, 2016b). A brief

148   description is given below. For many deformed materials, there are elastic and viscous

149   components. The viscous component can be modeled as a dashpot such that the

150   stress–strain rate relationship is: $\sigma = \upsilon(d\varepsilon/dt)$ where σ and ε are the stress and the strain,

151   respectively. Two simple models (shown in Fig. 2) commonly used to describe the

152   viscous materials are the Maxwell model and the Kelvin-Voigt model (or the Voigt

153   model). The first one can be represented by a purely viscous damper (denoted by "D")

154   and a purely elastic spring (denoted by "S") connected in series,. Its constitution

155   equation is: $d\varepsilon/dt = d\varepsilon_D/dt + d\varepsilon_S/dt = \sigma/\upsilon + E^{-1}d\sigma/dt$ where E is the elastic modulus and

156    $\sigma = E\varepsilon$. The constitutive relation of the second model is: $\sigma(t) = E\varepsilon(t) + \upsilon d\varepsilon(t)/dt$.

157   For the Maxwell model, the strain will increase, without a upper limit, with time;

158   while the Kelvin-Voigt model the strain will increases, with a upper limit, with time.

159   Wang (2016b) assumed that the latter is more appropriate than the former to be

160   applied to the seismological problems as suggested by Hudson (1980). Hence, the

161   Kelvin-Voigt model is taken in this study. To simplify the problem, only a constant

162   viscosity is considered below. The viscous stress at the slider is represented by $-\upsilon v$.

163   However, it is not easy to directly implement viscosity in a dynamical system as

164   used in this study. Wang (2016b) represented the viscosity coefficient in an alternative

165   way. Viscosity leads to the damping of oscillations of a body in viscous fluids. The

166   damping coefficient, η, depends on the viscosity coefficient, υ, and the linear

167   dimension, R, of the body in a viscous fluid. According to Stokes' law, the η of a

168   sphere of radius R in a viscous fluid of υ is $\eta=6\pi R\upsilon$ (cf. Kittel et al., 1968). In order

169   to simplify the problem, the damping coefficient is taken in this study. Hence, the

170   viscous force is $\Phi=\eta v$. Noted that the unit of η is $N(m/s)^{-1}$.

171   2.3 Friction caused by thermal pressurization

172   Numerous factors can affect friction (see Wang, 2009, 2016b; and cited

173   references herein). When fluids are present and temperature changes in faults, thermal

174   pressurization will yield resistance on the fault plane and thus play a significant role

175   on earthquake rupture (Sibson, 1973; Lachenbruch, 1980; Chester and Higgs, 1992;

176   Fialko, 2004; Fialko and Khzan, 2005; Bizzari and Cocco, 2006a,b; Rice, 2006; Wang,

177   2000, 2006, 2009, 2011, 2013, 2016b, 017; Bizzarri, 2010; Bizzarri, 2011a,b).

178   Rice (2006) proposed two end-members models for thermal pressurization: the

179   adiabatic-undrained-deformation (AUD) model and slip-on-a-plane (SOP) model. He

180   also obtained the shear stress-slip functions caused by the two models. The first model

181   corresponds to a homogeneous simple shear strain ε at a constant normal stress σ_n on

182   a spatial scale of the sheared layer that is broad enough to effectively preclude heat or

183   fluid transfer. The second model shows that all sliding is on the plane with $\tau(0)=$

184    $f(\sigma_n-p_o)$ where p_o is the pore fluid pressure on the sliding plane (y=0). For this second

185   model, heat is transferred outwards from the fault plane. Although the stress $\tau_{sop}(u)$

186   also shows slip-weakening (Wang, 2009), the SOP model is not appropriate in this

187   study because of the request of a constant velocity for this model.

188   The shear stress-slip functions, $\tau(u)$, caused by the AUD model is:

189

190        $\tau_{aud}(u)= f(\sigma_n-p_o)\exp(-u/u_c).$ (3)

191

192   The parameters u_c is the characteristic displacements associated with the thickness

193   and some physical properties of fault zone. The stress $\tau_{aud}(u)$ displays exponentially

194   with u and thus exhibits slip-weakening friction. Based on the AUD model, Wang

195   (2009) proposed a simplified slip-weakening friction law (denoted by the TP law

196   hereafter): $F(u)=F_o\exp(-u/u_c)$, where F_o is the static frictional force, to study seismic

197   efficiency. Wang (2016b, 2017) applied the law to simulate slip of one-body and

198   two-body spring-slider models. Fig. 3 exhibits F(u) versus u for five values of u_c, i.e.,

199   0.1, 0.3, 0.5, 0.7, and 0.9 m. The friction force decreases with increasing u and it

Nonlinear Processes
in Geophysics

Open Access

Discussions

EGU

200   decreases faster for smaller u_c than for larger u_c. Meanwhile, the force drop decreases

201   with increasing u_c. For small u, exp($-u/u_c$) can be approximated by $1-u/u_c$ (Wang,

202   2016a,b, 2017). The parameter u_c^{-1} is almost the decreasing rate, γ, of friction force

203   with slip at small u. Small (large) u_c is related to large (small) γ.

204   **2.4. Predominant Frequency and Period of the System**

205   To conduct marginal analyses of slip of one-body model with friction, Wang

206   (2016b) used the friction law: $F(u)=F_o-\gamma u$. His results show that the natural periods

207   are $T_o=2\pi/(K/m)^{1/2}$ when friction and viscosity are excluded and

208

209            $T_n=T_o/[1-T_o^2(\eta^2+4m\gamma)/(4\pi m)^2]^{1/2}$.   (4)

210

211   when friction and viscosity are included. Clearly, T_n is longer than T_o. Eq. (4) shows

212   that T_n increases with η and γ, thus indicating that friction and viscosity both lengthen

213   the natural period of the system.

214

215   **3. Normalization of Equation of Motion**

216   Substituting the TP law and the linear viscous law into Eq. (1) leads to

217

218            $md^2u/dt^2=-K(u-u_o)-F_o exp(-u/u_c)-\eta v$.   (5)

219

220   To simplify numerical computations, Eq. (5) is normalized based on the following

221   normalization parameters: $D_o=F_o/K$, $\omega_o=(K/m)^{1/2}$, $\tau=\omega_o t$, $U=u/D_o$, $U_c=u_c/D_o$, and

222    $\Gamma_D=F_D/K$. This gives $du/dt=[F_o/(mK)^{1/2}] dU/d\tau$, $d^2u/dt^2=(F_o/mK)d^2U/d\tau^2$. The driving

223   velocity becomes $V_p=v_p/D_o\omega_o$ Hence, the normalized acceleration and velocity are,

224   respectively, $A=d^2U/d\tau^2$ and $V=dU/d\tau$. The phase ωt is replaced by $\Omega\tau$, where

225    $\Omega=\omega/\omega_o$ is the dimensionless angular frequency. Note that $\eta/(mK)^{1/2}$ is simply

226   denoted by η below. Clearly, all normalization parameters are dimensionless. Hence,

227   Eq. (5) becomes:

228

229            $d^2U/d\tau^2=-U-\eta dU/d\tau-exp(-U/U_c)+\Gamma_D$.   (6)

230

231   When $F_D=v_p t$ or $\Gamma_D=V_p\tau$, Eq. (6) is transformed to a set of three first-order

Nonlinear Processes
in Geophysics
Discussions

Open Access

EGU

232   differential equations by defining $x=U/U_c$, $y=V/V_p$, and $z=-U+V_p\tau-\eta V_p y_\tau$

233   ($y_t=dy/d\tau$):

234

235         $x_\tau=(V_p/U_c)y$ (7a)

236

237         $y_\tau=(z-e^{-x})/V_p,$ (7b)

238

239         $z_\tau=V_p(1-y-\eta y_\tau).$ (7c)

240

241   As $x\ll1$, $e^{-x}\approx1-x$ and thus Eq. (7b) can be approximated by $y_\tau\approx(z-1+x)/V_p$. The

242   condition of $x\ll1$ shows $U/U_c\ll1$. Differential of this equation leas to

243    $y_{\tau\tau}\approx(z_\tau+x_\tau)/V_p$, where $y_{\tau\tau}=d^2y/d\tau^2$. Substituting Eqs. (7a) and (7b) into this equation

244   gives

245

246         $y_{\tau\tau}+\eta y_\tau+(1-1/U_c)y=1.$ (8)

247

248   The homogeneous equation of Eq. (8) is

249

250         $y_{\tau\tau}+\eta y_\tau+(1-1/U_c)y=0.$ (9)

251

252   Let the general solution be $y\sim e^{\lambda\tau}$. This leads to $[\lambda^2+\eta\lambda+(1-/U_c)]y=0$ or

253

254         $\lambda^2+\eta\lambda+(1-/U_c)=0.$ (10)

255

256   The solutions of Eq. (10) are

257

258         $\lambda_\pm=-\eta/2\pm[\eta^2-4(1-1/U_c)]^{1/2}/2.$ (11)

259

260   The term $-\eta/2$ of Eq. (11) leads to $e^{-\lambda/2}$ which yields attenuation of y. Define $D(\eta,1/U_c)$

261   to be $\eta^2-4(1-1/U_c)$. As mentioned above, U_c^{-1} is the normalized decreasing rate of

262   friction, Γ, at $U=0$. Fig. 4 shows the plot of η versus $1/U_c$ and thus exhibits the root

263   structure of the system. Because $\eta>0$ and $U_c>0$, only the plot in the first quadrant is

264 present in Fig. 4. The solid line displays the function: $D(\eta, 1/U_c) = \eta^2 - 4(1-1/U_c) = 0$.

