## 1 Identification of magnetic anomalies based on ground

# 2 magnetic data analysis using multifractal modeling: A case

3 study in Qoja-Kandi, East Azerbaijan Province, Iran

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## 12 Abstract

13 Ground magnetic anomaly separation using Reduction-To-the-Pole (RTP) technique and the 14 fractal concentration-area (C-A) method have has been applied to the Qoja-Kandi 15 prosepecting area in NW Iran. The geophysical survey that resulted resulting in the ground 16 magnetic data was conducted for magnetic elements exploration. Firstly, RTP technique was 17 applied for recognizing underground magnetic anomalies. RTP anomalies was classified into 18 different populations based on this the current method. For this reason, drilling points point 19 areas determination with by RTP technique was complicated. Next, C-A method was applied 20 on the RTP-Magnetic-Anomalies (RTP-MA) for demonstrating magnetic susceptibility 21 concentrations. This identification was appropriate for increasing the resolution of the drilling 22 points determination and decreasing the drilling risk issue, due to the economic costs of 23 underground prospecting. In this study, the results of C-A Modeling on the RTP-MA are 24 compared with 8 boreholes data. The results shows that there is a good 25 correlationsatisfying correlation between anomalies derived via C-A method and log report of 26 boreholes. Two boreholes were drilled in magnetic susceptibility concentrations, based on 27 multifractal modeling data analyses, between 63533.1 to 66296 nT. Drilling results showshowed appropriate magnetite thickness with the grades greater than 20% Fe total. 28

**Comentado** [N1]: Here, you should explain in what area is complicated?

Also, total associated with anomalies associated with containing and esite units host iron
 mineralization.

#### 3 1 Introduction

4 Mineral exploration aims to at discovering new mineral deposits in a region of interest (Abedi 5 et al., 2013). These mineral deposits could be related to magnetic anomalies which are 6 situated within underground. In the first steps of identification underground magnetic 7 anomalies, few boreholes should be drilled after interpretation Ground magnetic data. 8 Obviously, using new methods could increase the resolution of the drilling points 9 determination and decrease the drilling risk. A cursory look at magnetic maps will indeed 10 tellwould present more information much about about the shape of such a buried features. 11 However, the information acquired an informed study of a from map can provide additional 12 details about the specification of underground magnetic anomalies especially exact locations. 13 Because of depending magnetic anomaly on the inclination and declination of the body's magnetization, the inclination and declination of the local earth's magnetic field, and the 14 15 orientation of the body with respect to magnetic north (Baranov, 1957) and (Baranov and Naudy, 1964) proposed a mathematical approach known as reduction-to-the-pole (RTP) for 16 simplifying anomaly shape and determining anomaly exact location. As a result of increasing 17 the resolution of RTP technique, concentration-area (C-A) fractal method was applied. Fractal 18 19 geometry is a Non-Euclidean geometry established by Mandelbrot (1983) and has been 20 applied in geosciences and mineral exploration, especially in geophysical and geochemical 21 exploration since 1980s, e.g.( Turcotte (1989), Bolviken et al. (1992), Korvin (1992), Cheng 22 et al. (1994), Agterberg et al. (1996), Cheng (1999), Turcotte (2004), Dimri (2005) and Shen 23 et al. (2009)).

In this study, concentration-area (C-A) fractal method was used to gridded RTP data set, for better classification of RTP map which generated from RTP technique. This procedure was applied to the ground magnetic data of Qoja-Kandi, Zanjan Province, Iran.

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#### 28 2 The concentration-area fractal method

The concentration-area (C-A) method serves to illustrate the <u>correlated</u> relationship <del>correlated</del>
between the obtained results. Its most useful features are the easy implementation and the
ability to compute quantitative anomalous thresholds (<u>cheng Cheng</u> et al., 1994).

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Cheng et al. (1994) proposed the concentration-area (C-A) method for separating
 geochemical anomalies from background in order to characterize the distribution of elemental
 concentrations. Equations (1) Shows the general form of this model.