265 Along the line, we have $\eta^2 = 4(1-1/U_c)$, and thus $\lambda_\pm = -\eta/2$. In other word, the roots are

266 equal and real, and thus the solution is a stable inflected node displayed by a solid

267 circle in Fig. 4. As $D(\eta, 1/U_c) > 0$ or $\eta^2 > 4(1-1/U_c)$, the roots are both real and negative.

268 The solution shows no oscillation and thus is a stable node shown by a solid square in

269 Fig. 4. As $D(\eta, 1/U_c) < 0$ or $\eta^2 < 4(1-1/U_c)$, the roots are complex with negative real part.

270 This results in oscillations of exponentially decaying amplitude. The solution is a

271 stable spiral or a stable focus shown by an open circle in Fig. 4.

272

273 **4. Numerical Simulations**

274 Let $y_1 = U$ and thus $y_2 = dU/d\tau$. Eq. (6) can be re-written as two first-order

275 differential equations:

276

277 $\qquad dy_1/d\tau = y_2$ (12a)

278

279 $\qquad dy_2/d\tau = -y_1 - \eta y_2 - \exp(-y_1/U_c) + \Gamma_D$. (12b)

280

281 Eq. (12) will be numerically solved using the fourth-order Runge-Kutta method (Press

282 et al., 1986). To simplify the following computations, the value of Γ_D is set to be a

283 small constant of 10^{-5}, which can continuously enforce the slider to move.

284 A phase portrait, denoted by $y = f(x)$, is a plot of a physical quantity versus

285 another of an object in a dynamical system (Thompson and Stewart, 1986). The

286 intersection point of the bisection line, i.e., $y = x$, and $f(x)$ is called the fixed point, that

287 is, $f(x) = x$. If the function $f(x)$ is continuously differentiable in an open domain near a

288 fixed point x_f and $|f'(x_f)| < 1$, attraction is generated. In other words, an attractive fixed

289 point is a fixed point x_f of a function $f(x)$ such that for any value of x in the domain

290 that is close enough to x_f, the iterated function sequences, i.e., x, $f(x)$, $f^2(x)$, $f^3(x)$, ...,

291 converges to x_f. An attractive fixed point is a special case of a wider mathematical

292 concept of attractors. Chaos can be generated at some attractors. The details can be

293 seen in Thompson and Stewart (1986) or other nonlinear literatures. In the following

294 plots, the intersection points of the bisection line (denoted by a thin solid line) with

295 the phase portrait of V/V_{max} versus U/U_{max} are the fixed points. To explore nonlinear

296 behavior of a system, the Fourier spectrum $F[V(\Omega_k)]$, where $\Omega_k = k/\delta\tau$ is the

Nonlinear Processes
in Geophysics
Discussions

Open Access

EGU

297   dimensionless angular frequency at k=0, ..., N-1, is calculated for the simulation

298   velocity waveform through the fast Fourier transform (Press et al., 1986). The

299   bifurcation from a predominant period to others will be seen in the Fourier spectra.

300   Numerical simulations for the time variation in V/V_{max}, the phase portrait of

301   V/V_{max} versus U/U_{max}, and Fourier spectrum based on different values of model

302   parameters are displayed in Figs. 5–12. In the figures, V_{max} and U_{max} are, respectively,

303   the maximum velocity and displacement for case (a) of each figure, because the

304   maximum values of U and V decrease from case (a) to case (d) in this study.

305   First, the cases excluding viscosity, i.e., $\eta=0$, are explored. Fig. 5 is numerically

306   made for four values of U_c: (a) for $U_c=0.1$; (b) for $U_c=0.4$; (c) for $U_c=0.7$; and (d) for

307    $U_c=0.9$ when $\eta=0$. Fig. 6 is numerically made for four values of U_c: (a) for $U_c=1.00$;

308   (b) for $U_c=1.01$; (c) for $U_c=1.15$; and (d) for $U_c=2.00$ when $\eta=0$. A comparison

309   between Fig. 5 and Fig. 6 suggests that $U_c=1$ is a transition value of the friction law

310   between two modes of slip as displayed in Fig. 4. Only $U_c<1$ is considered below.

311   Secondly, the cases including both friction and viscosity are studied. Fig. 7 is

312   numerically made for four values of η: (a) for $\eta=0.20$; (b) for $\eta=0.50$; (c) for $\eta=0.87$;

313   and (d) for $\eta=0.90$ when $U_c=0.20$. Obviously, the time variation in V/V_{max} exhibits

314   cyclic oscillations with a particular period when $\eta<\eta_1=0.86$ and has intermittent slip

315   with shorter periods when $\eta>\eta_1$. Such a phenomenon holds also for $\eta<5.5$.

316   Fig. 8 is numerically made for four values of η: (a) for $\eta=0.46$; (b) for $\eta=0.47$; (c)

317   for $\eta=0.98$; and (d) for $\eta=0.99$ when $U_c=0.55$. The Fourier spectrum is not calculated

318   for case (d), because the velocity becomes negative infinity at a certain time. The time

319   variation in V/V_{max} exhibits cyclic oscillations specified with three main frequencies

320   when $\eta<\eta_1=0.47$. There is intermittency slip with shorter periods when

321    $\eta_1<\eta<\eta_u=0.98$. There are unstable slip when $\eta>\eta_u$. This phenomenon holds also

322   when $0.55<U_c<1$.

323   Four examples for η varying from $\eta<\eta_u$ to $\eta>\eta_u$ for different values of U_c are

324   displayed in Figs. 9–12. Fig. 9 is made for four values of η: (a) for $\eta=0.39$; (b) for

325    $\eta=0.83$; (c) for $\eta=0.84$; and (d) for $\eta=0.85$ when $U_c=0.6$. Fig. 10 is made for four

326   values of η: (a) for $\eta=0.34$; (b) for $\eta=0.71$; (c) for $\eta=0.72$; and (d) for $\eta=0.73$ when

327    $U_c=0.7$. Fig. 11 is made for four values of η: (a) for $\eta=0.25$; (b) for $\eta=0.53$; (c) for

328    $\eta=0.54$; and (d) for $\eta=0.55$ when $U_c=0.8$. Fig. 12 is made for four values of η: (a) for

329   η=0.14; (b) for η=0.35; (c) for η=0.36; and (d) for η=0.37 when U_c=0.9. The Fourier

330   spectrum is not calculated for case (d) in each example, because the velocity becomes

331   negative infinity at a certain time.

332   Fig. 13 exhibits the data points of η_l (with a solid square) and that of η_u (with a

333   solid circle) for several values U_c. The values of η_l and η_u for several values of U_c

334   are given in Table 1. The figure exhibits a stable regime when $\eta \leq \eta_l$, an intermittency

335   or transition regime when $\eta_l < \eta \leq \eta_u$, and unstable regime when $\eta > \eta_u$.

336

337   **5. Discussion**

338   As mentioned above, the natural period of the one-body system at low

339   displacements is $T_o = 2\pi/\omega_o = 2\pi(m/K)^{1/2}$ in the absence of friction and viscosity and

340    $T_n = 2\pi/\omega_n = T_o/[1 - T_o^2(\eta^2 + 4m\gamma)/(4\pi m)^2]^{1/2}$ in the presence of friction and viscosity.

341   Due to $\gamma = 1/u_c$ at u=0, T_n increases with decreasing u_c. Obviously, T_n is longer than

342   T_o and increases with η and γ, thus indicating that friction and viscosity both lengthen

343   the natural period of the system.

344   Based on the marginal analysis of the normalized equation of motion, i.e., Eq.

345   (11), the plot of η versus $1/U_c$ is displayed in Fig. 4 which exhibits the phase portrait

346   and root structure of the system. Since η and U_c are both positive, only the plot of η

347   versu $1/U_c$ in the first quadrant is displayed. In Fig. 4, the solid line displays the

348   function: $D(\eta,1/U_c)=\eta^2-4(1-1/U_c)=0$. Along the line, the solution $\lambda_\pm=-\eta/2$ and thus

349    $\exp(\lambda t)=\exp(-\eta/2)$. In other word, the roots are equal and real, and, thus, the phase

350   portrait is a stable inflected node displayed by a solid circle in Fig. 4. Because of $\eta \geq 0$,

351   we have $1/U_c \leq 1$. As $D(\eta,1/U_c)>0$ or $\eta^2>4(1-1/U_c)$, the roots are both real and

352   negative. The solution shows no oscillation and thus phase portrait is a stable node

353   shown by a solid square in Fig. 4. Because of $\eta \geq 0$, we have $1/U_c \leq 1$. As $D(\eta,1/U_c)<0$

354   or $\eta^2<4(1-1/U_c)$, the roots are complex with a negative real part. This results in

355   oscillations with exponentially decaying amplitude. The phase portrait is a stable

356   spiral or a stable focus shown by an open circle in Fig. 4.

357   Fig. 5 exhibits the time variation in V/V_{max}, the phase portrait of V/V_{max} versus

358   U/U_{max}, and Fourier spectrum for four values of U_c: (a) for U_c=0.1; (b) for U_c=0.4; (c)

359   for U_c=0.7; and (d) for U_c=0.9 when η=0. In the first panels, the time variation in

360   V/V_{max} exhibits cyclic behavior and the amplitude of V/V_{max} decreases and the

361   predominant period of signal increases with increasing U_c. This is consistent with Eq.

362   (5) in which T_n increases with U_c. Although the four phase portraits are almost similar,

363   yet their size decreases with increasing U_c. The second panels exhibit two fixed points:

364   one at V=0 and U=0 and the second one at larger V and larger V. The slope values at

365   the first fixed points decrease with increasing U_c, thus suggesting that the fixed point

366   is not an attractor for small U_c and can be an attractor for larger U_c. The slope values

367   at the fixed points for smaller U_c are greater than 1 and thus they cannot be an

368   attractor. The third panel for each case displays the Fourier spectrum. Fourier spectra

369   exhibit that in addition to the peak related to the predominant frequency, there are

370   numerous peaks associated with higher frequencies. This shows nonlinear behavior

371   caused by nonlinear friction. The frequency related to the first peak decreases with

372   increasing U_c. The amplitude of a peak decreases with increasing U_c. The amplitude

373   of a peak decreases with increasing Ω for small U_c; while it first increases up to the

374   maximum and then decreases with increasing Ω for large U_c. The amplitude of a peak

375   becomes very small when $\Omega > 0.25$.

376   Fig. 6 exhibits the time variation in V/V_{max}, the phase portrait of V/V_{max} versus

377   U/U_{max}, and Fourier spectrum for four values of U_c: (a) for U_c=1.00; (b) for U_c=1.01;

378   (c) for U_c=1.15; and (d) for U_c=2.0 when η=0. In the first panels, the time variation in

379   V/V_{max} exhibits cyclic behavior and the amplitude of V/V_{max} remarkably decreases

380   with increasing U_c when $U_c > 1$. In the second panels, the size of phase portrait

381   decreases with increasing U_c and there are two fixed points: the first one at V=0 and

382   U=0 and the second one at larger V and larger V. With comparison to the phase

383   portrait of U_c=1.0, the phase portrait becomes very small when $U_c \geq 1.15$. In contrast

384   to Fig. 5, the absolute values of slope at the fixed points in Fig. 6 increase with U_c.

385   Hence, the fixed points cannot be an attractor for $U_c \geq 1$. In the third panels, Fourier

386   spectra exhibit that except for U_c=1, there is only one peak and the predominant

387   frequency increases or the predominant period decreases with increasing U_c. This is

388   consistent with Eq. (5). Results show that nonlinear behavior disappears when $U_c > 1$.

389   In addition, the amplitude of a peak decreases with increasing U_c when $U_c > 1$.

390   Obviously, U_c=1 is the critical value of the friction law as displayed in Fig. 4.

391   Fig. 7 exhibits the time variation in V/V_{max}, the phase portrait of V/V_{max} versus

392   U/U_{max}, and Fourier spectrum for four values of η: (a) for η=0.20; (b) for η=0.50; (c)

393   for η=0.87; and (d) for η=0.90 when U_c=0.20. In the first panels, the time variation in

394   V/V_{max} exhibits cyclic behavior and the amplitude of V/V_{max} decreases with

395   increasing η. The predominant period of signal only slightly increases with increasing

396   η, because η changes in a small range. In the second panels, the size of phase portrait

397   decreases with increasing U_c and there are two fixed points: the first one at $V=0$ and

398    $U=0$ and the second one at larger V and larger V. Since the slope values of fixed

399   points are clearly all higher than 1, they are not an attractor. In the third panels, the

400   Fourier spectra exhibit that in addition to the peak related to the predominant

401   frequency, there are numerous peaks associated with higher Ω. This shows nonlinear

402   behavior, mainly caused by nonlinear friction, of the model. The highest peak for case

403   (a) appears at the second frequency. When $\eta<0.9$, the amplitude of a peak decreases

404   with increasing η. The frequencies related to the peaks do not change remarkably,

405   because η varies in a small range. Except for case (a), the amplitude of a peak

406   decreases with increasing Ω. The third peak amplitude disappears when $\eta>0.5$. The

407   amplitude of a peak becomes very small when $\Omega>0.25$. Except for $U_c=0.1$, the

408   frequencies related to the peaks in Fig. 7 are different from and slightly smaller than

409   those in Fig. 5. Note that when $U_c<0.55$ the simulation results in Fig. 5 are similar to

410   those in Fig. 6.

411   Fig. 8 shows the time variation in V/V_{max}, the phase portrait of V/V_{max} versus

412   U/U_{max}, and Fourier spectrum for four values of η: (a) for $\eta=0.46$; (b) for $\eta=0.47$; (c)

413   for $\eta=0.98$; and (d) for $\eta=0.99$ when $U_c=0.55$. When $\eta\leq0.47$, the time variation in

414   V/V_{max} exhibits cyclic oscillations specified with different main angular frequencies.

415   When $\eta>0.47$ (for example $\eta=0.98$ in the figure), in addition to cyclic behavior there

416   is small intermittent slip with shorter periods. This phenomenon also exists when

417    $\eta_l<\eta<\eta_u=0.98$. There are unstable (or chaotic) slip when $\eta>\eta_u$. Hence, the phase

418   portraits in the second panels display unstable slip at small V and U when

419    $\eta_l<\eta\leq\eta_u=0.98$. When $\eta=0.99$, the velocity becomes negative infinity at a certain time

420   and the phase portrait also displays unstable or chaotic slip at small V and U. Since

421   the slope values of fixed points at large V and U are clearly higher than 1, they are not

422   an attractor. Due to the appearance of infinity velocity when $\eta=0.99$, the Fourier

423   spectrum is not calculated for $\eta=0.99$. The Fourier spectra exhibit that when $\eta<0.47$,

424   in addition to the peak related to the predominant frequency, there are numerous peaks

425   associated with higher Ω. This shows nonlinear behavior of the model caused by

426   nonlinear friction. The amplitude of a peak decreases with increasing U_c and the peak

Nonlinear Processes
in Geophysics
Discussions

427   amplitude decreases with increasing Ω. When $\eta=0.98$, the amplitude of the highest

428   peak is much larger than others. For the first three cases, the amplitude of a peak

429   becomes very small when $\Omega>0.25$. The frequencies related to the peaks in Fig. 8 are

430   different from and slightly smaller than those in Fig. 7.