## 4 $A(\rho \leq v) \propto \rho^{-a_1}; A(\rho \geq v) \propto \rho^{-a_2}$ (1)

5 Where  $A(\rho)$  denotes the area with concentration values greater than the contour value  $\rho$ ;  $\nu$  represents the threshold; and a1 and a2 are characteristic exponents. The breaks between 6 7 straight line segments in C-A log-log plot and the corresponding values of p are known as 8 thresholds to separate geophysical values into different components representing different 9 causal factors such as, lithological differences, geochemical processes and mineralizing 10 events (Lima et al., 2003). Thus, applying C-A fractal model to the geochemical data, 11 improves resolution of the data helping to explore the deposits. It seems that, applying this 12 model to ground magnetic data improves the accuracy of magnetite deposit exploration. The 13 most useful feature of the C-A method is its capability to compute anomaly thresholds 14 (Goncalves et al., 2001). Using fractal theory, Cheng et al. (1994) derived similar power-law 15 relationships and equations in extended form. The area A ( $\rho$ ) for a given  $\rho$  is equal to the 16 number of cells multiplied by cell area with concentration values greater than p. Average 17 concentration values are used for those boxes containing more than one sample. Area-18 concentration [A ( $\rho$ )] with element concentrations greater than  $\rho$  usually shows a power-law 19 relation (cheng et al., 1994). The study area and geological setting

20 The Qoja-Kandi area is located within the Orumieh-Dokhtar magmatic arc in northwest 21 of Iran (Fig. 1); This magmatic arc is the most important exploratory area for metals, and 22 hosts the majority of the larger metals deposits such as copper and iron (Hassan-Nezhad and 23 Moore, 2006). The investigated area characterized by Precambrian to Jurassic units and 24 Oligo-Miocene volcanic rocks. Different types of metal ore deposits, such as iron have 25 already been documented near studied area. The lithology of this part includes schist and 26 shale (Kahar formation), dolomite and limestone (Elika formation), shale, sandstone and 27 limestone (Shemshak formation), limestone, marl, sandstone, conglomerate and andesit. A magnetite dyke which has outcrops in andesite units has already been seen near studied area. 28 29 It seems that this magnetite dyke presence in Qoja-Kandi area.

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#### 1 3 Ground magnetic data analysis

- 2 Ground magnetic data are acquired in the region at 15 m spacing along lines in the north 3 direction and spaced 10 m apart. 6997 geophysical ground data were collected by **GSM**-
- 4 19T proton.

#### 5 3.1 The TMI anomaly map

6 The Total-Magnetic-Intensity (TMI) map of the Qoja-Kandi area was obtained to delineate 7 the subsurface anomaly. Fig. 2 indicates TMI with ground magnetic data points. The ground 8 magnetic anomalies range from 38633 to 69509 nT and are characterized by both low and 9 high frequencies of anomalies. The map reveals that dipolar (anomalies having positive and 10 negative components) magnetic anomalies have a general E-W direction, which is in the 11 center and north of studied area. There are three obvious dipolar magnetic anomalies (two 12 anomalies in the east and west of the center and one anomaly in the north) in the Qoja-Kandi 13 prospecting area which are expected to depend on two magnetite dyke in andesite units.

#### 14 **3.2** Reduction to the pole technique

15 An <u>difficulty in interpretation</u> difficulty with TMI anomalies is that they are dipolar 16 (anomalies having positive and negative components) such that the shape and phase of the 17 anomaly depends in-on the part on of the magnetic inclination and the presence of any 18 remanent magnetization. Because of depending magnetic anomaly on the inclination and 19 declination of the body's magnetization, the inclination and declination of the local earth's 20 magnetic field, and the orientation of the body with respect to magnetic north, (Baranov, 21 1957) and (Baranov and Nudy, 1964) proposed a mathematical approach known as reduction 22 to the pole for simplifying anomaly shape.

The reduction-to-the-pole (RTP) technique transforms TMI anomalies to anomalies that would be measured if the field were vertical (assuming there is only an inducing field). This RTP transformation makes the shape of magnetic anomalies more closely related to the spatial location of the source structure and makes the magnetic anomaly easier to interpret, as anomaly maxima will be located centrally over the body (provided there is no remanent magnetization present). Thus, the RTP reduces the effect of the Earth's ambient magnetic field and provides a more accurate determination of the position of anomalous sources. It is

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therefore understood that the total magnetization direction is equivalent to that of the current
 inducing filed.