431   Figs. 9–12 show a variation from stable slip to intermittent slip and then to

432   unstable or chaotic slip when η increases from a smaller value to a larger one for

433    $U_c=0.6$, 0.7, 0.8, and 0.9. The values of η_u for $U_c=0.20$–0.95 with a unit difference of

434   0.05 are given in Table 1. Like Fig. 8, when $\eta\leq\eta_l$, the time variation in V/V_{max}

435   exhibits only cyclic oscillations specified with different frequencies. When $\eta_l<\eta\leq\eta_u$,

436   there are small intermittent displacements appear in the cyclic oscillations. Hence, the

437   phase portraits display that unstable slip at small V and U when $\eta_l<\eta\leq\eta_u$. When

438    $\eta>\eta_u$, the velocity becomes negative infinity at a certain time and the phase portrait

439   displays unstable slip at small V and U. Due to the appearance of infinity velocity, the

440   Fourier spectrum is not calculated for $\eta>\eta_u$. When $\eta<\eta_l$, in addition to the peak

441   related to the predominant frequency, there are numerous peaks related to higher Ω.

442   This shows nonlinear behavior, mainly caused by nonlinear friction, of the model. The

443   amplitude of a peak decreases with increasing U_c and the amplitude of a peak

444   decreases with increasing Ω. For the first three cases, the amplitude of a peak

445   becomes very small when $\Omega>0.25$. Figs. 7–12 show that the frequencies related to the

446   peaks slightly decrease with increasing U_c and the decreasing rate decreases with

447   increasing U_c. In other word, the frequencies related to the peaks for large U_c are

448   almost similar. The number of higher peaks and the amplitudes of peaks at higher Ω

449   both decrease with increasing η. This indicates that viscosity makes a stronger effect

450   on higher- frequency waves than on lower ones, and the effect increases with η.

451   Fig. 13 exhibits the data points of η_l (with a solid square) and that of η_u (with a

452   solid circle) for several values U_c. The values of η_l and η_u for several values of U_c

453   are given in Table 1. The figure exhibits a stable regime when $\eta\leq\eta_l$, an intermittency

454   (or transition) regime when $\eta_l<\eta\leq\eta_u$, and unstable (or chaotic) regime when $\eta>\eta_u$.

455   When $U_c<0.55$, there is no η_l, in other word, unstable slip does not exist. Clearly, η_l,

456   η_u, and their difference $\eta_u-\eta_l$ all decrease with increasing U_c. This means that it is

457   easier to yield unstable slip for larger U_c than for smaller U_c. Since smaller U_c is

458   associated to larger γ of decreasing rate of friction force with slip, it is easier to yield

459   unstable slip from smaller γ than from larger γ.

460   Huang and Turcotte (1990, 1992) observed intermittent phases in the

461   displacements based on a two-body model. In other word, some major events are

462   proceeded by numerous small events. Those small events could be foreshocks. They

463   also claimed that earthquakes are an example of deterministic chaos. Ryabov and Ito

464   (2001) also found intermittent phase transitions in a two-dimensional one-body model

465   with velocity-weakening friction. Their simulations exhibit that intermittent phases

466   appear before large ruptures. From numerical simulations of earthquake ruptures

467   using a one-body model with a rate- and state-friction law, Erickson et al. (2008)

468   found that the system undergoes a Hopf bifurcation to a periodic orbit. This periodic

469   orbit then undergoes a period doubling cascade into a strange attractor, recognized as

470   broadband noise in the power spectrum. From numerical simulations of earthquake

471   ruptures using a two-body model with a rate- and state-friction law, Abe and Kato

472   (2013) observed various slip patterns, including the periodic recurrence of seismic and

473   aseismic slip events, and several types of chaotic behavior. The system exhibits

474   typical period-doubling sequences for some parameter ranges, and attains chaotic

475   motion. Their results also suggest that the simulated slip behavior is deterministic

476   chaos andt ime variations of cumulative slip in chaotic slip patterns can be well

477   approximated by a time-predictable model. In some cases, both seismic and aseismic

478   slip events occur at a slider, and aseismic slip events complicate the earthquake

479   recurrence patterns. The present results seem to be comparable with those made by

480   the previous authors, even though viscosity was not included in their studies. This

481   suggests that nonlinear friction and viscosity play the first and second roles,

482   respectively, on the intermittent phases. The intermittent phases could be considered

483   as foreshocks of the mainshock which is associated with the main rupture. Simulation

484   results exhibit that foreshocks happen for some mainshcoks and not for others.

485

486   **6. Conclusions**

487   In this work, multistable slip of earthquakes caused by slip-weakening friction

488   and viscosity is studied based on the normalized equation of motion of a one-degree-

489   of-freedom spring-slider model in the presence of the two factors. The friction is

490   caused by thermal pressurization and decays exponentially with displacement. The

491   major model parameters are the normalized characteristic distance, U_c, for friction

Nonlinear Processes
in Geophysics
Discussions
Open Access

492   and the normalized viscosity coefficient, η, between the slider and the background

493   moving plate, which exerts a driving force on the former. Analytic results at small U

494   suggest that there is a solution regime for η and γ (=$1/U_c$) to make the slider slip

495   steadily. Numerical simulations lead to the time variation in V/V_{max}, the phase portrait

496   of V/V_{max} versus U/U_{max}, and Fourier spectrum. Results show that the time variation

497   in V/V_{max}, obviously depends on U_c and η. The effect on the amplitude is stronger

498   from η than from U_c. When $U_c>1$, the time variation of V/V_{max} exhibits cyclic

499   oscillations with a single period and the amplitude of V/V_{max} remarkably decreases

500   with increasing U_c. When $U_c<1$, slip changes from stable motion to intermittency and

501   then to unstable motion when η increases. For a certain U_c, the three regimes are

502   controlled by a lower bound, η_l, and an upper bound, η_u, of η. When $U_c<0.55$, η_u is

503   absent and thus unstable or chaotic slip does not exist. When $U_c\geq0.55$, the plots of η_l

504   and η_u versus U_c exhibit a stable regime when $\eta\leq\eta_l$, an intermittency (or transition)

505   regime when $\eta_l<\eta\leq\eta_u$, and unstable (or chaotic) regime when $\eta>\eta_u$. The values of η_l,

506   η_u, and η_u-η_l all decrease with increasing U_c, thus suggesting that it is easier to yield

507   unstable slip for larger U_c than for smaller U_c or larger η than for smaller η. The

508   phase portraits of V/V_{max} versus U/U_{max} exhibit that there are two fixed points: The

509   first one at large V/V_{max} and large U/U_{max} is not an attractor for all cases in study;

510   while the second one at small V/V_{max} and small U/U_{max} can be an attractor for some

511   values of U_c and η. When $U_c<1$, the Fourier spectra calculated from simulation

512   velocity waveforms exhibit several peaks rather than one, thus suggesting the

513   existence of nonlinear behavior of the system. When $U_c>1$, the related Fourier spectra

514   show only one peak, thus suggesting linear behavior of the system.

515

519

520   **References**

521   Abe, Y. and N. Kato (2013). Complex earthquake cycle simulations using a two-

522   degree-of-freedom spring-block model with a rate- and state-friction law. Pure

523   Appl. Geophys., 170, 745-765.

524   Belardinelli, M.E. and E. Belardinelli (1996). The quasi-static approximation of the

Nonlinear Processes
in Geophysics
Open Access
Discussions
EGU

525   spring-slider motion. Nonl. Process Geophys., 3, 143-149.

526   Bizzarri, A. (2009). What does control earthquake ruptures and dynamic faulting? A

527   review of different competing mechanism. Pure Appl. Geophys., 166, 741-776.

528   Bizzarri, A. (2010). An efficient mechanism to avert frictional melts during seismic

529   ruptures. Earth Planet. Sci. Lett., 296, 144-152, doi:10.1016/j.epsl.2010.05.012.

530   Bizzarri, A. (2011a). Dynamic seismic ruptures on melting fault zones. J. Geophys.

531   Res., 116, B02310, doi:10.1029/2010JB007724.

532   Bizzarri, A. (2011b). Temperature variations of constitutive parameters can

533   significantly affect the fault dynamics. Earth Planet. Sci. Lett., 306, 72-278, doi:

534   10.1016/j.epsl.2011.04.009.

535   Bizzarri, A. (2011c). On the deterministic description of earthquakes. Rev. Geophys.,

536   49, RG3002, doi:10.1029/2011RG000356.

537   Bizzarri, A. and M. Cocco (2006a). A thermal pressurization model for the

538   spontaneous dynamic rupture propagation on a three-dimensional fault: 1.

539   Methodological approach. J. Geophys. Res., 111, B05303, doi:10.1029/

540   2005JB003862.

541   Bizzarri, A. and M. Cocco (2006b). A thermal pressurization model for the

542   spontaneous dynamic rupture propagation on a three-dimensional fault: 2.

543   Traction evolution and dynamic parameters. J. Geophys. Res., 111, B05304,

544   doi:10.1029/2005JB003864.

545   Brun, J.L. and A.B. Gomez (1994). A four-parameter, two degree-of-freedom

546   block-spring model: Effect of the driver velocity. Pure Appl. Geophys., 143(4),

547   633-653

548   Burridge, R. and L. Knopoff (1967). Model and theoretical seismicity. Bull. Seism.

549   Soc. Am., 57, 341-371.

550   Byerlee, J.D. (1968). Brittle-ductile transition in rocks. J. Geophys. Res., 73,

551   4711-4750.

552   Carlson, J.M. and J.S. Langer (1989). Mechanical model of an earthquake fault. Phys.

553   Rev. A, 40, 6470-6484.

554   Chester, F.M., and H.G. Higgs (1992). Multimechanism friction constitutive model

555   for ultrafine quartz gouge at hypocentral conditions. J. Geophys. Res., 97, B2,

556   1859-1870.

557   Cohen, S. (1979). Numerical and laboratory simulation of fault motion and

558   earthquake occurrence. Rev. Geophys. Space Phys., 17(1), 61-72.

Nonlinear Processes
in Geophysics

Discussions

Open Access

EGU

559   Diniega, S., S.E. Smrekar, S. Anderson, and E.R. Stofan (2013). The influence of

560   temperature-dependent viscosity on lava flow dynamics. J. Geophys. Res., 118,

561   1516-1532, doi:10.1002/jgrf.20111.

562   Dragoni, M. and S. Santini (2015). A two-asperity fault model with wave

563   radiation. Phys. Earth Planet. Inter., 248, 83-93.

564   Erickson, B., B. Birnir1, and D. Lavall´ee (2008). A model for aperiodicity in

565   earthquakes. Nonlin. Process Geophys., 15, 1-12.

566   Erickson, B.A., B. Birnir, and D. Lavall´ee (2011). Periodicity, chaos and localization

567   in a Burridge–Knopoff model of an earthquake with rate-and-state friction.

568   Geophys. J. Int., 187, 178-198, doi:10.1111/j.1365-246X.2011.05123.x.

569   Fialko, Y.A. (2004). Temperature fields generated by the elastodynamic propagation

570   of shear cracks in the Earth. J. Geophys. Res., 109, B01303, doi:10.1029/

571   2003JB002496.

572   Fialko, Y.A. and Y. Khazan (2005). Fusion by the earthquake fault friction: stick or

573   slip?. J. Geophys. Res., 110, B12407, doi.org/10.1029/2005JB003869.

574   Galvanetto, U. (2002). Some remarks on the two-block symmetric Burridge–Knopoff

575   model. Phys. Letts. A, 251-259.

576   Gu, J.C., J.R. Rice, A.L. Ruina, and S.T. Tse (1984). Slip motion and stability of a

577   single degree of freedom elastic system with rate and state dependent friction. J.

578   Phys. Solid, 32, 167-196.

579   Huang, J. and D.L. Turcotte (1990). Are earthquakes an example of deterministic

580   chaos?. Geophys. Res. Lett., 17(3), 223-226.

581   Huang, J. and D.L. Turcotte (1992). Chaotic seismic faulting with a mass-spring

582   model and velocity-weakening friction. Pure Appl. Geophys., 138(4), 549-589.

583   Hudson, J.A. (1980). The excitation and propagation of elastic waves. Cambridge

584   Monographs on Mechanics and Applied Mathematics, Cambridge Univ. Press,

585   224pp.

586   Jaeger, J.C. and N.G.W. Cook (1977). Fundamentals of Rock Mechanics. John Wiley

587   & Sons, Inc., New York, 585pp.

588   Jeffreys, H. (1942). On the mechanics of faulting. Geol. Mag., 79, 291.

589   Kittel, C., W.D. Knight, and M.A. Ruderman (1968). Mechanics, Berkeley Physics

590   Course Volume 1, McGraw-Hill Book Co., New York, N.Y., 480pp.

591 Knopoff, L. and X.X. Ni (2001). Numerical instability at the edge of a dynamic
592   fracture. Geophys. J. Int., 147, F1-F6.

593 Knopoff, L., J.Q. Mouton, and R. Burridge (1973). The dynamics of a one-
594   dimensional fault in the presence of friction. Geophys. J. R. astro. Soc., 35,
595   169-184.

596 Kostić, S., I. Franović, K. Todorović, and N. Vasovi (2013). Friction memory effect
597   in complex dynamics of earthquake model. Nonlin. Dyn., 73, 1933-1943, DOI:10.
598   1007/s11071-013-0914-8.

599 Lachenbruch, A.H. (1980). Frictional heating, fluid pressure, and the resistance to
600   fault motion. J. Geophys. Res., 85, 6097-6122.

601 Lorenz, E.N. (1963). Deterministic non-periodic flow. J. Atmos. Sci., 20(2), 130-141.

602 Lu, X., A.J. Rosakis, and N. Lapusta (2010). Rupture modes in laboratory earthquakes:
603   Effect of fault prestress and nucleation conditions. J. Geophys. Res., 115,
604   B12302, doi:10.1029/2009JB006833.

605 Nussbaum, J. and A. Ruina (1987). A two degree-of-freedom earthquake model with
606   static/dynamic friction. Pure Appl. Geophys., 125(4), 629-656.

607 Parez Pascual, R. and J. Lomnitz-Adler (1988). Coupled relaxation oscillators and
608   circle maps. Physica D, 30, 61-82.

609 Press, W.H., B.P. Flannery, S.A. Teukolsky, and W.T. Vetterling (1986). Numerical
610   Recipes. Cambridge Univ. Press, Cambridge, 818pp.

611 Rice, J.R. (2006). Heating and weakening of faults during earthquake slip. J. Geophys.
612   Res., 111, B05311, doi:10.1029/2005JB004006.

613 Rice, J.R., N. Lapusta, and K. Ranjith (2001). Rate and state dependent friction and
614   the stability of sliding between elastically deformable solids. J. Mech. Phys.
615   Solids, 49, 1865-1898.

616 Ryabov, V.B. and K. Ito (2001). Intermittent phase transitions in a slider-spring
617   model as a mechanism for earthquakes. Pure Appl. Geophys., 158, 919-930.

618 Scholz, C.H. (1990). The Mechanics of Earthquakes and Faulting. Cambridge Univ.
619   Press, Cambridge, 439pp.

620 Sibson, R.H. (1973). Interaction between temperature and pore-fluid pressure during
621   earthquake faulting and a mechanism for partial or total stress relief. Natural
622   Phys. Sci., 243, 66-68.

623 Shkoller and Minster (1997). Reduction of Dieterich-Ruina attractors to unimodal
624   maps. Nonl. Process Geophys., 4, 63-69.

Nonlinear Processes
in Geophysics
Discussions

Open Access

EGU

625   de Sousa Vieira, M. (1999). Chaos and synchronized chaos in an earthquake model.
626   Phys. Rev. Letts., 82(1), 201-204.

627   Spray, J.G. (1993). Viscosity determinations of some frictionally generated silicate
628   melts: Implications for fault zone rheology at high strain rates. J. Geophys. Res.,
629   98(B5), 8053-8068.

630   Spray, J.G. (1995). Pseudotachylyte controversy: Fact or friction?. Geology, 23(12),
631   1119-1122.

632   Spray, J.G. (2005). Evidence for melt lubrication during large earthquakes. Geophys.
633   Res. Lett., 32, L07301, doi:10.1029/2004GL022293.

634   Thompson, J.M.T. and H.B. Stewart (1986). Nonlinear Dynamics and Chaos. John
635   Wiley and Sons, New York, 376pp.

636   Turcotte, D.L. (1992). Fractals and chaos in geology and geophysics. Cambridge Univ.
637   Press, London, 221pp.

638   Turcotte, D.L. and G. Schubert (1982). GEODYNAMICS – Applications of
639   Continuum Physics to Geological Problems. Wiley, 450pp.