3 Before applying the methods, the total field anomaly data were converted to RTP using a magnetic inclination of 55.43° and a declination of 4.93°. RTP anomalies, shows three 4 5 obvious magnetic anomalies (two anomalies in the east and west of the south and one 6 anomaly in the north) in the studied area, elongated in approximate E-W direction. The 7 highest class of RTP-Magnetic-Anomalies (RTP-MA) based on Reduction to the pole technique is > 55370.7 nT with 24941.79 square meters area. Also, RTP anomalies was 8 9 classified to different populations based on this method, as illustrated in Fig. 3. Based on this method, drilling points determination with RTP technique was complicated. 10

#### 11 3.3 Application of C-A Modeling on the RTP-MA

Multifractal models are utilized to quantify patterns such as geophysical data. Fractal and 12 13 multifractal modeling are widely applied to distinguish the different mineralized zones 14 (Cheng, 2007). Multifractal theory could be interpreted as a theoretical framework that 15 explains the power law relationships between areas enclosing concentrations below a given 16 threshold value and the actual concentrations itself. To demonstrate and prove that data 17 distribution has a multifractal nature, an extensive computation is required (Halsey et al., 1986; Evertz and Mandelbrot, 1992). This method has several constrains especially when the 18 19 boundary effects on irregular geometrical data sets are involved (Agterberg et al., 1996; Goncalves, 2001; Cheng, 2007; Xie et al., 2010). Multifractal modelings in geophysical and 20 geochemical exploration help to find exploration targets and mineralization potentials in 21 22 different types of deposits (Yao and Cheng, 2011). The C-A method seems to be equally 23 applicable to all cases which means that geophysical distributions mostly satisfy the 24 properties of a multifractal function. There is some evidence that geophysical and 25 geochemical data distributions have fractal behavior in nature, e.g. Bolviken et al. (1992), Turcotte (1997), Goncalves (2001), Gettings (2005) and Li and Cheng (2006). This theory 26 27 improves the development of an alternative interpretation validation and useful methods to be 28 applied to geophysical distributions analysis.

In this study, 57307 transformed RTP data were processed for identification of magnetic
 anomalies. Statistical results reveal that RTP-MA mean value is 48441 nT, as depicted in Fig.

4, and the RTP-MA domain shows a wide range. RTP-MA distribution map was generated

with minimum curvature method. The estimated RTP-MA model in terms of RTP data values 1 2 was intended to build of the C-A log-log plot for RTP-MA. Based on linear segments and 3 breakpoints log-log plot, as shown in Fig. 5, geophysical population were divided. RTP threshold values are 45383, 47424.2, 49493.7, 56493.7 and 635331.1 which are very low, 4 low, moderate, high and very high intensity anomaly threshold values, respectively, as 5 illustrated in Table 1. These geophysical populations were determined based on the 6 7 breakpoints in log-log plot. Actually Actually the length of the tangent, demonstrate the 8 extents of geophysical populations in fractal model. It is mentioned that the number of 9 population in fractal model could be more or less than five, but actually the extent of the last 10 class population isn't highly dependent on the number of population in fractal model. Hence, there are five populations for RTP-MA which illustrate that fifth class of RTP-MA based on 11 12 fractal method is > 63533.1 nT with very high priority for drilling. Consequently, the 13 locations of RTP-MA (two anomalies) based on fractal method are situated in the east of 14 southern part of the area, as depicted in Fig. 6.

#### 15 4 Control with borehole data

A method of investigating subsurface geology is, of course, drilling boreholes. For a more accurate results about identification of magnetic anomalies, the results of C-A Modeling on the RTP-MA are compared with borehole data (Table 2). There are 8 drilled boreholes in this area that are used for identification of magnetic anomalies obtained from boreholes (Fig. 7). The drilled boreholes were analyzed and studied by geologists. Hence, range of magnetite ores in each borehole were obtained and documented as log report in Table 2. The accepted lower limit for the ore length, is the grade 20% Fe total.