640   Wang, J.H. (1995). Effect of seismic coupling on the scaling of seismicity. Geophys. J.
641   Int., 121, 475-488.

642   Wang, J.H. (1996). Velocity-weakening friction law as a factor in controlling the
643   frequency-magnitude relation of earthquakes. Bull. Seism. Soc. Am., 86, 701-
644   713.

645   Wang, J.H. (2000). Instability of a two-dimensional dynamical spring-slider model of
646   an earthquake fault. Geophys. J. Int., 143, 389-394.

647   Wang, J.H. (2006). Energy release and heat generation during the 1999 Chi-Chi,
648   Taiwan, earthquake. J. Geophys. Res., 111, B11312, doi:10.1029/2005JB004018.

649   Wang, J.H. (2007). A dynamic study of the frictional and viscous effects on
650   earthquake rupture: a case study of the 1999 Chi-Chi earthquake, Taiwan. Bull.
651   Seism. Soc. Am., 97(4), 1233-1244.

652   Wang, J.H. (2008). One-dimensional dynamical modeling of earthquakes: A review,
653   Terr. Atmos. Ocean. Sci., 19, 183-203.

654   Wang, J.H. (2009). Effect of thermal pressurization on the radiation efficiency. Bull.
655   Seism. Soc. Am., 99(4), 2293-2304.

656   Wang, J.H. (2011). Thermal and pore fluid pressure history on the Chelungpu fault at
657   a depth of 1111 meters during the 1999 Chi-Chi, Taiwan, earthquake. J. Geophys.
658   Res., 116, B03302, doi:10.1029/2010JB007765.

659 Wang, J.H. (2012). Some intrinsic properties of the two-dimensional
660   dynamical spring-slider model of earthquake faults. Bull. Seism. Soc. Am.,
661   102(2), 822-835.

662 Wang, J.H. (2013). Stability analysis of slip of a one-body spring-slider model in the
663   presence of thermal pressurization. Ann. Geophys., 56(3), R03332, doi:10.4401/
664   ag-5548.

665 Wang, J.H. (2016a). A dynamical study of frictional effect on scaling of earthquake
666   source displacement spectra. Ann. Geophys., 59(2), S0210, 1-14, doi:10.4401/
667   ag-6974.

668 Wang, J.H. (2016b). Slip of a one-body spring-slider model in the presence of
669   slip-weakening friction and viscosity. Ann. Geophys., 59(5), S0541, DOI:10.
670   4401/ag-7063.

671 Wang, J.H. (2017). Slip of a two-degree-of-freedom spring-slider model in the
672   presence of slip-weakening friction and viscosity., Ann. Geophys. (in press)

673 Xu, H.J. and L. Knopoff (1994). Periodicity and chaos in a one-dimensional
674   dynamical model of earthquakes. Phys. Rev. E, 50(5), 3577-3581.

675

676

677

678

679

680   Table 1. Values of η_l, η_u, and V_{max} for various U_c.

U_c	η_l	η_u	V_{max}
0.20	0.87	1.00	0.4068
0.25	0.86	1.00	0.3611
0.30	0.86	1.00	0.3149
0.35	0.77	1.00	0.2905
0.40	0.69	1.00	0.2649
0.45	0.57	1.00	0.2497
0.50	0.51	1.00	0.2216
0.55	0.43	0.98	0.1989
0.60	0.39	0.84	0.1684
0.65	0.38	0.78	0.1338
0.70	0.34	0.72	0.1071
0.75	0.26	0.69	0.0879
0.80	0.25	0.55	0.0604
0.85	0.18	0.48	0.0423
0.90	0.14	0.37	0.0234
0.95	0.12	0.25	0.0076

681

682

Nonlinear Processes
in Geophysics

Open Access

Discussions

EGU

683

684

685

686
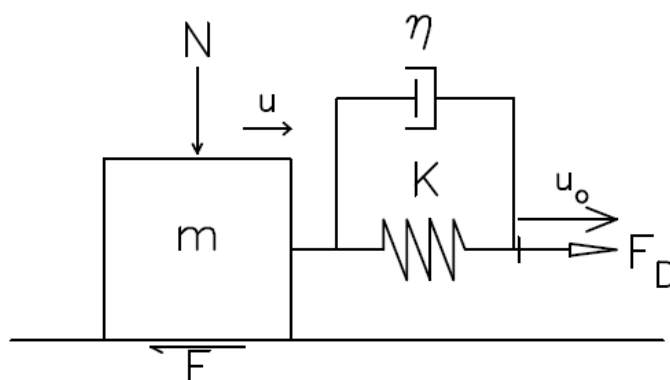
687
688   Figure 1. One-body spring-slider model. In the figure, u, K, η, F_D, N, and F denote,
689   respectively, the displacement, the spring constant, the viscosity coefficient, the
690   driving force, the normal force, and the frictional force.
691

692

693

694

695
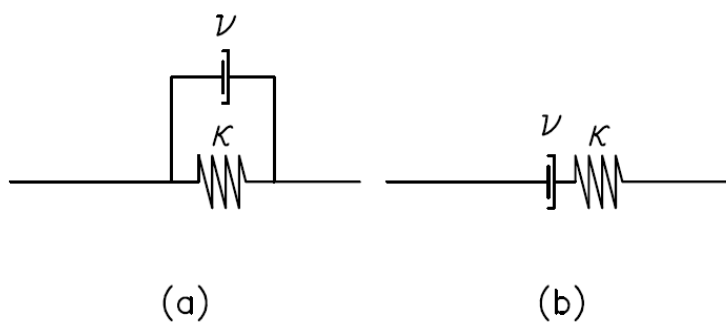
Figure 2. The two types of viscous materials: (a) for the Kelvin–Voigt model and (b) for the Maxwell model. (κ=spring constant and υ=coefficient of viscosity)

696

697

698

699

700

701

702

703
704

705
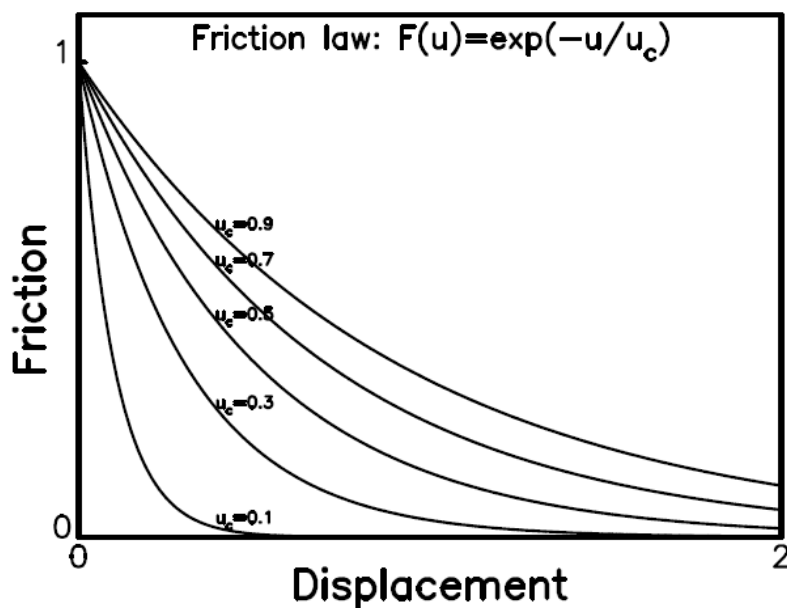
706

707 Figure 3. The variations in friction force with displacement for F(u)=exp(-u/u_c) when
708   u_c=0.1, 0.3, 0.5, 0.7, and 0.9 m (after Wang, 2016b).

709

710

711

712

713
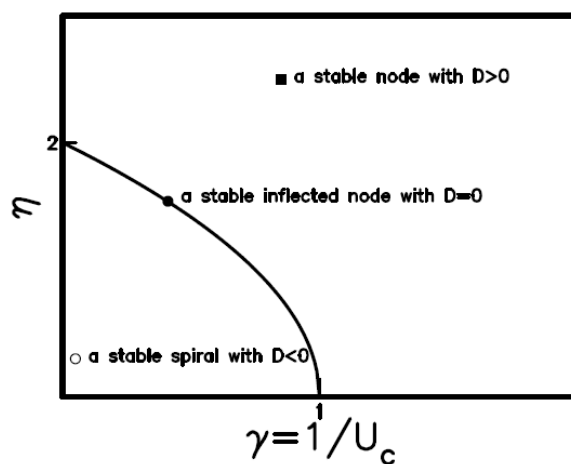
714
715 Figure 4. The plot of η versus $1/U_c$ exhibits the phase portrait and root structure of the
716 system. The solid line displays the function: $D(\eta,1/U_c)=\eta^2-4(1-1/U_c)=0$. The
717 solid circle, open circle, and solid square represent, respectively, a stable
718 inflected node with D=0, a stable spiral with D<0, and a stable node with D>0.