23 RTP transformed data based on ground magnetic anomaly data collected from C-A moderate 24 anomalies in Qoja-Kandi prospecting area show magnetic susceptibility concentration 25 between 63533.1 to 66296 nT with 1957.64 square meters area. This study shows that the 26 areas with very high priority obtained by C-A method have magnetite concentration with 27 appropriate thickness. This point is significant that borehole 1 and 2 were drilled in mentioned places and confirmed the results of C-A model (Fig. 8) for increasing the resolution of drilling 28 29 point determination and decreasing the drilling risk. Fig. 8 shows 3D RTP map of Qoja-Kandi 30 based on C-A method with pictures from magnetite zones in the surface of drilled borehole1 and 2, in addition of mentioned boreholes log plots. It is necessary to mention that, the 31 32 TERRA satellite has a back-looking telescope with a resolution of 15 m in the VNIR that

matches with the wavelength of the band 3 that is used to extract 3D information for provided
 Fig. 8.

3 The results confirmed there is affirmative correlation between anomalies derived via C-A 4 method and log report of boreholes. Furthermore, the ratio of the ore length and total core 5 length is calculated in Table 2. The number of this ratio is between ranges of 0 to 1. Whatever this number is larger and close to 1, the resolution of the drilling point determination increase 6 7 and the drilling risk decrease. The results shows positive correlation between the ratio of the 8 ore and total core column, and Priority areas for drilling column. Based on this study, 9 anomalies associated with andesite units host iron mineralization. Also, there isn't any 10 mineralization in other geological units such as limestones and conglomerates in northwest of 11 the studied area. It should be noted that, magnetite ores have outcrops in andesite units (Fig. 12 8).

## 13 5 Conclusions

Separation of magnetic anomalies using combine of RTP technique and C-A fractal modelling has been used in Qoja-Kandi prospecting area as a new geophysical method for increasing the resolution of the drilling points determination. This study demonstrates that C-A method utilizaing for ground magnetic anomaly separation is an appropriate manner for geophysical prospecting.

There was a multifractal model for RTP-MA, based on Log-log plots in the prospecting area.In this paper, RTP anomalies results from C-A method and RTP technique were compared. Anomalies resulting from RTP technique show huge anomalies in three parts, but C-A method show two small anomalies. RTP anomalies based on RTP technique are similar to anomalies from C-A method because of normal distribution in Qoja-kandi area. According to correlation between geological particulars and RTP anomalies obtained from C-A method, andesite units host the anomalies in the studied area.

There is an appropriate correlation between the calculated anomalous threshold values and ore thicknesses in total cores. Also, the ratio of the ore length and total core length is related to anomalous threshold, calculated with C-A method. Based on RTP technique, three anomalies (two RTP anomalies were identified in the east and west of the southern part of the area and one anomaly in the northern part). Also, according to the C-A method, two small anomalies are situated in the east of southern part of the prospecting area with very high priority for drilling. Borehole 1 and 2 were drilled in mentioned places and confirmed the results of C-A

- 1 model for increasing the resolution of drilling point determination and decreasing the drilling
- 2 risk.
- 3 Hence study geophysical magnetic anomalies with the C-A method can be a proper way for
- 4 geophysists to find targets with enriched magnetic elements. Also, applying C-A log-log can

- 5 increase the resolution of the drilling point determination and decrease the drilling risk.
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- 30

1 Table 1. RTP classification of magnetic anomalies based on fractal method.

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Class ID	Classes range (nT)	Priority areas for drilling
1	45383 - 47424.2	Very low
2	47424.2 - 49493.7	Low
3	49493.7 – 56493.7	Moderate
4	56493.7 - 63533.1	High
5	63533.1 - 66296	Very high

# 1 Table 2. Log report of boreholes with RTP classification based on fractal method.

		Magnetite		Magnetite	e range	
Borehole Total ID (v)	thickness (m) in total Ore / Total			(m) Priority areas fo		
	core(grades	core	_	_	drilling	
	(m)	greater than		From	То	
		20% Fe total)				
				19.3	25.2	
BH1	136.5	52.4	0.38	60.7	85.2	Very high
				109.4	131.4	
				4	12.2	
BH2	171.2	47.2	0.27	50.2	53.5	Very high
				130.6	166.3	
ВНЗ 151.2	151.0	2 32	0.21	80	102	High
	131.2			112	122	
	10.0	12.5	0.11	44	48	Moderate
BH4	106			81	89.5	
BH5	58.9	0	0	-	-	Very low
BH6	136.5	3	0.02	69	72	Low
				44	47	
BH7	172	14	0.08	61.5	63.5	Moderate
				156	164	
BH8	157	29	0.18	70	90	
				133	142	High

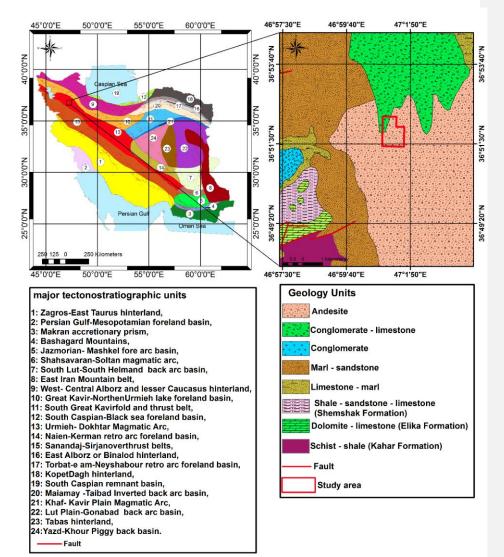
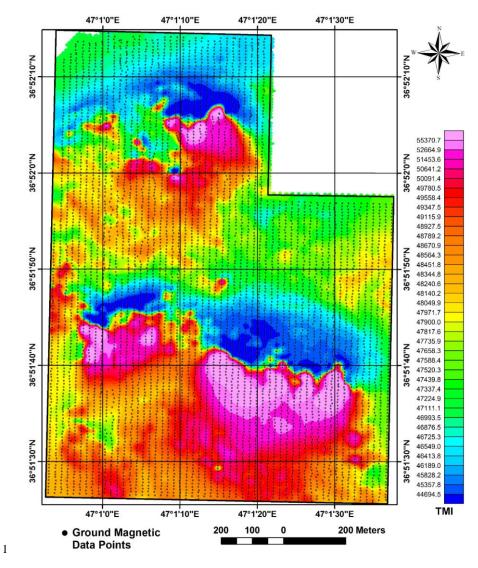






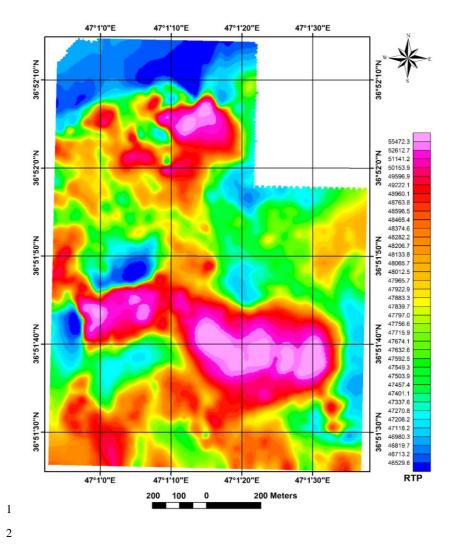
Figure 1. Physiographic-tectonic zoning map of Iran's sedimentary basins (Arian, 2013) and