719
720

721

722

723
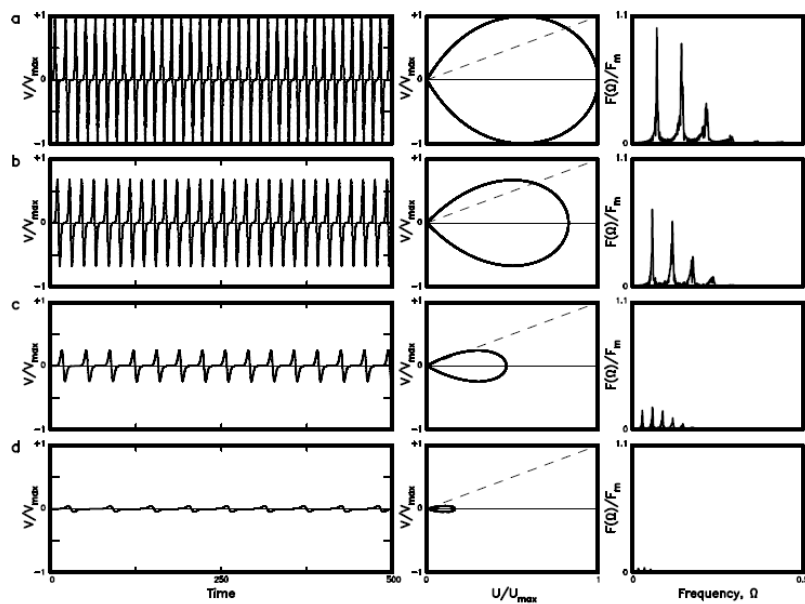
724
725   Figure 5. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
726   and power spectrum for four values of U_c: (a) for $U_c=0.1$; (b) for $U_c=0.4$; (c) for
727         $U_c=0.7$; and (d) for $U =0.9$ for the TP law of $F(U)=\exp(-U/U_c)$ when $\eta=0$.
728
729
730
731
732
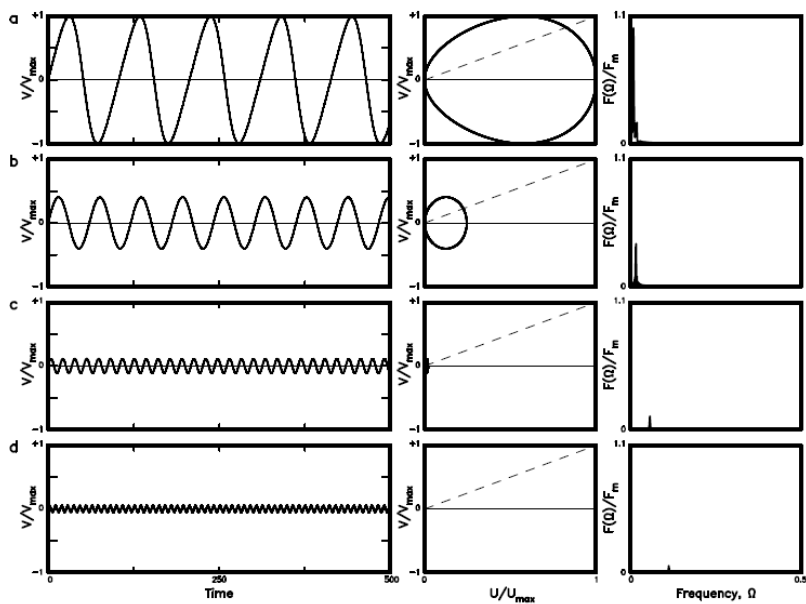
733
734   Figure 6. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
735   and power spectrum for four values of U_c: (a) for $U_c=1.00$; (b) for $U_c=1.01$; (c)
736   for $U_c=1.15$; and (d) for U $=2.00$ for the TP law of $F(U)=\exp(-U/U_c)$ when $\eta=0$.
737
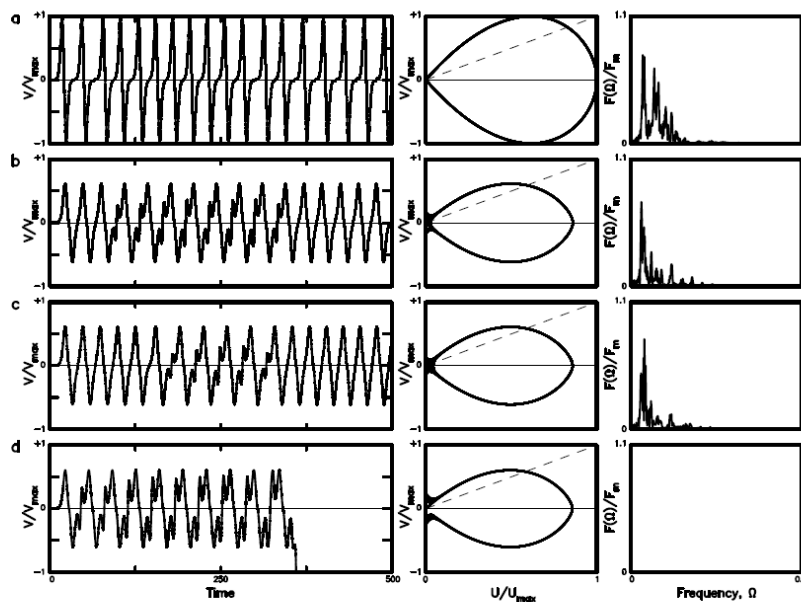738

739

740

741
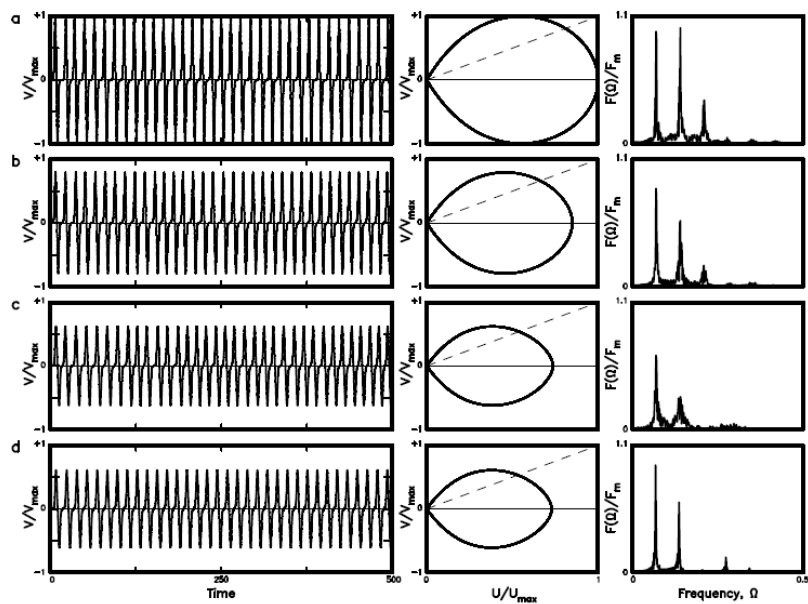
742
743   Figure 7. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
744   and power spectrum for four values of η: (a) for η=0.20; (b) for η=0.50; (c) for
745   η=0.87; and (d) for η=0.90 when U_c=0.20 for the TP law of $F(U)=\exp(-U/U_c)$.
746