- 4 location of study area.
- 5

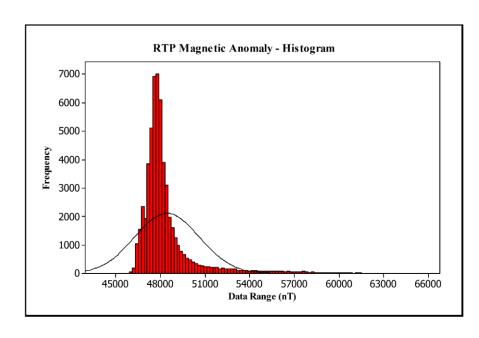




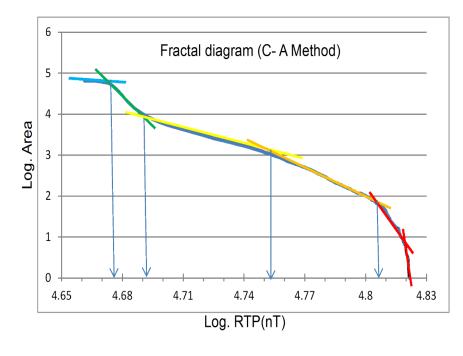
3 Figure 2. TMI map of Qoja-Kandi with ground magnetic data points.



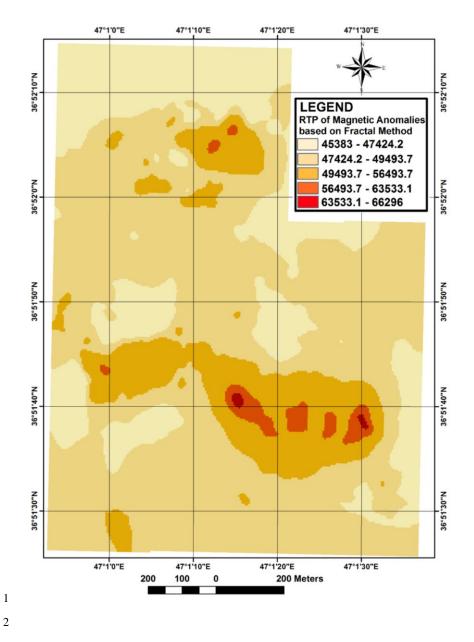
3 Figure 3. RTP map of Qoja-Kandi based on Reduction to the pole technique.



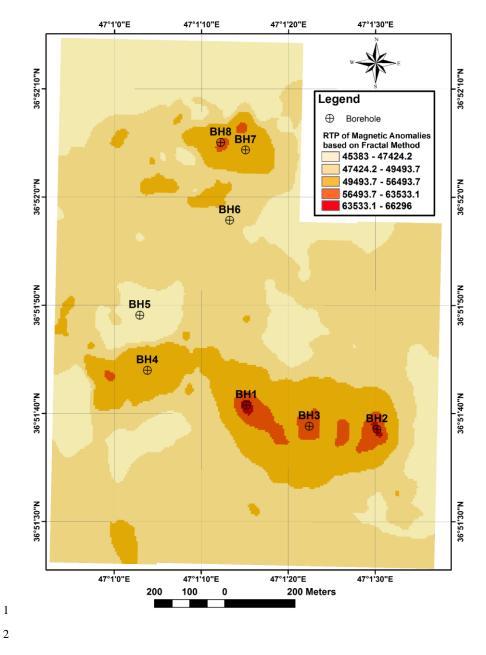
3 Figure 4. Histogram of RTP-MA data in Qoja-Kandi.



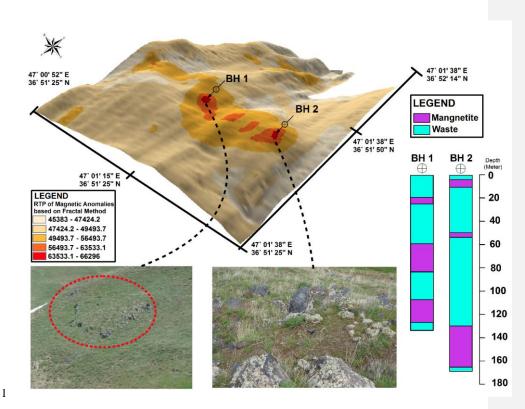
3 Figure 5. Log-log plot for RTP-MA data in Qoja-Kandi.



3 Figure 6. RTP map of Qoja-Kandi based on C-A method.



3 Figure 7. RTP map of Qoja-Kandi based on C-A method with drilled boreholes.



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3 Figure 8. 3D RTP map of Qoja-Kandi based on C-A method with pictures from magnetite

4 zones in the surface of drilled borehole1 and 2, in addition of mentioned boreholes log plots.