747

748

749

750
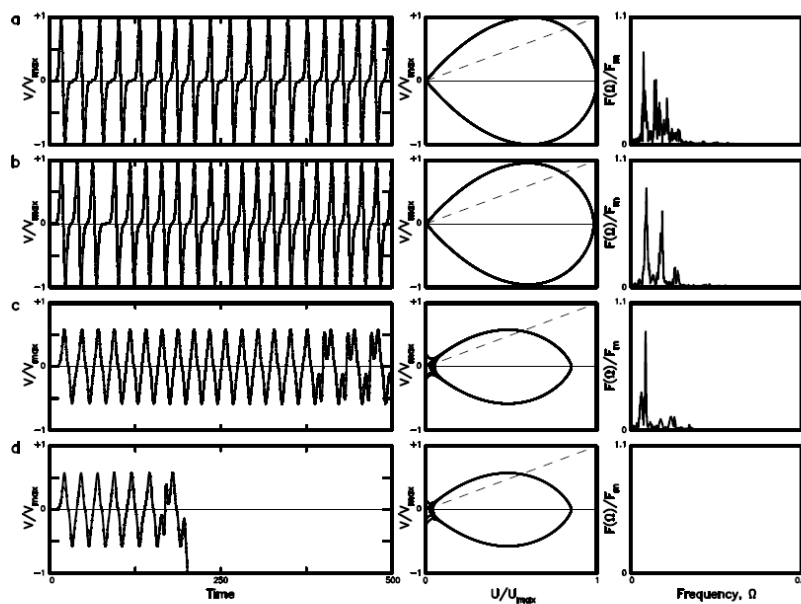
751
752   Figure 8. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
753   and power spectrum for four values of η: (a) for η=0.43; (b) for η=0.47; (c) for
754   η=0.98; and (d) for η=0.99 when U_c=055 for the TP law of $F(U)=exp(-U/U_c)$.
755

Nonlinear Processes
in Geophysics

Open Access

Discussions

EGU

756

757

758

759
760   Figure 9. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
761   and power spectrum for four values of η: (a) for $\eta=0.39$; (b) for $\eta=0.83$; (c) for
762         $\eta=0.84$; and (d) for $\eta=0.85$ when $U_c=0.6$ for the TP law of $F(U)=\exp(-U/U_c)$.
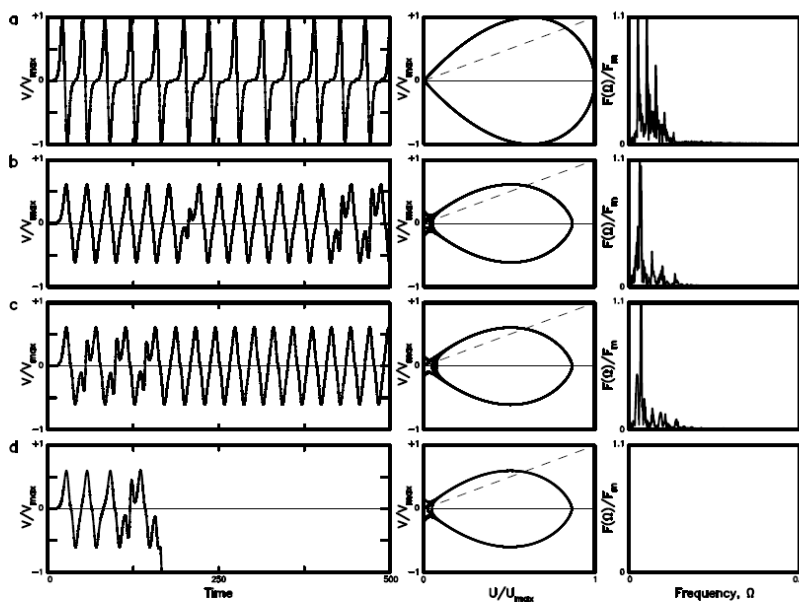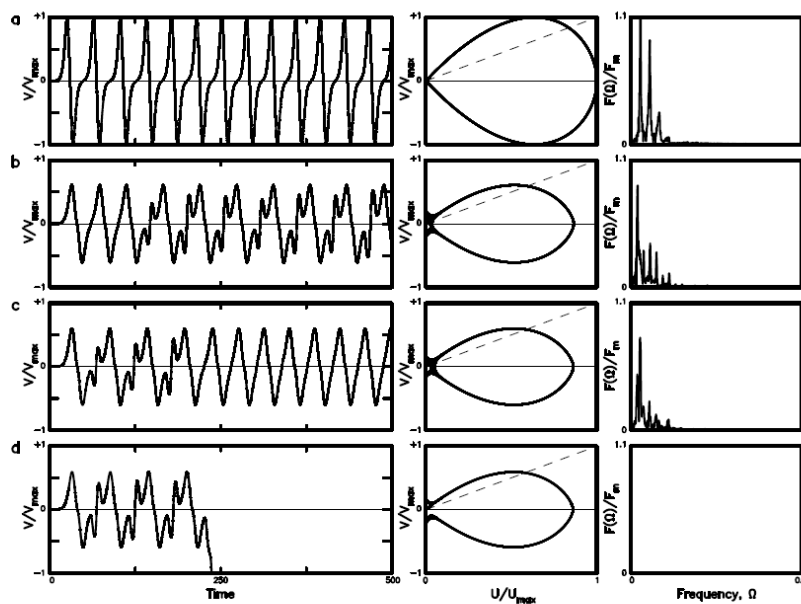763

764
765

Nonlinear Processes
in Geophysics
Discussions

766

767

768

769
770   Figure 10. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
771   and power spectrum for four values of η: (a) for η=0.34; (b) for η=0.71; (c) for
772   η=0.72; and (d) for η=0.73 when U_c=0.7 for the TP law of $F(U)=\exp(-U/U_c)$.
773

774
775
776

777
778   Figure 11. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
779   and power spectrum for four values of η: (a) for $\eta=0.25$; (b) for $\eta=0.54$; (c) for
780         $\eta=0.55$; and (d) for $\eta=0.56$ when $U_c=0.8$ for the TP law of $F(U)=\exp(-U/U_c)$.
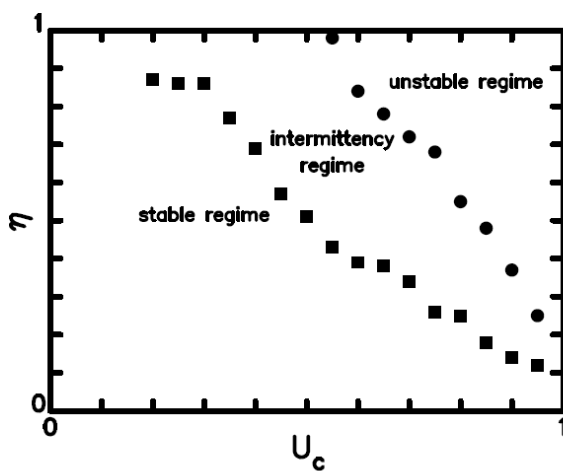
781
782

783

784

785
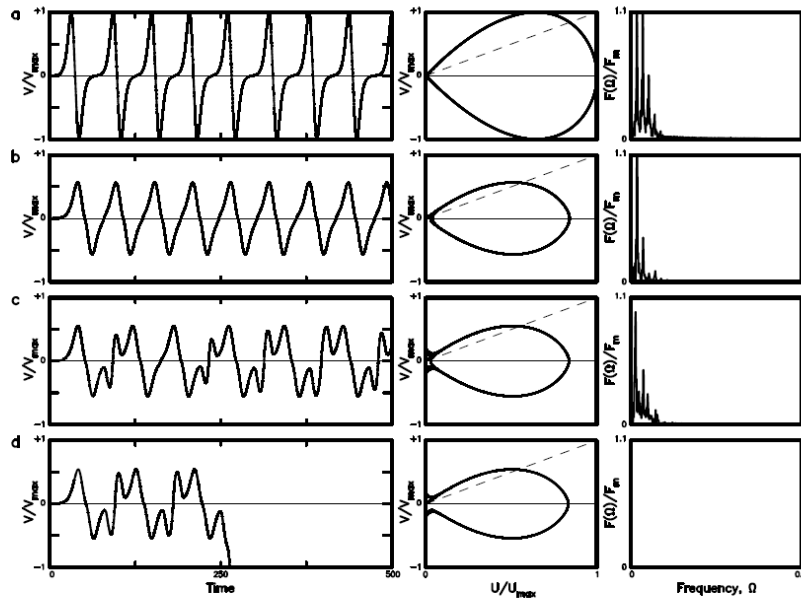
786
787   Figure 12. The time variation in V/V_{max}, the phase portrait of V/V_{max} versus U/U_{max},
788   and power spectrum for four values of η: (a) for $\eta=0.14$; (b) for $\eta=0.36$; (c) for
789        $\eta=0.37$; and (d) for $\eta=0.38$ when $U_c=0.9$ for the TP law of $F(U)=\exp(-U/U_c)$.
790
791

792

793

794

795
796   Figure 13. The plot of η_l (with a solid square) and η_u (with a solid circle) versus U_c.